

Article

Modelling Hydrological Components of the Rio Maipo of Chile, and Their Prospective Evolution under Climate Change

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Abstract: We used the *Poly-Hydro* model to assess the main hydrological components of the snow-ice melt driven Maipo River in Chile, and glaciers' retreat under climate change therein until 2100. We used field data of ice ablation, ice thickness, weather and hydrological data, and precipitation from TRMM. Snow cover and temperature were taken from MODIS. We forced the model using weather projections until 2100 from three GCMs from the IPCC AR5, under three different radiative concentration pathways (RCPs 2.6, 4.5, 8.5). We investigated trends of precipitation, temperature, and hydrology until 2100 in the projection period (PR, 2014–2100) and the whole period (CM 1980–2100, composite), against historical trends in control period (CP, 1980–2013). We found potentially increasing temperature until 2100, except for Spring (OND). In the PR period, yearly flow decreases significantly under RCP85, on average $-0.25 \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{year}^{-1}$, and down to $-0.48 \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{year}^{-1}$, i.e., $-0.4\% \text{ year}^{-1}$ against CP yearly average ($120 \text{ m}^3 \text{ s}^{-1}$). In the long run (CM) significant flow decrease would, occur under almost all scenarios, confirming persistence of a historical decrease, down to $-0.39 \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{year}^{-1}$ during CM. Large flow decreases are expected under all scenarios in Summer (JFM) during PR, down to $-1.6 \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{year}^{-1}$, or $-1\% \text{ year}^{-1}$ against CP for RCP8.5, due to increase of evapotranspiration in response to higher temperatures. Fall (AMJ) flows would be mostly unchanged, while Winter (JAS) flows would be projected to increase significantly, up to $0.7 \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{year}^{-1}$ during 2014–2100, i.e., $+0.9\% \text{ year}^{-1}$ vs. CP under RCP8.5, due to large melting therein. Spring (OND) flows would decrease largely under RCP8.5, down to $-0.67 \text{ m}^3 \text{ s}^{-1} \cdot \text{year}^{-1}$, or $-0.4\% \text{ year}^{-1}$ vs. CP, again due to evapotranspiration. Glacier down wasting is projected to speed up, and increasingly so with RCPs. Until 2100 ice loss would range from -13% to -49% (-9% , and -39% at 2050) of the estimated volume at 2012, which changed by -24% to -56% (-21% , and -39% at 2050) vs. ice volume in 1982, thus with rapider depletion in the first half of the century. Policy makers will have to cope with modified hydrological cycle in the Maipo River, and greatly decreasing ice cover in the area.

Keywords: climate change; Andes; hydrological modelling; ice down wasting

1. Introduction

Water security in the Andean region is at stake under demographic pressure, the changing climate, and glaciers' shrinking under climate change lately [1–6]. Central Chile is semi-arid with mostly dry Summer, and rain in Winter, and snow and glaciers' melt contributes runoff in the Summer up

to 70–80%. Since the 70s, ~50% of Andean cryosphere wasted under a +0.7 °C increase [7,8], with an initial increase in discharge, but likely a decrease in the long run, also in response to decrease of snowfall [9–11]. Santiago region (with 7 M+ inhabitants) takes most (70–90%) agricultural water from the Maipo River [12], with water use expected to increase along the century [13].

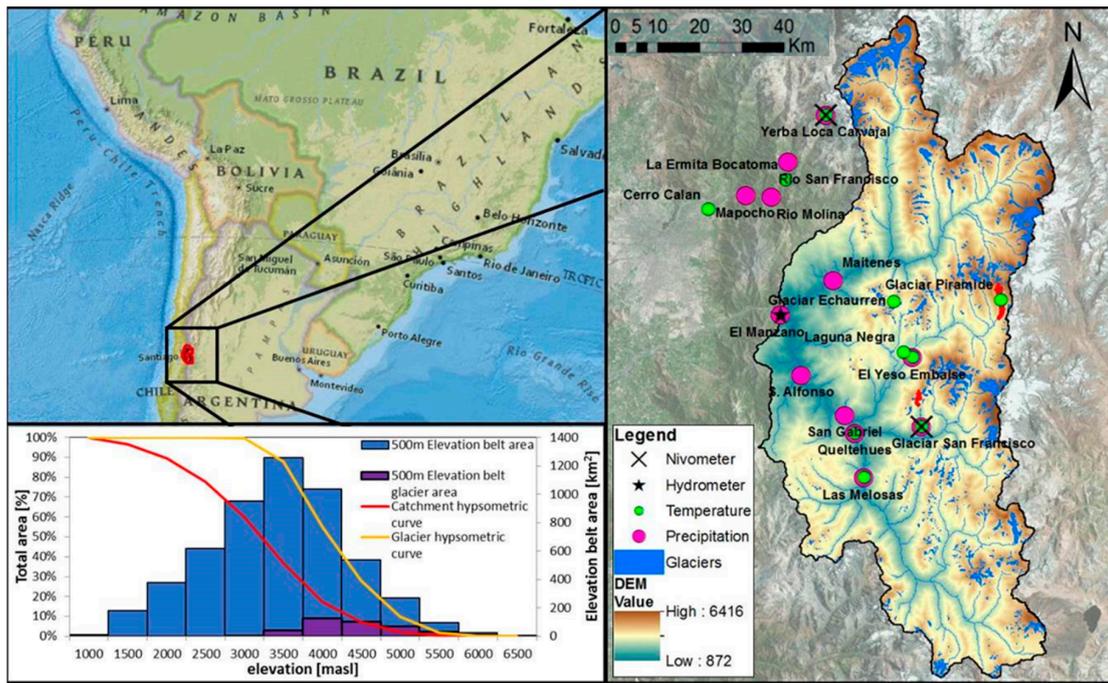
Among others recently, Pellicciotti et al. [6] reviewed recent status of glaciers of Chile. They found glaciers shrinking rapidly in the central Andes, i.e., the Juncal Norte, Juncal Sur, and Olivares Gamma glaciers, the two latter dwelling in the Maipo catchment, wasting to −2.4%, −10.9%, and −8.2% in area since 1955 [1]. Our study here aims at setting up a hydrological model to sketch the present and future hydrological regime, and hydrological components of the basin of Rio Maipo Alto, a high altitude area, where gathering of information is complex given harsh environmental conditions [14–17]. To make up for the sparseness of data at the highest altitudes, we used series of precipitation from the Tropical Rainfall Measurement Mission (TRMM), and temperature from the Moderate Resolution Imaging Spectroradiometer (MODIS). The manuscript structure is as follows. (i) In Section 2 we describe the case study area, data base, the climate projections from the chosen general circulation models GCMs; (ii) in Section 3 we describe hydro-glaciological modelling, projections until 2100, and assessment of future meteo-hydrological trends; (iii) in Section 4 we report the main results of our exercise; (iv) in Section 5 we provide discussion, and benchmarking against recent findings; and (v) in Section 6 we report our conclusions, and outlooks.

The work partly builds on a former study [18] focusing upon the Maipo River. Here we provided much wider insight, namely by (i) projecting future hydrology under two more GCMs, so exploring a wider range of climate uncertainty; (ii) explicitly investigating past, and future trends of glaciers' cover; (iii) explicitly investigating the importance, and potential future evolution of flow components; (iv) explicitly investigating past, and potential future flow trends under the defined variability range; and (v) performing a correlation analysis between in stream flows and climate drivers, for the historical and projected data, to assess changes between present and future flow generating mechanisms, and the climate. Accordingly, our results here are original and of large interest for scientists, and policy makers in the Andean region.

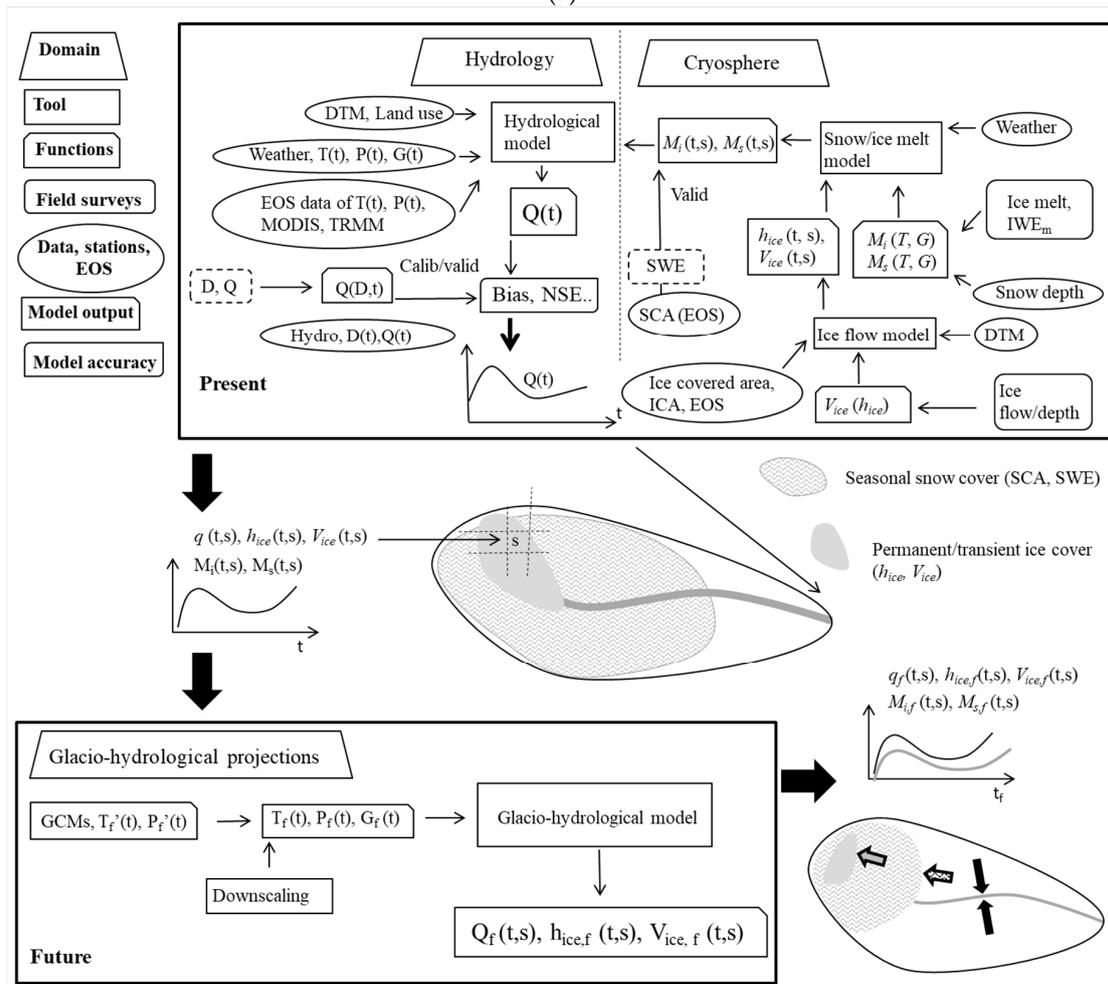
2. Case Study and Data Base

2.1. Maipo River

The case study area was the basin of Maipo River (Central Chile), closed at El Manzano (a small town along the river, ca. 15 km upstream Santiago, 890 m a.s.l., 4839 km², in Figure 1), nesting the upper Maipo, and its main tributaries (Olivares, Colorado, Yeso and Volcàn). The catchment covered 33°4' S to 34°15' S, and altitude varied between 890 m a.s.l. to more than 6500 m a.s.l. of Tupungato. About 90% of the area was above 2000 m a.s.l., and 60% above 3000 m a.s.l. There was 364 km² of glaciers, i.e., 8% of the area. Average flow was 120 m³ s^{−1} (1980–2013).



(a)



(b)

Figure 1. (a) Case study. Rio Maipo. Location of gauging/weather stations, glaciers, hypsometry. Field campaigns were carried out in 2012 on glaciers San Francisco and Pyramid. (b) Flow chart of the modelling approach and components, data requirements, tuning methods, and water resources

projections under climate change. The necessary tools are reported, as per seven categories, i.e., domain (e.g., hydrology, cryosphere), tools (e.g., hydrological model, snow melt model), functions (e.g., snow melt M_s as a function of temperature, and radiation $M_s(T, G)$, etc.), necessary data from field surveys or other sources (ice melt from stakes, ice volume loss from topographic methods, earth observation from space EOS, etc.) data (weather, snow depth, SCA from remote sensing, etc.), model outputs (e.g., ice melt in time and space $M_i(t, s)$), and model accuracy (e.g., Bias, NSE). $T(t)$ is daily temperature, $P(t)$ daily precipitation, $G(t)$ is solar radiation, $D(t)$ daily flow depth at hydro station(s), $Q(t)$ is daily discharge at outlet section. $M_i(t, s)$ is daily ice melt in a given place (cell) s , $M_s(t, s)$ is daily snow melt, $q(t, s)$ is daily runoff in cell s , $h_{ice}(t, s)$ is daily ice depth, and $V_{ice}(t, s)$ daily ice flow velocity. SCA is snow covered area. SWE is snow water equivalent. ICA is ice covered area, IWE_m is water equivalent of ice melt. Bias is systematic error on average, NSE is Nash-Sutcliffe Efficiency. $T_f^f(t)$, $P_f^f(t)$ are (future/projected) temperature and precipitation from GCMs before downscaling (biased); $T_f(t)$, $P_f(t)$ future daily temperature and precipitation after downscaling (unbiased). Dashed lines indicate methods/data that could be used generally as alternatives, but were not used here.

Soil was bare (54%), with bushes (24%), snowfields and glaciers (8%), and elsewhere forest, prairie, and steppe (MODIS, MCD12Q1, 500 m resolution). At the lowest altitudes the climate was temperate Mediterranean, with dry Summer and wet Winter (May to August, 95% of annual precipitation). Summers were hot with 30 °C+, and strong thermal excursion and cold climate at high altitudes (above 2500 m a.s.l.) were seen. Winter temperature was largely below 0 °C and considerable snow falls occur [19]. Rainfall increased from North to South [20]. Two main reservoirs were present in the catchment, i.e., Laguna Negra (600 Mm³) and el Yeso (250 Mm³, on the Rio Yeso River).

2.2. Historical Weather and Hydro Data

Poly-Hydro model needs daily temperature and precipitation. The grid size adopted here was 3 × 3 km², somewhat large given the size of glaciers within the catchment (the largest being Juncal Sur, ca. 21 km²). However, given the size of Maipo catchment, such grid size was a reasonable trade-off between description distributed processes including ice flow, and computational burden for long term simulation of climate change scenarios. Accordingly, input data of precipitation, and temperature were distributed over a 3 × 3 km² grid, and fed to the model. To aid proper spatial interpolation on the model grid, and to capture local patterns in temperature, and precipitation, we complemented the information from ground stations with data from TRMM, and MODIS satellites, reported below in Section 3. In Figure 1b a flow chart is reported, resuming the modelling approach and components, data requirements, tuning methods, and water resources projections under climate change (see also [17] for reference).

We used data from ten temperature stations, seven in the Maipo basin Rio Alto, and three northeast of Santiago (Rio Mapocho, Figure 1, Table 1). There were 11 rainfall stations available (Figure 1, Table 1) below snow line. Data from the station Yerba Loca Carvajal could not be used because the station was not heated, and snow water equivalent was not accounted for systematically. The stations were property of the Dirección General de Aguas DGA (except for one on the Pyramid glacier, property of EvK2CNR), and three were higher than 3000 m a.s.l. Monthly temperature lapse rates were estimated from stations with at least eight years of data, and varied from −6.9 °C km^{−1}, to of −5.2 °C km^{−1}, in February and December. Precipitation increases with altitude, from +3 mm km^{−1} in December (Spring) to +200 mm km^{−1} in July (Winter). The DGA made available daily flows (m³ s^{−1}) at El Manzano (Figure 1). In the Maipo catchment one snow depth station was present near the San Francisco glacier at 2220 m a.s.l. (Figure 1), and we could also use the Yerba Loca Carvajal snow station, close to the divide (3250 m a.s.l.).

Table 1. Weather and hydrometric stations available. Measured variables are temperature T, precipitation P, snow depth HS, discharge Q. See Figure 1 for location.

Station	Altitude (m a.s.l.)	Variable	Period Available
Cerro Calan	848	T	(1975–2013)
El Manzano	890	P, Q	P(2012–2013), Q(1995–2013)
Mapocho	966	P	(2012–2013)
San Alfonso	1040	P	(1965–1973, 2012–2013)
Maitenes Bocatoma	1143	P	(1979–2013)
Rio Molina	1158	P	(2010–2013)
San Gabriel	1266	P	(1977–2013)
Le Ermita Bocatoma	1350	T	(1987–2011)
Queltehues	1450	T, P	T(1987–2011), P(1972–1980)
Las Melosas	1527	T, P	T(1977–1978), P(1962–2006)
Rio San Francisco	1550	P	April–July 2013
Glaciar San Francisco	2220	T, P, HS	T(2012–2013), P(2012), HS(2012–2013)
El Yeso embalse	2475	T, P	T(1963–2013), P(1998–2013)
Laguna Negra	2780	T	T(2012–2013)
Yerba Loca Carvajal	3250	T, P, HS	T(2011–2013), P(2013), HS(2012–2013)
Glaciar Piramide	3587	T	T(March–April 2012)
Glaciar Echaurren	3850	T	T(1999–2001)

2.3. Satellite Data

We used temperature data from MOD11C3 product of MODIS (Moderate Resolution Imaging Spectroradiometer) on board of Terra (EOS AM), during 2002–2012. Rainfall distribution in space was assessed using data from TRMM (Tropical Rainfall Measuring Mission, TRMM 2B31, [21,22]) during 1998–2009 ($4 \times 4 \text{ km}^2$), obtained by fusion of Precipitation Radar (PR) data, and TRMM Microwave Imager data [23]. Simulated snow coverage (i.e., water equivalent SWE) was validated against MOD10A2 product at 500 m resolution (snow covered area), 8 days composite [24–26] during 2010–2012.

2.4. Field Campaigns

Field data were gathered in 2012 by personnel of EvK2CNR in the project “Action plan for the safeguarding of the glaciers against climate change” of DGA, including measurements on San Francisco, and Pyramid glaciers [18]. San Francisco glacier had seven ablation stakes, at 2890–3425 m a.s.l. Pyramid is a debris covered glacier, and supraglacial rock debris strongly affects ice melt [27], so we assessed melt factors against debris thickness [15,16,28].

2.5. Climate Projections

The Fifth Assessment Report (AR5) by the Intergovernmental Panel on Climate Change IPCC presented scenarios describing four different Representative Concentration Pathways (RCPs, 2.6, 4.5, 6.5, 8.5 [29]). Temperature and precipitation were taken according to RCP2.6 (peak in radiative forcing at 3 Wm^{-2} or 490 ppm CO_2 equivalent at 2040, with decline to 2.6 Wm^{-2}), RCP4.5 (stabilization without overshoot pathway to 4.5 Wm^{-2} , or 650 ppm CO_2 eq. at 2070), and RCP8.5 (rising radiative forcing up to 8.5 Wm^{-2} , or 1370 ppm CO_2 eq. by 2100). Three GCM models were used for this study (Table 2), i.e., CCSM4 ([30], <https://www.earthsystemgrid.org>), EC-Earth ([31], <http://eearth.knmi.nl/>) and ECHAM6 ([32], <http://cera-www.dkrz.de>). The low spatial resolution of GCMs requires downscaling to perform hydrological/impact studies at the basin scale [33].

Table 2. Features of the three adopted GCMs.

Model	Institute	Country	Grid Cell Size	Layers	Cells
CCSM4	National Center for Atmospheric Research	U.S.A.	$1.25^\circ \times 1.25^\circ$	26	288×144
ECHAM6	Max Planck Institute for Meteorology	GER	$1.875^\circ \times 1.875^\circ$	47	192×96
EC-EARTH	Europe-wide consortium	E.U.	$1.125^\circ \times 1.125^\circ$	62	320×160

3. Methods

3.1. Glacio-Hydrological Modelling

Poly-Hydro model [17,34] was used here, developed at Politecnico di Milano. The *Poly-Hydro* model can be used to mimic hydrology of high altitude catchments [15,26,35,36]. We do not report here all model's equations, and the reader is referred to the literature above. Glacier flow is modeled as driven by gravity [15]. Ice flow velocity is depicted by a simplified force balance, proportional to shear stress to a power n , i.e., with Glen's flow law ($n = 3$, e.g., [37,38]). Basal shear stress τ_b (Pa) is either known or estimated, and accounting for deformation and sliding velocity as governed by τ_b one can approximate depth averaged ice velocity as [39]:

$$V_{ice,i} = K_d \tau_{b,i}^n h_{ice,i} + K_s \frac{\tau_{b,i}^n}{h_{ice,i}}, \quad (1)$$

with $h_{ice,i}$ (m) ice thickness in the cell i , and K_s ($\text{m}^{-3} \cdot \text{year}^{-1}$) and K_d ($\text{m}^{-1} \cdot \text{year}^{-1}$) parameters of basal sliding and internal deformation. Model tuning was carried out using observed flow velocities for the San Francisco, and Pyramid glaciers, where ice thickness was known via GPR measurements [40]. Initialization was pursued by estimating ice thickness $h_{ice,i}$ for each cell (with ice cover) within our catchment, to subsequently estimate basal shear τ_b as:

$$\tau_{b,i} = \rho_i g h_{ice,i} \sin \alpha_i, \quad (2)$$

with ρ_i ice density (kg m^{-3}), g gravity (9.8 m s^{-2}), and α_i slope. DGA provided estimated ice thickness at 2012 for all glaciers of the Maipo catchments. Avalanching on the glaciers is accounted for by considering the terrain slope (linearly increasing within 30° – 60° , [34]). The cell size used here was $3 \times 3 \text{ km}^2$ as reported, giving a reasonable trade-off between accuracy, and computational burden.

3.2. Temperature and Precipitation Correction Using Satellite Data

We used monthly temperature from ground stations to assess monthly lapse rates. We then evaluated monthly temperatures and lapse rates from MOD11C3 data in the same cells ($3 \times 3 \text{ km}^2$) of the temperatures gages, which gave values acceptably close to those from ground stations ($-5.3 \text{ }^\circ\text{C km}^{-1}$ against $-5.8 \text{ }^\circ\text{C km}^{-1}$ yearly), so we could use it to provide spatial distribution of temperature [41]. For each cell ($3 \times 3 \text{ km}^2$) we estimated average yearly temperature from MOD11C3. Benchmarking temperature from MOD11C3 against ground based values, we obtained a distributed maps of temperature corrections to be applied. To aid assessment of rainfall distribution we used TRMM 2B31 [21]. Monthly average TRMM rainfall was estimated in the same cells ($3 \times 3 \text{ km}^2$) as the rainfall gages, with acceptable agreement below 2500 m a.s.l. We thus decided to use TRMM correction in Summer (January February March, JFM), i.e., in the absence of snow. We calculated mean rainfall from TRMM, and we obtained a distributed map of (multiplicative) correction to be applied to rainfall.

3.3. Snow and Ice Ablation Modelling

Snow melt M_s was estimated using degree day above a threshold T_t , with melt factor D_D ($\text{mm }^\circ\text{C}^{-1} \cdot \text{day}^{-1}$) [14]. Snow melt factors were estimated at Yerba Loca station (year 2013), and at San Francisco glacier (2012). We found $D_D = 5.9 \text{ mm }^\circ\text{C}^{-1} \cdot \text{day}^{-1}$ for Yerba Loca, and $D_D = 6.5 \text{ mm }^\circ\text{C}^{-1} \cdot \text{day}^{-1}$ for San Francisco. A threshold $T_t = -1 \text{ }^\circ\text{C}$ was taken from data analysis. Similarly ice melt M_i was estimated using a degree day with melt factor D_I factor ($\text{mm }^\circ\text{C}^{-1} \cdot \text{day}^{-1}$). Bare ice ablation from seven stakes on the San Francisco glacier gave a constant (with altitude, 2890 to 3425 m a.s.l.) $D_I = 4.1 \text{ mm }^\circ\text{C}^{-1} \cdot \text{day}^{-1}$ [40]. Pyramid glacier was instead debris covered. From ablation data of 2012 at four stakes with debris cover thickness (ca. 5 to 50 cm, [40]) we calculated $D_I = 4.5 \text{ mm }^\circ\text{C}^{-1} \cdot \text{day}^{-1}$ at 10 cm, $D_I = 2.5 \text{ mm }^\circ\text{C}^{-1} \cdot \text{day}^{-1}$ above 25 cm, and $D_I = 2.6 \text{ mm }^\circ\text{C}^{-1} \cdot \text{day}^{-1}$ at 0 cm (bare ice), not very large, given that melt factors may

reach up to $20 \text{ mm } ^\circ\text{C}^{-1}\cdot\text{day}^{-1}$ [42,43]. Debris cover mapping was not available for Maipo River, unless specified locally (e.g., Laguna negra [4]). A rough estimation of debris cover would indicate 45% or so of ice area [44], but accurate assessment requires large effort [45,46]. Being both D_D D_I values site-specific, and possibly not spatially representative, we decided to use them as parameters for model tuning, with the assumption that they would not differ too much from those obtained in the field study (as done e.g., in [26]).

3.4. Downscaling of GCM Projections

We pursued downscaling of the climate scenarios. For precipitation we used stochastic time random cascades (SSRCs) [33], tuned for each GCM using 1994–2003 daily precipitation data at S. Gabriel station, most complete. A constant (multiplicative) term forces the average daily precipitation from the GCM to equate its ground based value. A β (binomial) generator is then used to evaluate the probability of wet (or dry) spells. Finally, a “strictly positive” generator mimics variability of precipitation. The so estimated parameters were then used to disaggregate future precipitation. Temperature downscaling was also carried out, using the 1994–2003 temperature series in Yeso Embalse station. A monthly *Delta-T* approach was used [35]. Precipitation and temperature were also corrected for altitude using lapse rates, and further spatial distribution was modeled as reported in Section 3.2.

3.5. Glacio-Hydrological Projections

Glacio-hydrological projections were carried out by feeding *Poly-Hydro* with the precipitation and temperature scenarios obtained above. We estimated yearly, and seasonal average flows yearly until 2100, and the projected amount of ice (m^3) in every year under each scenario. The initial condition for ice volume was taken from DGA estimates at 2012 as reported in Section 3.1. To further benchmark future ice depletion against past conditions, we back-estimated ice volume until 1982, ever since when climate data were available. To do so, we used trial and error, iteratively increasing (as a percentage) initial ice thickness at 1982, and subsequently simulating ice flow and ablation until 2012 (see for a similar procedure on the Baltoro glacier, in [15]). Also for the purpose of benchmarking, we simulated potential future ice cover under a stationary climate, i.e., feeding the model with simulated climate series displaying the same statistics as now.

3.6. Trend Analysis, and Correlation against Climate Drivers

We performed a trend analysis (using linear regression LR) on temperature (since 1981), precipitation (since 1982), and stream flows (since 1980) in the past period, which we call CP, to verify the presence of any measurable change. We studied yearly, and seasonal trends of temperature (Yeso Embalse), precipitation (San Gabriel), and flow discharge (El Manzano). Also, we studied the climate/hydrologic trends from our scenarios during the projections period PR (2014–2100). To assess trends under potential climate change, we calculated LR trends for the whole considered period, i.e., CP + PR (1980/1/2–2100), named composite CM. Albeit more complex and accurate methods exist to assess non-stationarity of climate series, maybe combining more tests [47,48]; here the purpose was to evaluate the magnitude and significance of trends of climate and water resources, and the bearing of climate on such trends. We then performed a correlation analysis between in stream flows under each scenario, and the corresponding climate drivers of temperature and precipitation, and between these same drivers [48,49]. Also, we carried out the same analysis on historical data, to assess potential similarities, and changes between present and future flow generating mechanisms, and climate drivers.

4. Results

4.1. Models' Performance

In Figure 2 we report calibration of *Poly-Hydro* (1994–2003, 2008–2011) at El Manzano, With Figure 3 giving validation. As reported in Figure 1b, model calibration was pursued by optimization of two goodness of fit measures, namely Bias (systematic error on average), and NSE (Nash-Sutcliffe Efficiency). We manually calibrated the model (see Table 3 for calibration parameters) to obtain a highest value of NSE ($0 < \text{NSE} < 1$ in general), constrained by a small Bias (i.e., $|\text{Bias}| < 5\%$), so that, on average, the model depicted well water resource availability, and it also represents acceptably well their variability in time measured by NSE (see [34] for some discussion about calibration of hydrological models in high altitude environment). In Table 3, we report calibration/validation statistics. Also in Figures 2 and 3 we report daily flow components (precipitation, snow and ice melt). In Table 3 we also report statistics for calibration/validation without correction of ground temperature/rainfall with remote sensing ($\text{NSE}_{\text{NS}}, \text{RMSE}_{\text{NS}}$), and with use of the sole TRMM data ($\text{NSE}_{\text{TR}}, \text{RMSE}_{\text{TR}}$). We also report seasonal validation statistics, representing different flow generating mechanism (i.e., mostly snow melt during OND, ice melt during JFM, mixed during Fall, AMJ and Winter JAS). Generally, *Poly-Hydro* mimics acceptably seasonal flow dynamics. Snow melt factor was $D_D = 5.6 \text{ mm } ^\circ\text{C}^{-1} \cdot \text{day}^{-1}$, close to that from snow data ($D_D = 5.9 \text{ mm } ^\circ\text{C}^{-1} \cdot \text{day}^{-1}$ at Yerba Loca, $D_D = 6.5 \text{ mm } ^\circ\text{C}^{-1} \cdot \text{day}^{-1}$ at San Francisco). The modelled monthly snow cover in Winter and Spring (JAS, OND, 2010–2012) was benchmarked vs. MOD10A2, with acceptable results (not shown). Ice melt factor was set to $D_D = 7.2 \text{ mm } ^\circ\text{C}^{-1} \cdot \text{day}^{-1}$, larger than those observed upon our case study glaciers ($D_I = 4.1 \text{ mm } ^\circ\text{C}^{-1} \cdot \text{day}^{-1}$ on San Francisco, $D_I = 2.6\text{--}4.5 \text{ mm } ^\circ\text{C}^{-1} \cdot \text{day}^{-1}$ on Pyramid, depending on debris cover), and yet in line with values in literature [42,43,50]). In Figure 2 we report the flow contribution (mmd^{-1}) from snow melt q_s , ice melt q_i , and rainfall q_r ($q = q_s + q_i + q_r$), indicating flow generation mechanism in different periods (see [34] for the importance of separating flow generation mechanisms). A large increase occurs during Spring (OND) due to snow melt, and subsequently ice melt increases flows in Summer (JFM), and Fall (AMJ). Rainfall sporadically contributes to flow during Winter (JAS). Yearly average during 1994–2003 from *Poly-Hydro* was $E[q_s] = 1.26 \text{ mm} \cdot \text{day}^{-1}$, $E[q_i] = 0.33 \text{ mm} \cdot \text{day}^{-1}$, $E[q_r] = 0.10 \text{ mm} \cdot \text{day}^{-1}$, i.e., snow melt contributes more than ice melt, and much more than rainfall.

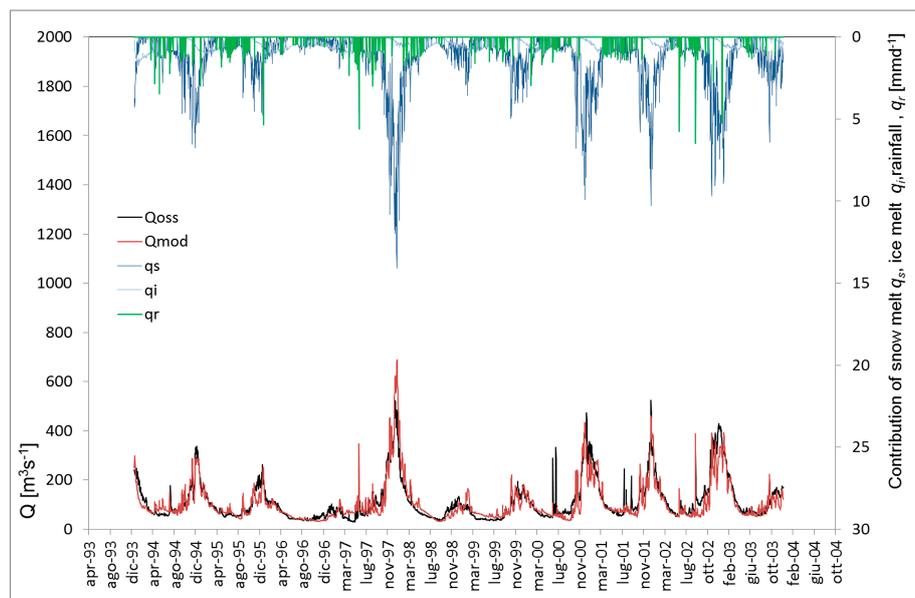


Figure 2. Rio Maipo at El Manzano. Model calibration (1994–2013). Flow contribution (mmd^{-1}), snow melt q_s , ice melt q_i , rainfall q_r .

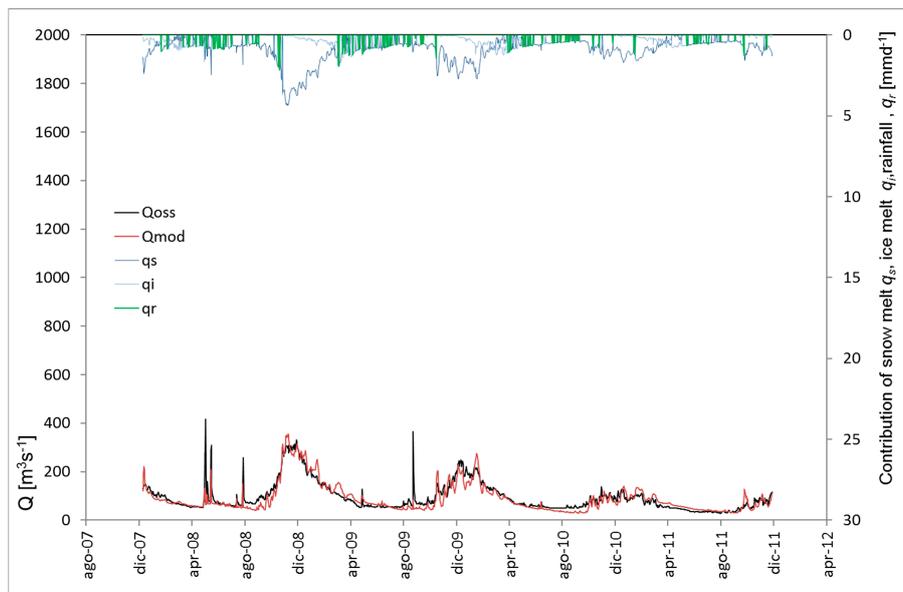


Figure 3. Rio Maipo at El Manzano. Model validation (2008–2011). Flow contribution ($\text{mm}\cdot\text{day}^{-1}$), snow melt q_s , ice melt q_i , rainfall q_r .

Table 3. Poly-Hydro parameters in calibration (1994–2003), and goodness of fit (calibration/validation). Parameter in *italics* are from literature, or defined a priori. Method of estimation explained. Goodness of fit measures reported with method of estimation (e.g., minimization of *Bias*, maximization of *NSE*, etc.). We also report yearly and seasonal flow statistic, observed, and modelled. Seasonal flows reported with confidence limits ($\pm 95\%$) to assess goodness of fit.

Parameter	Description	Value	Method
Calibration			
D_D ($\text{mm}\cdot\text{C}^{-1}\cdot\text{day}^{-1}$)	Snow Degree Day	5.6	Snow data/valid vs. MODIS
D_I ($\text{mm}\cdot\text{C}^{-1}\cdot\text{day}^{-1}$)	Ice Degree Day	7.2	Surveys/Calibration vs. flow
K_d ($\text{m}^{-1}\cdot\text{year}^{-1}$)	Ice flow deformation coefficient	0.98×10^{-16}	Ice stakes (SF, PI)/Literature
K_s ($\text{m}^{-3}\cdot\text{year}^{-1}$)	Ice flow basal sliding coefficient	1×10^{-14}	Ice stakes (SF, PI)/Literature
k (.)	Groundwater flow exponent	2	Max NSE, Min Bias
K ($\text{mm}\cdot\text{day}^{-1}$)	Hydraulic conductivity	4	Max NSE, Min Bias
W_{max} (mm)	Max soil water content (average)	244	Land use analysis
t_s (day)	Lag time surface	3	Max NSE, high flows
t_g (day)	Lag time subsurface	20	Max NSE, low flows
n (.)	Number of reservoir (sup./subsup.)	4/5	Literature
Goodness of fit (Calib., Valid.)			
<i>Bias</i> (%)	Daily average percentage error	−4.4, −4.7	Minimization (for Calib.)
<i>Bias_I</i> (%)	Percentage error ice flow vel.	−4	Minimization (for Calib.)
R^2_I (.)	Det. Coefficient ice flow vel	0.56	Maximization (for Calib.)
<i>NSE</i> (.)	Daily Nash Sutcliffe efficiency	0.81, 0.79	Maximization (for Calib.)
<i>RMSE</i> ($\text{m}^3\cdot\text{s}^{-1}$)	Daily Random mean square error	24.2, 17.2	-
<i>RMSE</i> (%)	Percentage RMSE	23, 19	-
<i>NSE_{NS}</i> (.)	<i>NSE</i> without satellite correction	0.62, 0.61	Maximization (for Calib.)
<i>RMSE_{NS}</i> ($\text{m}^3\cdot\text{s}^{-1}$)	<i>RMSE</i> without satellite correction	35.0, 23.53	-
<i>NSE_{TR}</i> (.)	<i>NSE</i> using only TRMM	0.77, 0.74	Maximization (for Calib.)
<i>RMSE_{TR}</i> ($\text{m}^3\cdot\text{s}^{-1}$)	<i>RMSE</i> using only TRMM	26.5, 19.5	-
Flow statistics obs/mod (Calib., Valid.)			
Q_{avY} ($\text{m}^3\cdot\text{s}^{-1}$)	Av. stream flow yearly ($\pm 95\%$)	113 \pm 17.1/108, 93 \pm 16/89	-
Q_{avJFM} ($\text{m}^3\cdot\text{s}^{-1}$)	Av. stream flow JFM ($\pm 95\%$)	156 \pm 43/155, 121 \pm 21/125	-
Q_{avAMJ} ($\text{m}^3\cdot\text{s}^{-1}$)	Av. stream flow AMJ ($\pm 95\%$)	65 \pm 8/73, 64 \pm 10/66	-

Table 3. Cont.

Parameter	Description	Value	Method
Q_{avJAS} ($m^3 s^{-1}$)	Av. stream flow JAS ($\pm 95\%$)	$68 \pm 12/62$, $58 \pm 10/46$	-
Q_{avOND} ($m^3 s^{-1}$)	Av. stream flow OND ($\pm 95\%$)	$163 \pm 40/142$, $130 \pm 30/119$	-
Q_{av} ($m^3 s^{-1}$)	Average flow discharge ($\pm 95\%$)	$113 \pm 3/108$, $94 \pm 3/89$	Best fitting (for Calib.)
σ_Q ($m^3 s^{-1}$)	Standard deviation of flow discharge	$84/83, 59/60$	-
CV_Q (.)	Coeff. of variation of flow discharge	$0.75/0.76$	-

4.2. Ice Flow Model

Poly-Hydro reproduces acceptably well the velocity field for both glaciers (see [18]). In Table 3, goodness of fit is reported. A *Bias* = -4% was observed, with determination coefficient $R_I^2 = 0.56$. Given this seemingly acceptable performance, we could use our model for description of ice flow of all glaciers in the area, for the purpose of hydrological modelling, and projections.

4.3. Climate and Hydrological Projections

In Figure 4a, we report projected yearly temperature until 2100, as from our three GCMs, and three RCPs (e.g., nine projections, plus mean value, and extremes), together with recent (CP 1981–2013) patterns, or control period CP. Also, yearly precipitation ($mm \cdot day^{-1}$) projections are reported in Figure 4b, including CP (1982–2013). Visibly, yearly temperature was rapidly increasing during CP, and until the end of the century (PR, 2014–2100), and similarly in all seasons (not shown). Precipitation instead was not visibly changing at the yearly scale, nor seasonally (not shown). Seasonal trends of temperature and precipitation under all scenarios were investigated in depth, reported in Section 4.5 below. Seasonal hydrological projections are reported in Figure 5, together with their CP (1980–2013) counterpart. During Summer (JFM, Figure 5a) decreases were seen in both CP, and PR under all scenarios. Fall (AMJ, Figure 5b), and Winter (JAS, Figure 5c) displayed visible decreases in CP, and increases in PR, and Spring (OND, Figure 5d) showed decreases in CP, and substantial stationarity in PR, on average. Again, seasonal trends of stream flows under all scenarios were investigated in depth, reported in Section 4.5 below.

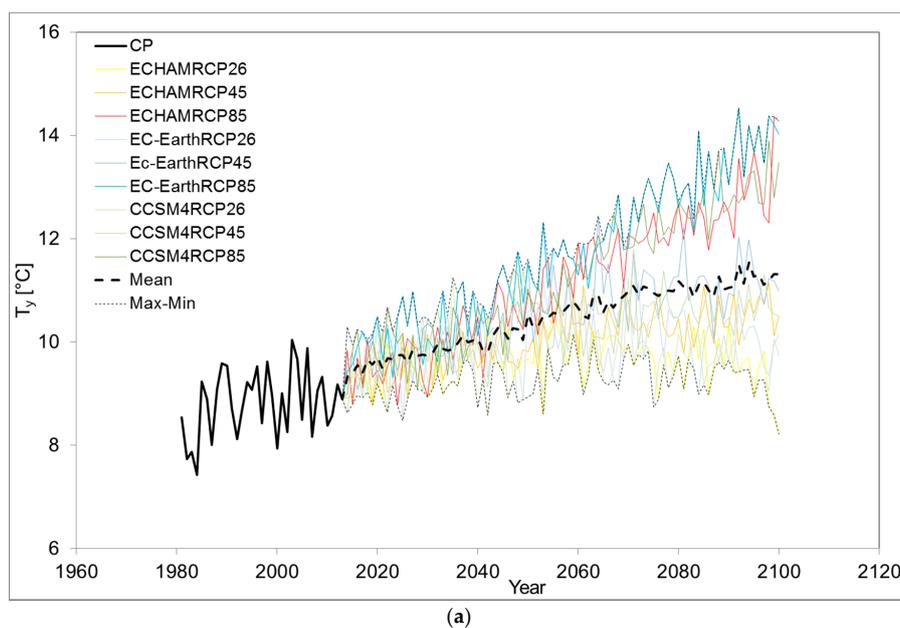


Figure 4. Cont.

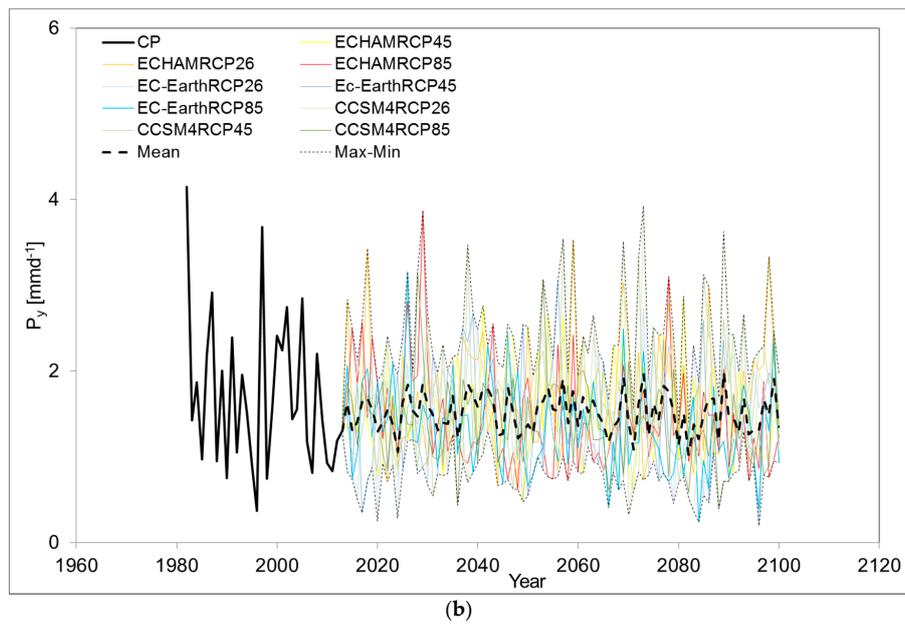


Figure 4. Yearly Climate projections until 2100 vs. CP (1980–2014). (a) Temperature at Yeso Embalse. (b) Precipitation at San Gabriel.

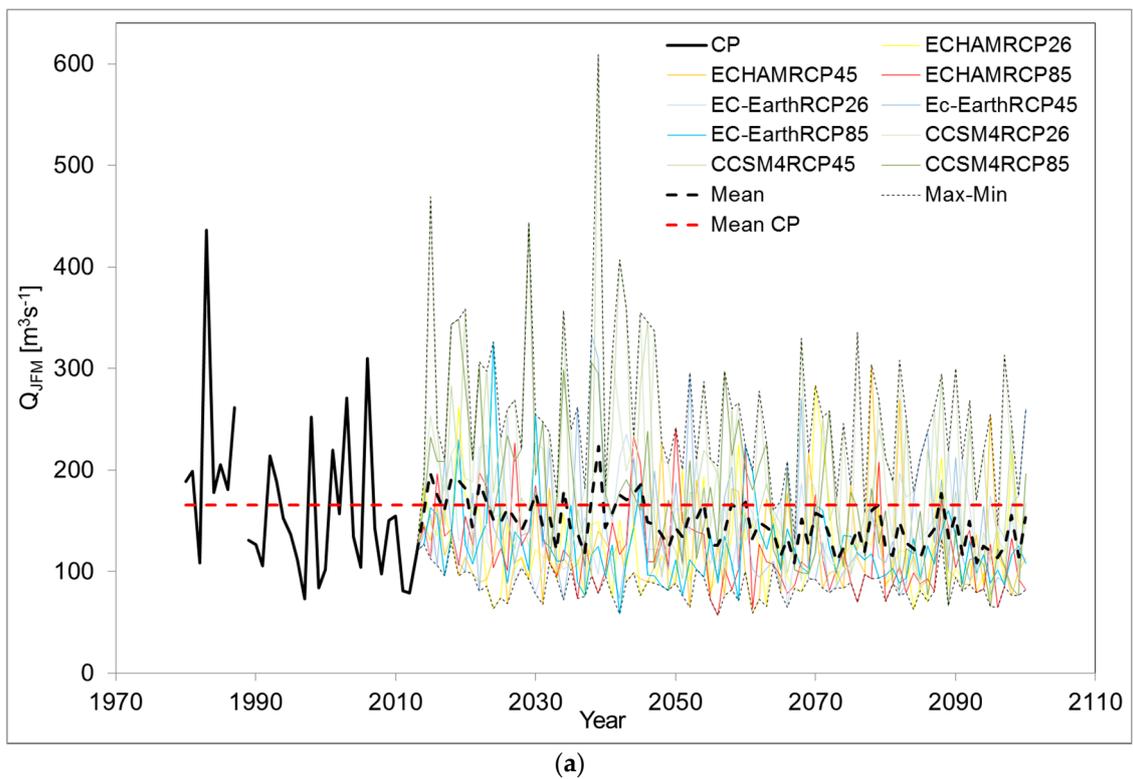


Figure 5. Cont.

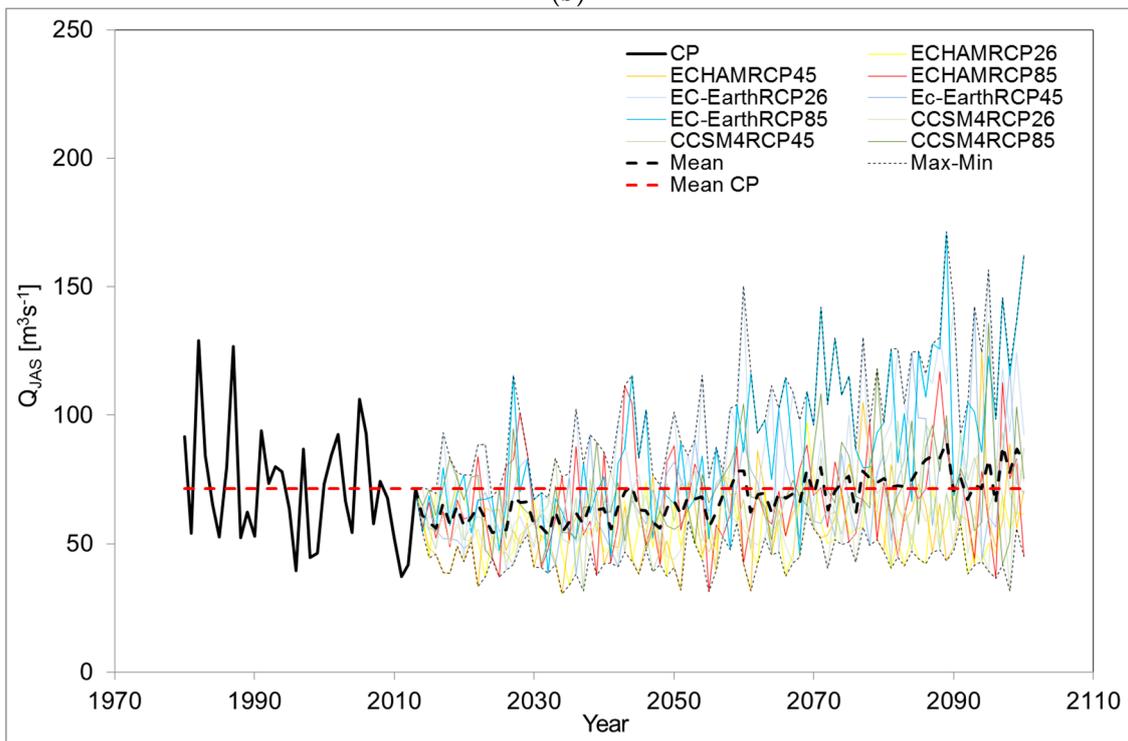
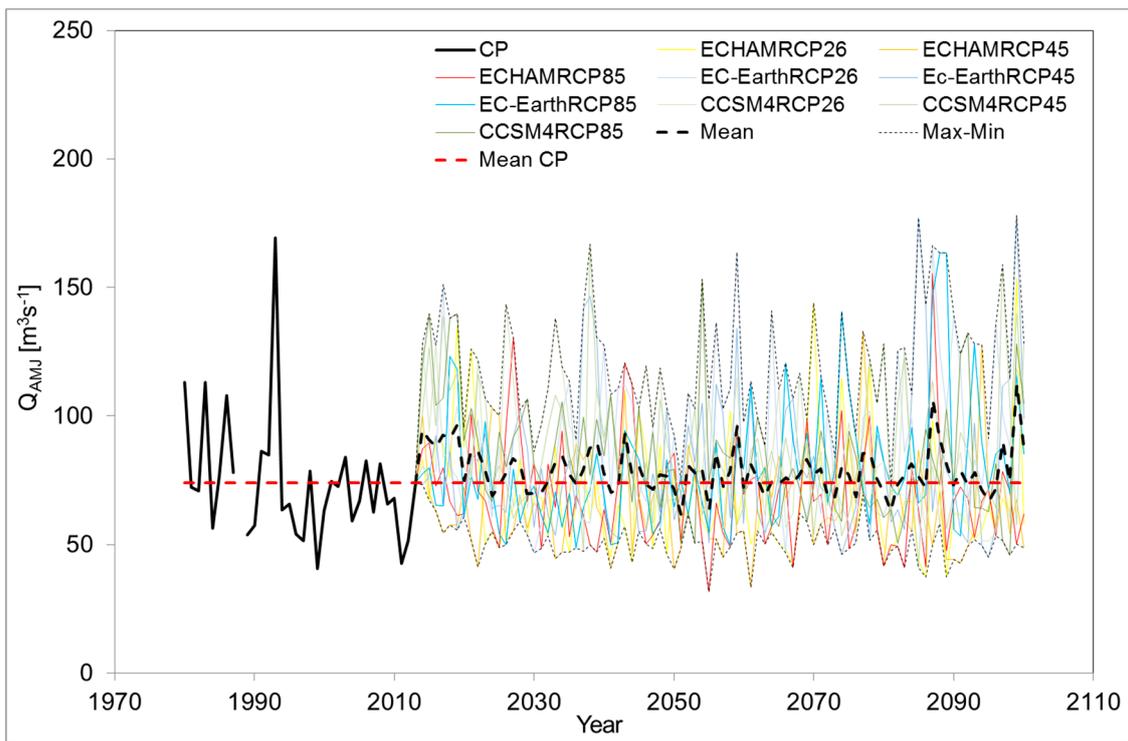


Figure 5. Cont.

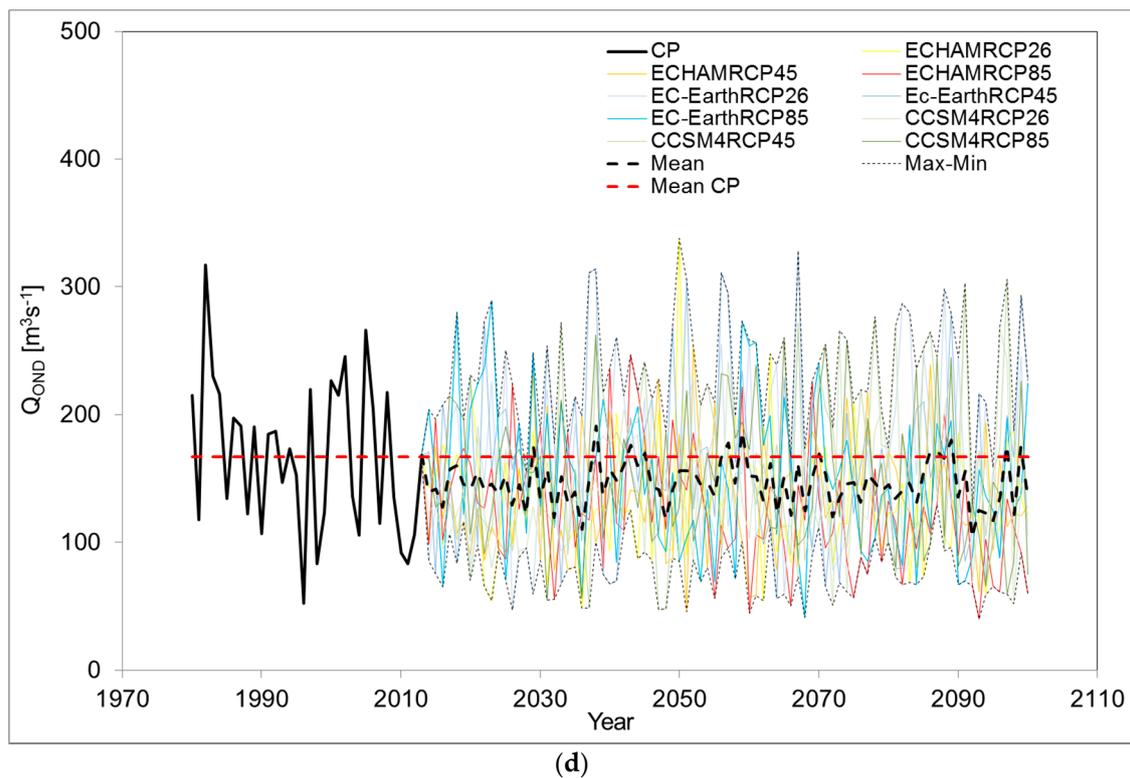


Figure 5. Rio Maipo at El Manzano. Stream flow projections until 2100 vs. CP (1980–2014). (a) Summer JFM. (b) Fall AMJ. (c) Winter JAS. (d) Spring OND.

4.4. Glaciers' Dynamics

Figure 6 reports the past (1982–2013, CP) estimated glaciers' depletion (i.e., ice volume) curve as obtained by back-estimation reported in Section 3.5, against potential future depletion under our climate scenarios. Also, potential ice evolution under present climate (stationary, STAT, is reported). Clearly after 2014, all GCMs models and RCPs projected a sudden increase of ice melting (i.e., a sudden loss in ice volume, as seen by the steeper curves in Figure 6). This was consistent with the large increase of temperature as simulated by all models (Figure 4a), and especially under RCP8.5. Particularly during Summer JFM, all models, and especially CCSM4, depicted large increases suddenly from 2014. Notice that all the GCMs used here provided projected series starting from 2006 (i.e., control runs end in 2005, and projections start in 2006, and no simulation is available starting after that year). Accordingly, projections from these GCMs may provide a large increase than actually happened in some regions, including here (i.e., making temperature in 2014 suddenly much higher than in 2013). As a result, in Figure 6 an abrupt transition is seen from the present behaviour (until 2013) to future dynamics (since 2014). However, this discontinuity only affects the curve locally, and does not impact on the long run glaciers' dynamics. Indeed, glaciers' down wasting as shown here is in response to increase of temperature in all seasons, especially Spring (OND), and Summer (JFM), which results into large ablation at higher altitudes than hitherto.

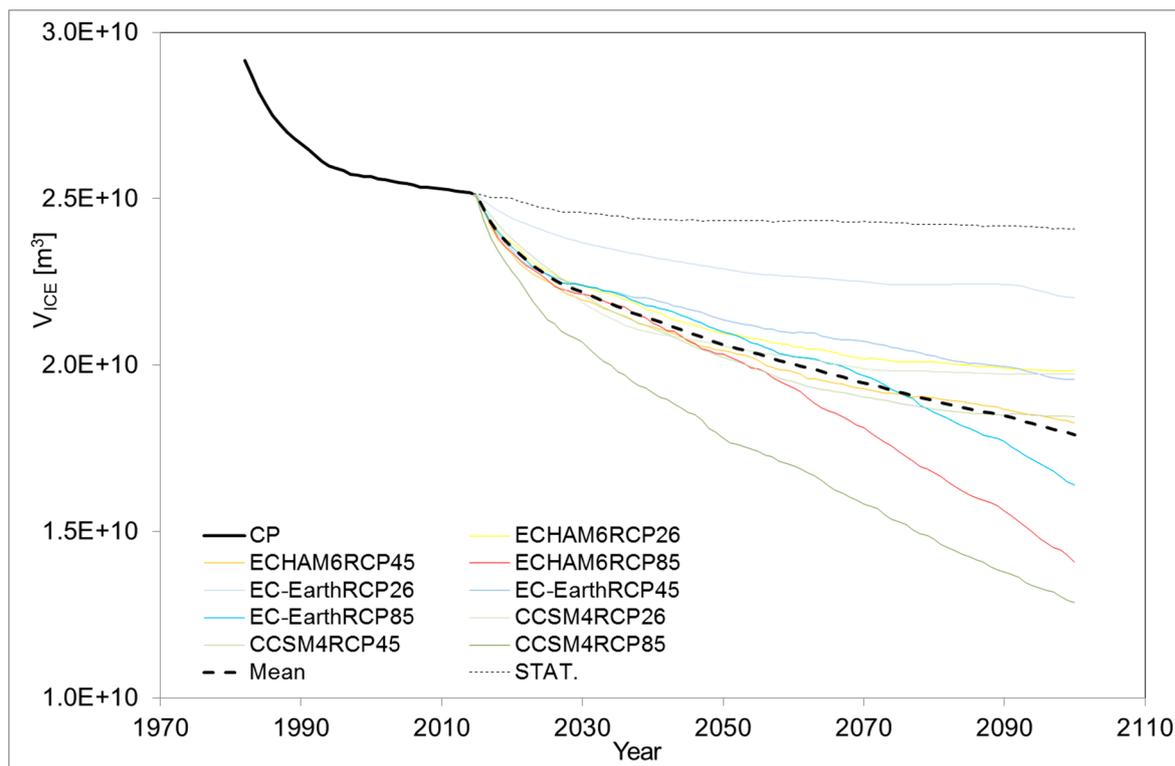


Figure 6. Projected ice volume until 2100 vs. simulated ice volume during CP. Also, projected ice volume is reported under stationary (STAT) climate conditions, for reference.

4.5. Climate and Hydrological Trends until 2100

We quantified climate and hydrological trends via linear regression (LR) during three periods, namely CP (1980–2013, for discharge, temperature, and precipitation), PR (2014–2100), and CM (CP+PR). In Table 4 we report seasonal indicators, including slope, and significance ($\alpha = 5\%$, bold values). In Figure 7 we resumed the so estimated trends of climate and hydrological fluxes for CP and CM situation, the latter of interest to gather a global (i.e., on a longer run) trend assessment, as per RCPs (average values and extremes), regardless of their significance. Also, the projected trends in PR period are reported in Figure 7. Temperature in Table 4 increased during CP (significantly in Summer JFM, and Fall AMJ), with a slight decrease during Spring OND. According to our nine scenarios (3 GCMs, three RCPs) T normally increased (mostly significantly) during all seasons. Decrease (significant) would be seen in few cases, in Summer and Fall under RCP2.6 of CCSM4. RCP2.6 depicts increase of temperature first (since 2014, until half century), then followed optimistically by mitigation of global warming, with decreasing temperature, as seen in Figure 4a. Decrease (not significant) occurred for RCP2.6 of EC-Earth during JAS. RCP4.5 of CCSM4 depicts significant decrease in Fall AMJ. When considering CM period, all scenarios depict significantly increasing temperature. Precipitation in CP decreased yearly, and during Winter JAS, and Spring AMJ, but never significantly. During PR and CM, precipitation was occasionally significantly increasing or decreasing, but no clear patterns were visible. Under RCP8.5, total (yearly), and Winter JAS precipitation decreased significantly. At the Y scale, decrease of flows ($-1.5 \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{year}^{-1}$) was observed in CP (1980–2013), and was statistically significant. From Figure 7, all models and scenarios (not RCP2.6 for ECHAM6) provided long run (CM, 1980–2100) decreasing yearly flows, and several significantly (all RCPs under CCSM4, RCP8.5 under Ec-Earth), with largest decrease projected at $-0.39 \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{year}^{-1}$ until 2100 under CCSM4, RCP8.5. This in spite of RCP2.6, and RCP4.5 projecting slightly increasing Q during PR (but RCP8.5 projecting large decrease on average, significant for CCSM4, and EC-Earth, and not significant increase for ECHAM6, Table 4). The largest flow changes would be during Summer, JFM. During CP significant

flow decrease was detected (down to $-2.6 \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{year}^{-1}$), and CM scenarios project further decreasing discharge until 2100, mostly significantly (Figure 7, Table 4), and similarly during PR period (Table 4), unless for the case of CCSM4 (RCP4.5). During Fall AMJ, CP data provide not significant decrease, and all scenarios depict no changes therein on average. However, ECHAM for RCP2.6, and RCP8.5 projects significant increase during PR period (and also during CM), while EC-Earth under RCP2.6, and RCP8.5 project significant decrease during CP (but not CM), from Table 4. In Winter JAS, no significant decrease was detected in CP. During PR all scenarios project large flow increase in Winter (except for EC-Earth under RCP2.6, with very slight decrease), in most cases significantly (CCSM4 in RCP4.5, ECHAM6 all RCPs, EC-Earth under RCP4.5 and RCP8.5). Given however the observed decrease in CP, on the long run of CM, more stable JAS flows were projected (with significant decrease under CCSM4 for RCP2.6, and significant increase for ECHAM6 under RCP2.6, and RCP8.5). During Spring OND, a somewhat large (albeit not significant, $p\text{-val} = 0.10$) flow decrease was seen in the CP period. During PR, discharges increased on average under RCPs 2.6 and 4.5, but decreased for RCP8.5 (significantly for CCSM4 and ECHAM6). However, on the long run Spring flow always decreased until the end of the century (and significantly in most cases, e.g., all RCPs for CCSM4, RCP8.5 for ECHAM6, and EC-Earth).

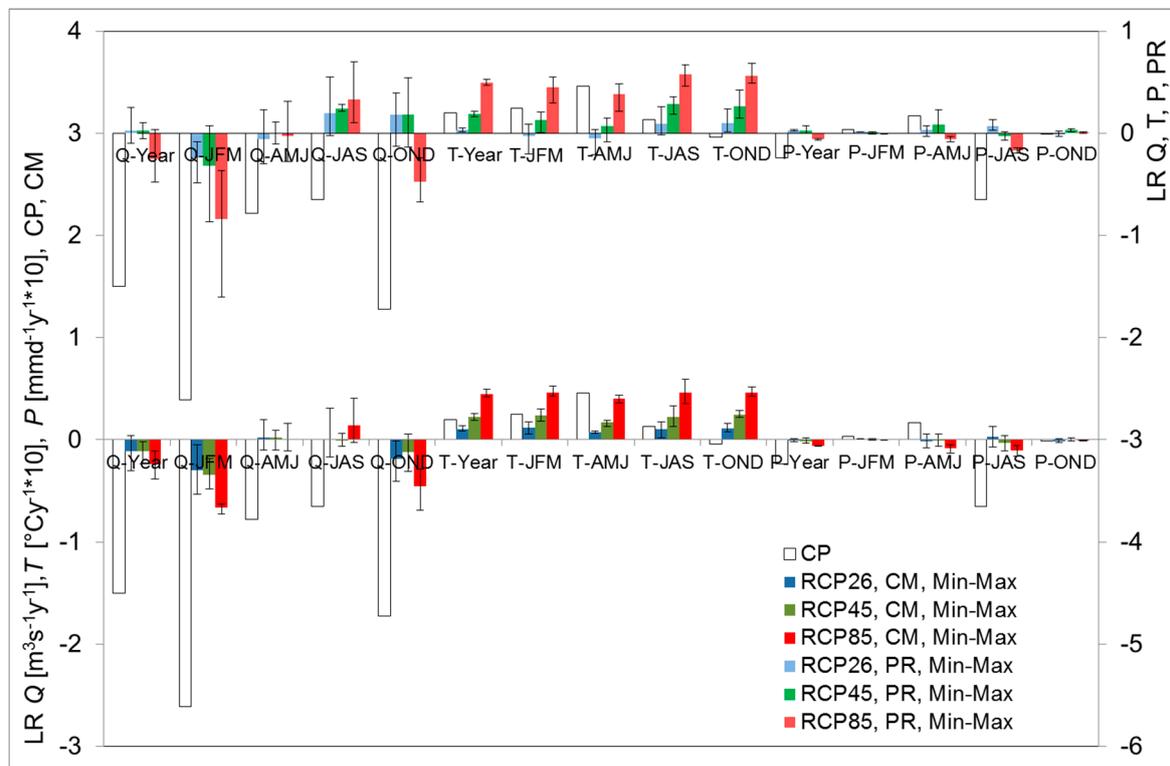


Figure 7. Projected trends of discharge yearly and seasonal discharge Q, precipitation P, temperature T until 2100 as per RCPs vs. CP period. CM period (bars below), and PR period (bars above). Average RCP values reported with maximum and minimum (i.e., among GCMs).

Table 4. Trends of climate and hydrology for CP, PR, and CM periods. In CP periods, average values are reported. In **bold** significant values ($\alpha = 5\%$).

	Season	P, CP	E[P] mm · day ⁻¹	T, CP	E[T] °C	Q, CP	E[Q] m ³ s ⁻¹
CP	Year	-2.4×10^{-2}	1.70	2.0×10^{-2}	8.8	-1.5	120
CP	JFM	3.3×10^{-3}	0.15	2.5×10^{-2}	14.7	-2.6	167
CP	AMJ	1.7×10^{-2}	3.01	4.6×10^{-2}	6.7	-7.8×10^{-1}	74
CP	JAS	-6.5×10^{-2}	2.93	1.3×10^{-2}	3.3	-6.5×10^{-1}	71
CP	OND	-9.3×10^{-4}	0.48	-3.8×10^{-3}	10.7	-1.7E	167

Table 4. Cont.

Scenario	Season	P, PR	P, CM	T, PR	T, CM	Q, PR	Q, CM
CCSM4RCP26	Year	2.2×10^{-3}	-5.0×10^{-4}	2.2×10^{-3}	8.7×10^{-3}	-9.9×10^{-2}	-3.0×10^{-1}
CCSM4RCP26	JFM	1.6×10^{-3}	1.3×10^{-3}	-2.0×10^{-2}	1.7×10^{-2}	-1.4×10^{-1}	-5.3×10^{-1}
CCSM4RCP26	AMJ	4.1×10^{-3}	-2.4×10^{-3}	-2.1×10^{-2}	7.6×10^{-3}	-1.1×10^{-1}	-1.0×10^{-1}
CCSM4RCP26	JAS	3.5×10^{-3}	4.2×10^{-3}	2.6×10^{-2}	2.0×10^{-3}	-2.1×10^{-2}	-1.7×10^{-1}
CCSM4RCP26	OND	-1.1×10^{-3}	-2.3×10^{-3}	2.4×10^{-2}	8.0×10^{-3}	-1.3×10^{-1}	-4.1×10^{-1}
CCSM4RCP45	Year	6.9×10^{-3}	1.7×10^{-3}	1.7×10^{-2}	2.1×10^{-2}	1.0×10^{-1}	-2.4×10^{-1}
CCSM4RCP45	JFM	-1.2×10^{-4}	4.6×10^{-4}	3.5×10^{-4}	3.0×10^{-2}	7.0×10^{-2}	-4.8×10^{-1}
CCSM4RCP45	AMJ	2.3×10^{-2}	5.6×10^{-3}	-8.3×10^{-3}	1.9×10^{-2}	-2.5×10^{-2}	-9.8×10^{-2}
CCSM4RCP45	JAS	1.1×10^{-3}	3.9×10^{-3}	3.5×10^{-2}	1.3×10^{-2}	2.2×10^{-1}	-6.4×10^{-2}
CCSM4RCP45	OND	3.9×10^{-3}	-4.1×10^{-5}	4.2×10^{-2}	2.2×10^{-2}	1.4×10^{-1}	-3.1×10^{-1}
CCSM4RCP85	Year	-5.5×10^{-3}	-5.4×10^{-3}	4.7×10^{-2}	4.3×10^{-2}	-3.1×10^{-1}	-3.9×10^{-1}
CCSM4RCP85	JFM	-2.5×10^{-5}	3.8×10^{-4}	3.0×10^{-2}	5.2×10^{-2}	-5.5×10^{-1}	-7.3×10^{-1}
CCSM4RCP85	AMJ	-3.0×10^{-3}	-8.0×10^{-3}	2.1×10^{-2}	4.0×10^{-2}	-1.2×10^{-1}	-1.1×10^{-1}
CCSM4RCP85	JAS	-2.0×10^{-2}	-9.7×10^{-3}	6.7×10^{-2}	3.5×10^{-2}	1.0×10^{-1}	-2.3×10^{-2}
CCSM4RCP85	OND	1.3×10^{-3}	-1.0×10^{-3}	6.9×10^{-2}	4.2×10^{-2}	-6.7×10^{-1}	-6.9×10^{-1}
ECHAMRCP26	Year	3.3×10^{-3}	8.7×10^{-4}	1.4×10^{-3}	8.5×10^{-3}	2.5×10^{-1}	4.3×10^{-2}
ECHAMRCP26	JFM	1.0×10^{-3}	1.8×10^{-4}	3.8×10^{-3}	5.8×10^{-3}	-8.2×10^{-2}	-3.3×10^{-1}
ECHAMRCP26	AMJ	-2.9×10^{-3}	-7.6×10^{-3}	1.9×10^{-3}	6.3×10^{-3}	2.3×10^{-1}	2.0×10^{-1}
ECHAMRCP26	JAS	1.3×10^{-2}	1.3×10^{-2}	-1.5×10^{-3}	1.1×10^{-2}	5.5×10^{-1}	3.1×10^{-1}
ECHAMRCP26	OND	1.7×10^{-3}	1.0×10^{-3}	1.3×10^{-3}	1.0×10^{-2}	2.9×10^{-1}	-1.3×10^{-2}
ECHAMRCP45	Year	-2.0×10^{-4}	-2.4×10^{-3}	1.6×10^{-2}	1.9×10^{-2}	2.4×10^{-2}	-7.5×10^{-2}
ECHAMRCP45	JFM	-9.4×10^{-4}	-4.7×10^{-4}	1.7×10^{-2}	1.8×10^{-2}	-1.6×10^{-1}	-3.7×10^{-1}
ECHAMRCP45	AMJ	3.0×10^{-3}	-6.6×10^{-3}	1.5×10^{-2}	1.3×10^{-2}	1.1×10^{-1}	9.4×10^{-2}
ECHAMRCP45	JAS	-6.8×10^{-3}	-1.4×10^{-3}	1.8×10^{-2}	2.2×10^{-2}	2.8×10^{-1}	6.6×10^{-2}
ECHAMRCP45	OND	4.4×10^{-3}	2.0×10^{-3}	1.5×10^{-2}	2.2×10^{-2}	-1.4×10^{-1}	-1.0×10^{-1}
ECHAMRCP85	Year	-6.6×10^{-3}	-5.6×10^{-3}	4.9×10^{-2}	4.2×10^{-2}	3.5×10^{-2}	-1.1×10^{-1}
ECHAMRCP85	JFM	-9.2×10^{-4}	-2.5×10^{-4}	5.5×10^{-2}	4.4×10^{-2}	-3.7×10^{-1}	-6.3×10^{-1}
ECHAMRCP85	AMJ	-8.2×10^{-3}	-1.3×10^{-2}	4.6×10^{-2}	3.6×10^{-2}	3.1×10^{-1}	1.6×10^{-1}
ECHAMRCP85	JAS	-1.7×10^{-2}	-5.9×10^{-3}	4.6×10^{-2}	4.4×10^{-2}	7.0×10^{-1}	4.0×10^{-1}
ECHAMRCP85	OND	-3.1×10^{-4}	-3.7×10^{-4}	4.9×10^{-2}	4.5×10^{-2}	-5.1×10^{-1}	-3.9×10^{-1}
EC-EarthRCP26	Year	2.3×10^{-3}	-1.7×10^{-3}	5.3×10^{-3}	1.4×10^{-2}	-8.6×10^{-2}	-7.9×10^{-2}
EC-EarthRCP26	JFM	1.3×10^{-3}	9.7×10^{-4}	9.1×10^{-3}	1.2×10^{-2}	-4.9×10^{-1}	-4.6×10^{-2}
EC-EarthRCP26	AMJ	7.5×10^{-3}	5.2×10^{-3}	3.9×10^{-3}	8.7×10^{-3}	-3.0×10^{-1}	-2.2×10^{-2}
EC-EarthRCP26	JAS	2.7×10^{-3}	-7.5×10^{-3}	4.1×10^{-3}	1.7×10^{-2}	5.4×10^{-2}	-1.2×10^{-1}
EC-EarthRCP26	OND	-3.4×10^{-3}	-2.6×10^{-3}	4.5×10^{-3}	1.6×10^{-2}	3.9×10^{-1}	-1.3×10^{-1}
Ec-EarthRCP45	Year	-4.0×10^{-4}	-3.4×10^{-3}	2.2×10^{-2}	2.6×10^{-2}	-5.4×10^{-2}	-1.6×10^{-2}
Ec-EarthRCP45	JFM	1.2×10^{-3}	8.3×10^{-4}	2.0×10^{-2}	2.3×10^{-2}	-8.7×10^{-1}	-1.8×10^{-1}
Ec-EarthRCP45	AMJ	-3.4×10^{-4}	3.9×10^{-4}	1.5×10^{-2}	1.9×10^{-2}	-1.0×10^{-1}	6.7×10^{-2}
Ec-EarthRCP45	JAS	-3.6×10^{-3}	-1.1×10^{-2}	3.1×10^{-2}	3.3×10^{-2}	2.1×10^{-1}	-1.2×10^{-2}
Ec-EarthRCP45	OND	9.3×10^{-4}	-9.0×10^{-4}	2.1×10^{-2}	2.8×10^{-2}	5.4×10^{-1}	5.6×10^{-2}
EC-EarthRCP85	Year	-5.2×10^{-3}	-6.1×10^{-3}	5.3×10^{-2}	4.9×10^{-2}	-4.8×10^{-1}	-2.2×10^{-1}
EC-EarthRCP85	JFM	-4.0×10^{-4}	4.0×10^{-4}	5.1×10^{-2}	4.3×10^{-2}	-1.6	-6.3×10^{-1}
EC-EarthRCP85	AMJ	-6.4×10^{-3}	-5.1×10^{-3}	4.8×10^{-2}	4.3×10^{-2}	-2.8×10^{-1}	-2.4×10^{-2}
EC-EarthRCP85	JAS	-1.4×10^{-2}	-1.6×10^{-2}	6.1×10^{-2}	5.9×10^{-2}	2.0×10^{-1}	5.0×10^{-2}
EC-EarthRCP85	OND	7.0×10^{-4}	-8.4×10^{-4}	5.1×10^{-2}	5.2×10^{-2}	-2.4×10^{-1}	-2.9×10^{-1}

4.6. Correlation against Climate Drivers

In Table 5, we report the yearly and seasonal correlation analysis between climate (temperature T, precipitation P) and hydrological (discharge Q) variables during past (1982–2013 for all series for consistency), and projected (PR 2014–2100) periods. Only first order correlation analysis was pursued (i.e., same year, same season) for simplicity. In Figure 8 we provide a resume for CP, and PR, as per RCPs, with indication of significance. In grey shades we report correlation analysis during model calibration period (1994–2003) to test consistence of climate to hydrology correlation in modelled and observed series, i.e., to (i) verify if modelled series display a correlation with climate similar to the observations during the same periods, necessary for making usable inferences for the future, and (ii) avoid correlation highlighted by the model which has no correspondence in the actual flow regime. This exercise indicated that the modelled discharges substantially displayed coherent correlation with the observed ones (albeit with less significance in some cases), so that one can use, for indicative purposes, correlation analysis upon modelled variables to investigate potential the future correlation of in stream flows against climate drivers. During CP, normally anti-correlation between flows and temperature was seen, and it was especially significant in Fall AMJ, and Winter JAS. Conversely, precipitation was positively correlated with stream flows, especially in Winter. Also, clear anti-correlation was seen between precipitation, and temperature, yearly, but especially in Winter

(when ca. 45% of yearly precipitation occurs, Table 4), and Spring, however dry. Accordingly, hot Winters may lead to low precipitation income. In the PR period, mostly the observed correlations were confirmed. Precipitation will be largely (and mostly significantly) anti-correlated against temperature during the wettest Winter season, and also yearly, especially under RCP8.5. Consistently, as reported above, yearly, and Winter precipitation would always decrease significantly under RCP8.5. Also, projected flows would be mostly positively correlated to precipitation, especially during Fall AMJ, and less during Winter under RCP8.5 (however variable). A small and not significant correlation, Q vs. P , would be seen during Summer, and Spring, however very dry as reported, was not projected to change in the future. More variable seems the correlation of stream flows against temperature in the future. While at the yearly scale stream flow remained negatively (and significantly on average) correlated against temperature as seen during CP, seasonally some changes may occur. During Summer JFM, RCP26, and RCP4.5 projected increasingly positive correlations of Q vs. T , according to what was found during CP. However, under RCP8.5, anti-correlation was projected. Similarly occurs during Fall AMJ, when small positive correlation coefficients were projected on average, against the significantly negative correlation during CP. Also in Winter, large and often significant (see Table 5) positive correlation coefficients are found between Q and T , increasingly with the RCP, and conversely to CP. During Spring OND, mostly not significant correlation was projected, except for some anti-correlation under RCP8.5. Accordingly, the projected correlation structure between flow and its climate drivers in the Maipo River mirrors partially the results found investigating the historical series, and some changes were seen. Such modified correlation structure may derive from the underlying patterns of flow formation, as driven by climate, and requires some discussion, which we report below.

Table 5. Yearly, and seasonal correlation analysis of climate (temperature T , precipitation P), and hydrological (discharge Q) variables, CP, and PR periods (RCP2.6–8.5). In **bold** significant values ($\alpha = 5\%$).

	CP	PR, CCSM4			PR, EC-Earth			PR, ECHAM6		
		RCP26	RCP45	RCP85	RCP26	RCP45	RCP85	RCP26	RCP45	RCP85
Q-T Y	-0.32	-0.04	-0.10	-0.48	-0.17	-0.15	0.00	-0.14	0.16	-0.33
Q-T JFM	0.08	0.22	0.22	-0.32	-0.10	0.23	-0.09	0.29	0.15	-0.13
Q-T AMJ	-0.40	0.30	0.21	-0.16	-0.35	-0.03	0.32	0.12	0.15	0.04
Q-T JAS	-0.44	0.38	0.26	0.36	-0.16	0.39	0.57	0.16	0.29	0.10
Q-T OND	-0.26	0.25	0.15	-0.13	-0.15	-0.16	-0.18	0.16	0.14	-0.40
Q-P Y	0.42	0.30	0.06	0.34	0.23	0.40	0.17	0.27	0.21	0.39
Q-P JFM	0.25	-0.09	-0.12	0.04	-0.11	-0.01	0.22	0.02	0.19	0.05
Q-P AMJ	0.31	0.18	0.04	0.26	0.33	0.35	0.13	0.18	0.27	0.32
Q-P JAS	0.66	0.21	0.18	-0.01	0.15	0.14	-0.17	0.11	0.21	0.39
Q-P OND	0.02	0.08	0.05	-0.06	-0.07	0.05	0.19	-0.01	-0.18	0.02
P-T Y	-0.46	-0.15	-0.09	-0.37	-0.20	-0.16	-0.30	-0.11	-0.17	-0.36
P-T JFM	-0.02	-0.15	-0.06	-0.06	-0.09	-0.05	-0.06	-0.36	0.00	-0.07
P-T AMJ	-0.03	-0.25	-0.14	-0.03	-0.25	-0.12	-0.29	-0.09	-0.15	-0.18
P-T JAS	-0.45	-0.16	-0.37	-0.40	-0.24	-0.11	-0.23	-0.23	-0.17	-0.39
P-T OND	-0.47	0.06	0.18	0.07	-0.16	-0.05	0.07	-0.29	-0.24	-0.09

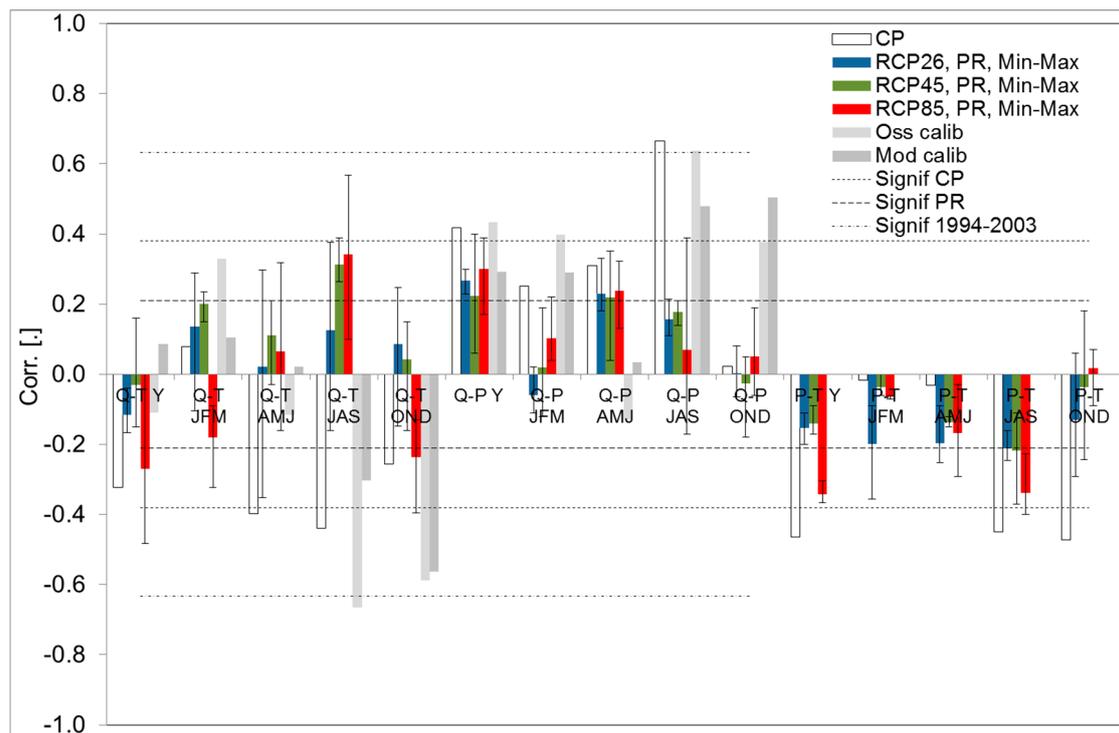


Figure 8. Projected yearly and seasonal correlation coefficients of discharge Q vs. climate drivers T, P, and between P-T until 2100, for PR period as per RCPs vs. CP period. Average RCP values reported with maximum and minimum (among GCMs). Significance values ($\alpha = 5\%$) given for CP, and for PR. In grey shade correlation coefficients of Q vs. T and P during calibration period (1994–2003) for observed, and modelled series, for reference.

5. Discussion

5.1. Glacio-Hydrological Trends, and Flow Components

The *Poly-Hydro* model provides an acceptable description of the stream flows in the Maipo River. Seasonal flows were acceptably simulated both in calibration, and validation phase. One may therefore confidently project forward the catchment response to future climate trends, in the hypothesis of unchanged response patterns. To complement our analysis with explicit investigation of flow components and hydrological fluxes, in Figure 9 we report, for the CP and PR periods, the (average seasonal, per scenarios) relative contribution (averaged at the catchment scale) to in stream flows of snow, and ice melt (lumped). This includes melt ratio $M_r = (q_i + q_s)/q$, rainfall ratio, or the share of stream flow given by rainfall $R_r = 1 - M_r$, and evapotranspiration ratio as $ET_r = ET/(R + M_i + M_s)$. Joint analysis of trends and correlations (Figures 7 and 8) with Figure 6 providing ice ablation, and M_r and ET_r in Figure 9 may provide an explanation of the observed and projected trends of stream flows into the Maipo River under past, and projected future climate conditions. During Summer JFM, flow decrease was projected in the future, continuing the observed significant trends of the recent past (CP period, Figure 7). During CP, no significant trend of P was seen in Summer (slight increase is detected), while T increased significantly (Table 4). From Figure 8 (and Table 5), Q was not significantly correlated to T or P. From Figure 9a, during RCP M_r was constant and nearby 0.99 on average. Snow melt was substantially constant (no large precipitation changes are detected anywhere, Figure 7), and ice depletion occurs during Summer (Figure 2), so the decrease of Q during JFM may likely be due to decreased ice melt (i.e., for lack of ice), say since 1990s' as calculated from our model, and reported in Figure 6. During PR, Q_{JFM} further decreases, especially under RCP8.5 (Figure 7). For RCP 2.6 and RCP4.5 positive correlation of Q_{JFM} vs. temperature was seen, and conversely for RCP8.5 (Figure 8).

Figure 6 indicates that initially in PR increasing ice melt was seen due to increasing temperature, but later decrease occurs especially under RCP2.6 and RCP4.5. From Figure 9a on the long run M_r would decrease slightly, and ET_r would increase instead, especially for RCP8.5. Accordingly stream flows during JFM would largely depends on snow and ice melt, the latter decreasing with time. Increasing ET_r driven by temperatures (especially under RCP8.5) would further decrease stream flow (hence, anti-correlation of Q vs. T in Figure 8). In Fall AMJ, the same trends as in JFM persist. Ice melt would occur in Fall (Figure 2), and decreasing ice melt since the 90s' (Figure 6) may have reduced Q_{AMJ} in the CP period. Also, the slightly increasing ET_r (due to significantly increasing T , Table 4) may have further decreased stream flows (again here, see anti-correlation of Q vs. T in Figure 8). In the PR period, AMJ would be substantially constant. Melt ratio M_r would decrease at the end of century given decreasing ice melt, still higher than now (and constant snow melt, not shown). ET_r would increase due to high temperatures (especially under RCP8.5). However, sustained melting from higher temperatures would still maintain stream flows at the present level during Fall for most of the century. During Winter JAS no significant trend was detected during CP. Stream flows slightly decreased during CP, possibly due to decreased precipitation (Table 4), and also given large significant correlation coefficient of Q vs. P in JAS (Figure 8). Q_{JAS} would slightly increase during PR period, with an eventually stationary behaviour. During PR, both ice and snow melt would increase due to increasing T , and being rainfall substantially constant, M_r would increase (Figure 9c). Increasing ET_r in response to increasing T would not be able to offset increased melting, and Winter stream flows would increase at the end of the century.

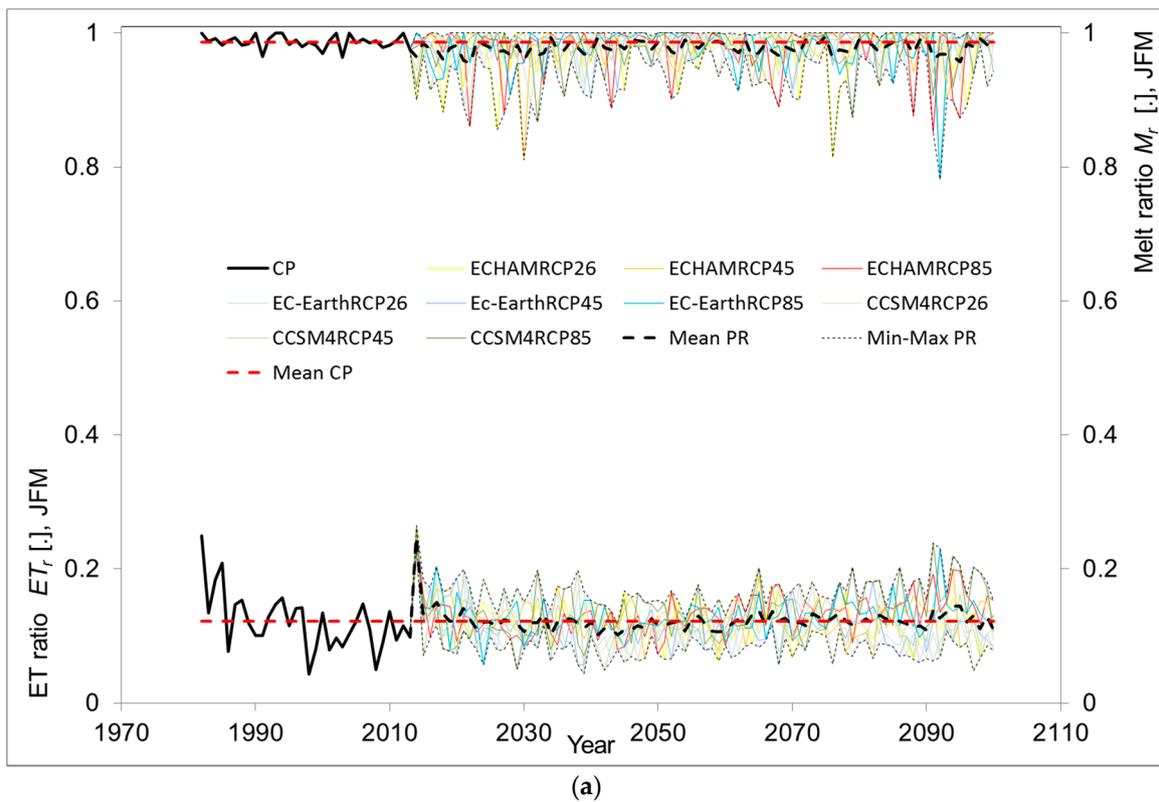
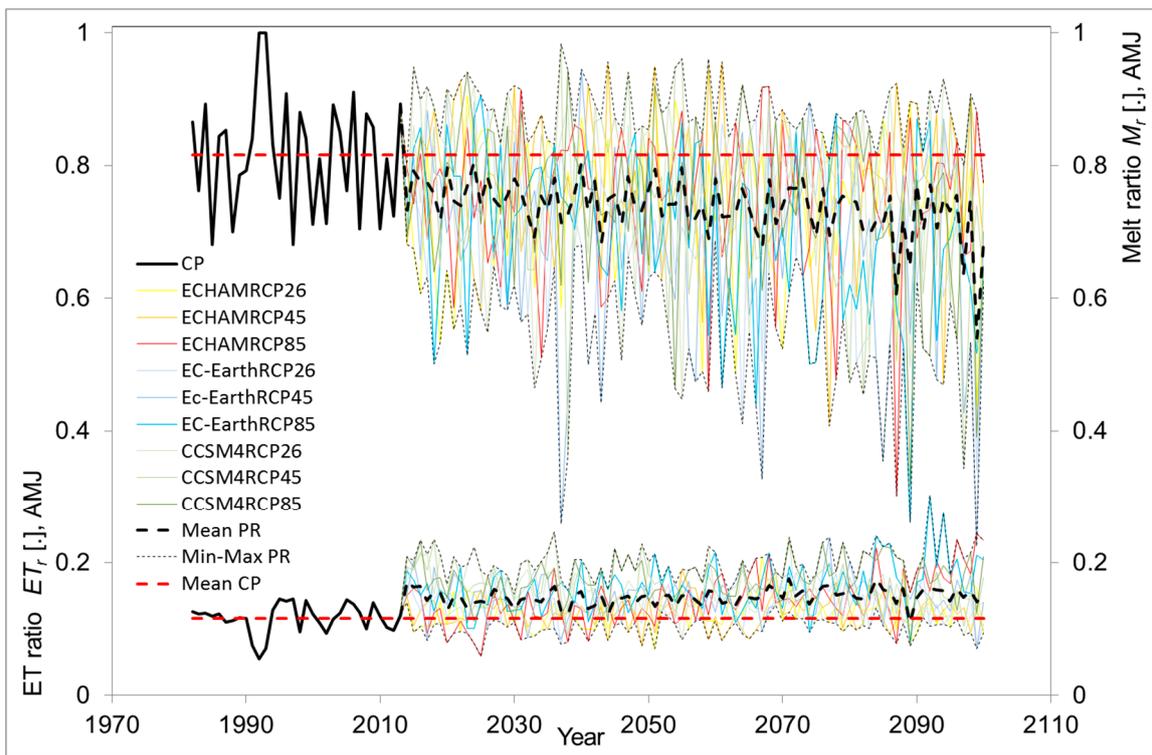
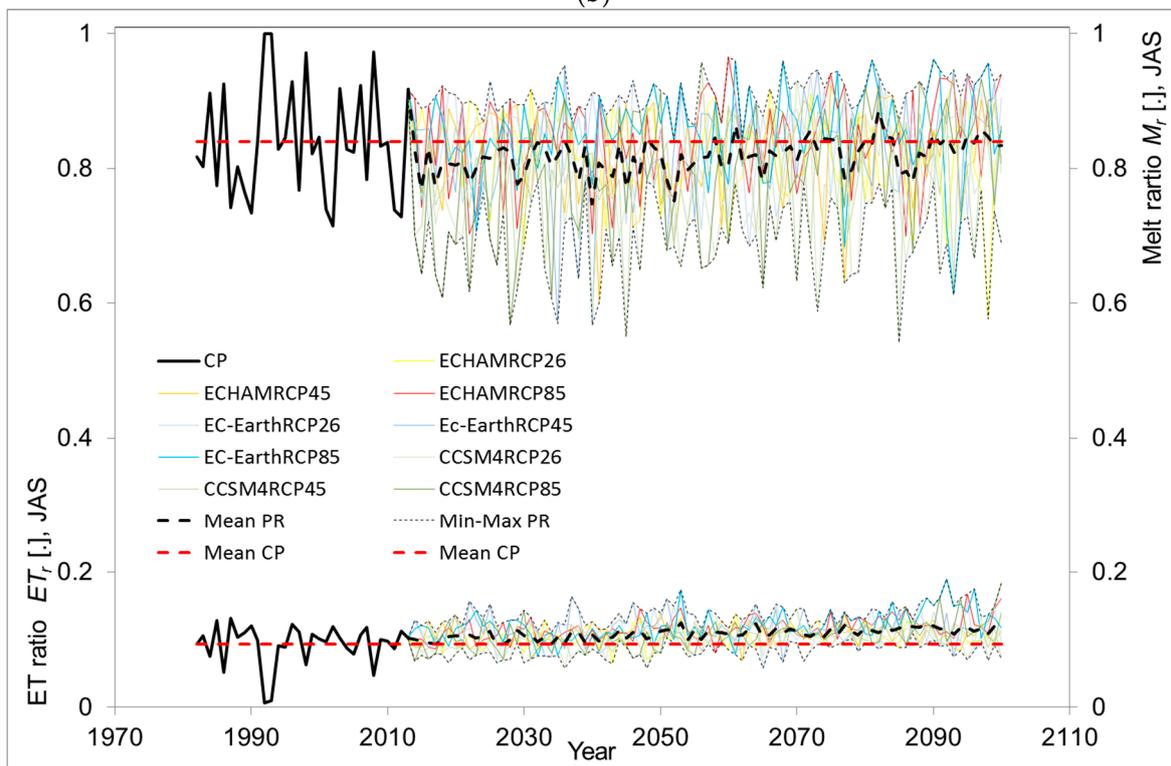


Figure 9. Cont.



(b)



(c)

Figure 9. Cont.

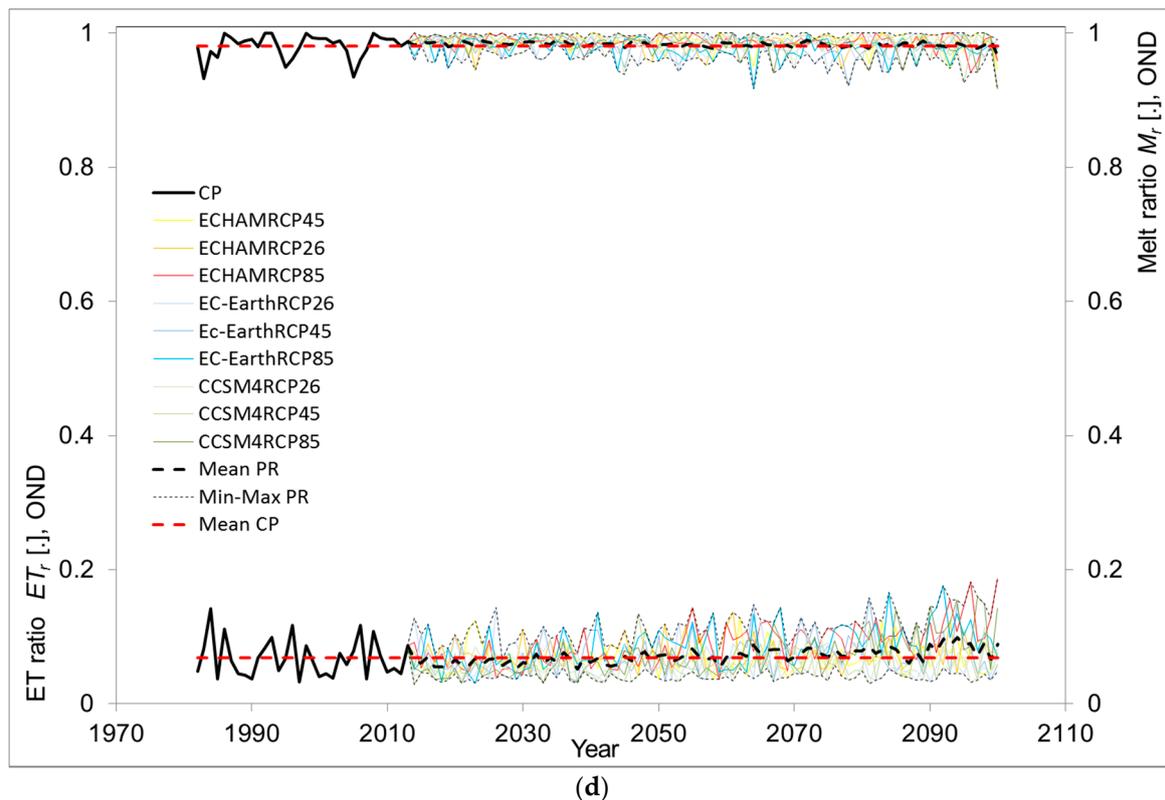


Figure 9. Rio Maipo at El Manzano. Evapotranspiration ratio ET_r (lower y axis), and melt. ratio M_r (upper y axis). (a) Summer JFM. (b) Fall AMJ. (c) Winter JAS. (d) Spring OND.

In Spring OND, historical discharge decreased visibly (albeit not significantly, Table 4, $p\text{-val} = 0.09$), in spite of substantially constant T and P (no visible correlation of Q is seen vs. either). Analysis of ice melt during Spring (not shown) displayed decreasing values during CP, consistent with ice volume decreasing as from Figure 6, which with substantially constant snow melt, may explain decreasing Q . In the PR period, a similar pattern would be seen during Summer JFM. Namely, under RCP2.6 and RCP4.5, higher temperature and increasing ice melt (Figure 6) would lead to high M_r , and to increasing flows (hence, positive albeit small correlation of Q vs. T for OND in Figure 8, RCP2.6, RCP4.5) towards the end of the century (yet with a negative net balance during CM). Under RCP8.5 however, increasing temperature would also largely increase ET_r (Figure 9d), thus providing decreasing Q_{OND} during CP (and significantly negative correlation of Q vs. T for OND in Figure 8, RCP8.5).

At the yearly scale, the projected trends would be consistent with past observations. During CP, significant flow decrease has been detected, given by decreased ice melt since the 90s, and possibly increasing evapotranspiration (as displayed by negative, albeit not significant correlation of Q vs. T for Y in Figure 8, CP). The projected trends indicate substantially constant, or slightly decreasing flows as given trading off by initial increased ice melting, followed by decreasing ice availability, and increasing evapotranspiration, especially under RCP8.5, with the largest negative correlation coefficient Q vs. T (Figure 8, Table 5, Y).

Eventually, the projected scenarios indicate evolution of the hydrology of the Maipo River as dominated by temperature increase, with a twofold effect of increasing ice and snow melt (the latter being faster, but substantially constant at the yearly scale given slight changes in precipitation), and changing ice melt availability, initially larger given the rise of temperature (and altitude of the melting area), but lately smaller for depletion, with evapotranspiration becoming more effective. On average (nine scenarios), at the end of the century, yearly mean discharge would change from

120 $\text{m}^3 \cdot \text{s}^{-1}$ now, to 106 $\text{m}^3 \cdot \text{s}^{-1}$ (−11.3%) at 2100, with a variability between 82 $\text{m}^3 \cdot \text{s}^{-1}$ and 128 $\text{m}^3 \cdot \text{s}^{-1}$ (−32% to +7%).

5.2. Benchmark against the Present Literature

The authors of [6] studied the Juncal Norte glacier, North of Maipo here. With climate scenarios from 2 GCMs (ECHAM5, HadCM3, AR4 of IPCC, 2007) under storyline A1B (medium optimistic) they projected mass balance of Juncal at 2050. Yearly, they projected runoff at −30% under HadCM3, but constant under ECHAM5 ([6], Figure 8). Until 2050 we depicted (Figure 5, decade 2045–2054) a change (vs. 1980–2013) of −11% on average, from +3% (EC-Earth, RCP4.5) to −24% (EC-Earth, RCP8.5). The authors in [6] did not include ice flow, resulting in accumulation at high altitude being unreasonable (see also [26]), and into ablation tongues rapidly down wasting, nor they included evapotranspiration, influencing the water budget under the warmest scenarios. In [13], the authors studied effects of climate change until 2065 on the Maipo River at El Manzano, also based on CEPAL [51]. They projected monthly flows during 2035–2065 to −12% yearly, similarly to here (here −8% on average for the same years, lowest −22%, highest +2.5%). The authors of [12] used a multisite stochastic weather generator to downscale A2 and B2 scenarios, and a hydrological model including crop production to simulate monthly flows at El Manzano for 2071–2100. They projected largely increased (ca. +15%) evapotranspiration like we did here (Figure 9). They further report that during irrigation season (OND+JFM) flow decrease would affect the area. Our runoff estimates for 2071–2100 during irrigation season provide on average −17% (lowest −38%, highest +3%). Krellenberg and Hansjürgens [52] studied water availability in the Maipo based on AR5 projections, and predicted, until 2050, a potential decrease from −19% (B1 storyline) to −30% (A2 storyline), even larger than our RCP4.5 (−8% on average), and RCP8.5 (−15% on average). Our results here are consistent with the available literature (see also [53]), and yet updated (i.e., based upon more recent climate projections from AR5 of IPCC), and physically based (i.e., on a hydrological model including the main physical processes governing water budget, and flow components).

5.3. Limitations and Outlooks

The present work has limitations. Permafrost and rock glaciers dwell in the Maipo catchment [5,54], and were not modelled here. Hydrological dynamics of permafrost, i.e., of the active layer [55,56] needs to be studied in situ, and no knowledge was available that we know of in the case study area. A simple melt model (degree-day) was used here for cryospheric flows, when more complex methods may be used, including mixed (e.g., with solar radiation) temperature index models [5,34], or energy budget models [28]. However, degree day models at the daily scale provide acceptable results for hydrological purposes [15,57], at the cost of limited data requirement, conversely to other more demanding methods (e.g., [58]). We lumped degree day of ice regardless of debris cover, given lack of debris thickness information for glaciers (as used e.g., in [15,26,59]). Marangunic [60] reported glaciers in the Rio Maipo (above Las Melosas, Figure 1) to possess ca. 44% debris cover. DGA [44] more recently reported in the whole catchment 43% debris cover. The large size of cells in *Poly-Hydro* model ($3 \times 3 \text{ km}^2$), necessary for computational reasons may have hidden the variability of glaciers' dynamics. However, we used partial cell coverage, and visually analysed flow velocity and ice flow dynamics, found to be consistent with those measured in situ, and acceptable thickness and shear stress during the simulation [18]. Uncertainty may occur in the initial (2012) ice thickness provided by DGA [40]. At El Manzano in 2012, we mapped 341 km^2 of ice cover, and an estimated volume of 25.2 km^3 of water (i.e., ca. a mean 74 m thickness). DGA [44] for the whole Maipo catchment mapped ca. 388 km^2 , and estimated a volume of ca. 37 km^3 (i.e., ca. a mean 96 m in thickness), largely variable depending on the estimation method. In spite of changes in the absolute thickness (affecting slightly ice flow dynamics), relative thickness change seems large especially under the warmest scenarios, and the broad picture seems not largely affected by the initial conditions. Our back estimation of ice (water) volume at 1982 was indicative. We obtained 29.1 km^3 in 1982 (ca. 85 m thickness on average),

while Marangunic [60] estimated for the Maipo superior (much smaller) ca. 67 km², and ca. 3.4 km³ of water (ca. 51 m in thickness), comparable given large uncertainty of estimation.

Variability of climate projections from different models, especially for precipitation (e.g., [35] for Italy, [15] in the Karakoram, [16,61] for the Himalayas) may influence the results. Seasonal temperature patterns here are consistent between models, but more variability is seen in precipitation. Here, until 2100, a decrease of precipitation was projected on average, and the more so with the warmer RCPs. Largest changes would happen in Fall and Winter, which are the two wet seasons (Table 4). Future steps may include sensitivity analysis against climate projections, e.g., using synthetic simulation (e.g., ensembles of climate patterns based upon GCM scenarios, e.g., [35]) beyond our scope in this manuscript. Even with uncertainty as reported, our results depict consistent pattern, in line with the present know how concerning future hydrology of the central Andes of Chile.

6. Conclusions

Our study assessed impacts of climate change on water resources, and hydrological components in the Maipo catchment of Santiago under updated climate scenarios from the AR5 of IPCC until 2100. During three decades recently Maipo underwent visible flow decrease yearly, most notably in Spring and Summer, when water is needed for irrigation. Decreasing trends were envisioned for the future in those seasons, and more under the warmest RCPs towards mid-century, when ice cover would be largely thinned, and evapotranspiration would draw large amount of moisture. Glacier cover would be increasingly depleted with rising of melting altitude, providing initially more water, but less water after. Stream flows would increase at the cost of ice down wasting during dry, increasingly warmer Fall and especially Winter, but the net effect would be flow decrease. Recent findings [62] indicate that global temperature, when compared against GCMs projections under IPCC AR5 launched in 2006, substantially overlaps with the projected pattern under RCP8.5, i.e., warming proceeded according to the most pessimistic scenario. Seemingly thus, if projections need be made now, globally one may expect that most credible scenarios here are those under RCP8.5. Notwithstanding uncertainty, our results seem consistent, and credible, also in the face of the present literature. Policy makers in Chile are therefore warned, climate change is acting, and will further act to decrease water availability in this region, and deplete ice cover, so that adaptation needs be tackled rapidly enough.

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