

An automatic locally-adaptive method to estimate heavily-tailed breakthrough curves from particle distributions

Daniele Pedretti, Daniel Fernàndez-Garcia

*Hydrogeology Group, Department of Geotechnical Engineering and Geosciences,
Universitat Politècnica de Catalunya (UPC BarcelonaTech), Ed. D2/004/01, C/ Jordi
Girona 1-3, E-08034 Barcelona, Spain*

Abstract

Particle tracking methods to simulate solute transport deal with the issue of having to reconstruct smooth concentrations from a limited number of particles. This is an error-prone process that typically leads to large fluctuations in the determined late-time behavior of breakthrough curves (BTCs). Kernel density estimators (KDE) can be used to automatically reconstruct smooth BTCs from a small number of particles. The kernel approach incorporates the uncertainty associated with subsampling a large population by equipping each particle with a probability density function. Two broad classes of KDE methods can be distinguished depending on the parametrization of this function: global and adaptive methods. This paper shows that each method is likely to estimate a specific portion of the BTCs. Although global methods offer a valid approach to estimate early-time behavior and peak of BTCs, they exhibit important fluctuations at the tails where fewer particles exist. In contrast, locally adaptive methods improve tail estimation while oversmoothing both early-time and peak concentrations. Therefore a new method is proposed combining the strength of both KDE approaches. The proposed approach is universal and only needs one parameter (α) which slightly depends on the shape of the BTCs. Results show that, for the tested cases, heavily-tailed BTCs are properly reconstructed with $\alpha \approx 0.5$.

Keywords: Kernel density estimator, Particle tracking, Heterogeneity,

Email address: daniele.pedretti@upc.edu (Daniele Pedretti)

1. Introduction

Among the several numerical methods that exist to solve the solute transport partial differential equation, Particle Tracking Methods (PTMs) (e.g., [1, 2, 3, 4]) provide a computationally efficient and versatile approach with many interesting applications. Processes such as non-Fickian dispersion [5, 6], rate-limited mass transfer [7, 8, 9, 10, 11, 12], incomplete mixing [13, 14, 15] and complex chemical reactions [16, 17, 18, 19, 20] can easily be addressed in a cost-effective manner without oversimplifying the description of the spatial variability of the physical and chemical properties. Thus, PTMs are especially indicated to simulate advection-dominated problems in natural heterogeneous formations, where large flow velocity contrasts typically translate into numerical problems for Eulerian methods, such as finite differences or finite elements (e.g., [21, 22, 23]).

A major disadvantage of PTMs relies on the need to convert particle distributions into smooth concentrations [18, 21]. Consequently, a challenge in this area is the inference of concentration breakthrough curves (BTCs) from particle attributes (travel times) determined at control locations. Since true concentrations can only be obtained in the limit, when an infinite number of particles are employed, there is always an estimation error in the final solution.

For computational reasons, a generic solute plume can be discretized only into a limited number of particles, usually ranging from 10^4 to 10^7 (e.g. [24, 25]). As a result, BTCs, obtained basically from histograms (counting particles in bins), normally produce discontinuous solutions that strongly depend on the number of particles used and the size of the bin selected. Large bins will yield biased estimates of the BTCs, while small bins will produce large fluctuations [26]. These fluctuations can be reduced by increasing the number of n particles. Unfortunately, the magnitude of these fluctuations is proportional to $n^{-1/2}$ and thereby an unfeasible large number of particles are necessary to minimize the estimation error.

The application of PTMs to many groundwater problems has been widely reported in the literature, especially with regards to risk assessment of polluted sites and solute transport phenomena [27, 28, 29, 30, 31, 32, 33]. A general conclusion from these works is that a proper quantification of the

35 actual threat that potential contaminants pose to human health and ecosys-
36 tems requires an adequate description of the BTCs. Different parts of the
37 BTC are important to different problems [34]. The early arrival of con-
38 centrations determines the potential hazard of a toxic compound, the central
39 part controls human exposure to site-related contaminants during risk assess-
40 ment, and the late arrival of concentrations is the key factor controlling the
41 end-point of remediation [27, 35, 36]. In this context, particle tracking sim-
42 ulations of real aquifer systems often produce highly asymmetric BTCs with
43 sustained low concentrations at large times [11, 24, 37, 38]. This late-time
44 behavior is sometimes known as tailing. Here, we will refer to the complete
45 distribution of concentrations associated to this behavior as ‘heavily-tailed’
46 BTCs. Heavily-tailed BTCs can be observed in porous media due to sev-
47 eral mechanisms. Among them, heterogeneity [25, 38, 39], flow configuration
48 [40], rate-limited mass transfer [35, 41, 42], kinetic reactions [43], and non-
49 linear sorption [44] have been largely addressed. Accurate reconstruction of
50 heavily-tailed BTCs predicted by particle tracking simulations in any of these
51 situations will typically require an enormous number of particles (millions)
52 to describe the entire concentration distribution [21].

53 Kernel density estimators (KDE) offer a valid option to properly recon-
54 struct BTCs with many interesting advantages [18, 45]. It is a non-parametric
55 method that can be used to infer the true structure of the BTC, it allows the
56 automatization of the reconstruction process and it provides a better quality
57 of results for the same number of injected particles [18]. The reader can find
58 complete and exhaustive reviews of these methods in Härdle [46] and Silver-
59 man [26]. Although several KDE methods exist, two main categories can
60 be distinguished depending on the parametrization chosen to characterize
61 the key parameter controlling the smoothing degree, namely the bandwidth:
62 Global KDE methods consider a constant bandwidth size, whereas, on the
63 contrary, adaptive KDE methods allow the bandwidth to change with the
64 particle location. According to Silverman [26], the selection of one of the two
65 methods must depend on the specific goal of the study, since each of them is
66 more likely to estimate a specific portion of the curves while penalizing the
67 other parts. Adaptive methods are preferred to estimate the tails of heavily-
68 tailed probability distributions, while global estimators can provide a more
69 robust estimation of symmetric probability distributions.

70 This paper develops a general method that provides accurate estimates of
71 all parts of the BTC. The approach combines the best of both KDE methods
72 by means of the cumulative distribution function (CDFs) of the particles’

73 travel times, which is a natural output from each model run. The perfor-
74 mance of all methods is analyzed by using benchmark solutions of heavily-
75 tailed breakthrough curves.

76 The paper is structured as follows: Firstly, Section 2 gives some general
77 concepts on particle tracking and KDE methods. Then, Sections 3 and 4
78 respectively introduce the global and adaptive kernel methods. Finally, after
79 a careful examination of the limitation of both methods, Section 5 introduces
80 a new approach, which is contrasted against the previous ones.

81 **2. Particle Tracking Methods and Kernel Density Estimation**

82 Particle tracking methods simulate transport by injecting a large number
83 of particles into the system. Each particle carries a portion of the solute mass
84 that moves with groundwater and reacts according to some fundamental
85 mechanisms. A large variety of mechanisms can be chosen to efficiently
86 simulate different transport phenomena. These mechanisms range from pure
87 advection and random walk motions to complex reaction and mass transfer
88 processes (e.g., [7, 23, 47, 48]).

89 In all cases, the outcome of PTMs is typically expressed as a distribu-
90 tion of particle travel times at control locations and/or positions observed at
91 different times. Since this information is discrete in nature a reconstruction
92 process is mandatory to finally convert particles into concentrations. This
93 reconstruction process is normally seen as the main disadvantage of PTMs
94 [18, 21]. This is based on the relationship between particle distributions
95 and concentrations, which determines that concentrations are proportional
96 to particle probability densities (PDFs) (see, e.g., [49, pag. 167]). Let us con-
97 sider, for instance, the reconstruction of a BTC from a record of particle travel
98 times obtained at a given well location. In this case, flux-concentrations, c_f ,
99 are related to the PDF travel times, $p(t)$, by

$$p(t) = \frac{Q}{m_0} c_f \quad m_0 = \int_0^\infty Q c_f dt, \quad (1)$$

100 where Q is the total flow at the extraction well, and m_0 is the mass under-
101 neath the BTC. The PDF travel time $p(t)$ is defined as the derivative of the
102 travel time cumulative distribution function $P(t)$, where $P(t)$ is the proba-
103 bility that the travel time is smaller than t . For small time increments dt ,
104 the term $p(t)dt$ is the probability that, at time t , the particle travel time falls
105 into the time interval $[t, t + dt]$. Unfortunately, the application of such simple

106 relationship (1) is filled with difficulties. An infinite number of particles is
 107 never used to properly fulfill the validity of this equation.

108 Kernel Density Estimation (KDE) methods were introduced in the late
 109 '50 [50] in order to improve the density estimation of the probability den-
 110 sity function $p(t)$. The need of KDE methods appears crucial when esti-
 111 mating smoothed PDFs with frequency histograms, which are determined
 112 by counting the number of particles within a given time interval $B_j =$
 113 $[t_0 + (j - 1)h, t_0 + jh]$, where h is the time interval (bin size) and t_0 is
 114 the origin of the histogram. The frequency at the j -th bin is determined by

$$p_j \approx \hat{p}_j \equiv \frac{1}{n} \sum_{i=1}^n \frac{1}{h} I\{t_i \in B_j\}, \quad (2)$$

115 where n is the number of particles (arriving at the well location), \hat{p}_j is the
 116 estimated density, t_i is the particle travel time data (t_1, \dots, t_n) , and $I\{\cdot\}$ is
 117 an indicator function defined as $I = 1$ if t_i is inside the bin B_j and $I = 0$
 118 otherwise. By definition, the use of a histogram provides a discontinuous box
 119 function that prevents the estimation of smooth probability density functions.
 120 Moreover, the histogram depends on its origin t_0 and the selection of the
 121 bin size h , which are arbitrary. KDE methods overcome these problems
 122 by replacing the fixed box model in the histogram by a moving weighting
 123 function K centered at the particle location,

$$p(t) \approx \hat{p}(t) = \frac{1}{n} \sum_{i=1}^n \frac{1}{h} K\left(\frac{t - t_i}{h}\right), \quad (3)$$

124 where h is the bandwidth, and K is a weighting function that should be
 125 chosen from symmetric functions of unitary area, i.e.,

$$\int K(\tau) d\tau = 1. \quad (4)$$

126 The density estimation can be seen as an attempt to calculate an his-
 127 togram where every point is the center of the sampling lag interval h . The
 128 area of influence of the i -th datum is restricted to a fixed area, determined by
 129 arbitrarily choosing h . Typical kernel functions are reported in Table 1. The
 130 choice of a specific type of kernel does not substantially influence the final
 131 results. However, Gaussian kernels are usually preferred for mathematical
 132 advantages. We refer to [26] and [46] for details in this subject.

133 In the following sections we present three methods to evaluate h . The
 134 first two methods will be referred herein as the Global Bandwidth Method
 135 (GB) and the Adaptive Bandwidth Method (AB), and they are commonly
 136 used in statistics to reconstruct density distributions. The third method,
 137 which we called the Universal Adaptive Bandwidth Method (UAB), is a new
 138 approach that we developed to merge the capabilities of GB and AB methods
 139 to reconstruct heavily-tailed breakthrough curves.

140 3. The global bandwidth method

141 3.1. Mathematical statement

142 Parameter h controls the degree of smoothing and relates to the area
 143 of influence around a datum. If h is too small, there cannot be sufficient
 144 data to find the proper density function $p(t)$ and the outcome will fluctuate.
 145 If h is too large, distant particles can also contribute to the estimation,
 146 smoothing out the local structure of the data. Thus, an optimum value of
 147 h must exist. Selecting an optimum bandwidth h must be established based
 148 on an objective criterion about the goodness of the estimator given by (3).
 149 A common criterion is the Mean Integrated Squared Error (MISE), formally
 150 written as

$$MISE(h) = E \int (p(t) - \hat{p}(t))^2 dt. \quad (5)$$

151 Knowing that the mean square error, $MSE = E[(p - \hat{p})^2]$, can be ex-
 152 pressed as the sum of the estimator bias, $bias(\hat{p}) = E(\hat{p}) - p$, and its variance,
 153 $var(\hat{p}) = E([\hat{p} - E(\hat{p})]^2)$, $MISE$ can be written as

$$MISE(h) = \int bias^2(\hat{p}(t))dt + \int var(\hat{p}(t))dt. \quad (6)$$

154 Approximate solutions of $MISE$ can only be derived asymptotically, i.e.,
 155 in the limit when $h \rightarrow 0$ and $nh \rightarrow \infty$. Assuming p to be twice differentiable,
 156 the bias and the variance can be approximated by [46]

$$bias(\hat{p}(t)) = \frac{h^2}{2}p''(t)\mu_2(K) + \mathcal{O}(h^2), \quad (7)$$

$$var(\hat{p}(t)) = \frac{1}{nh}\|K\|_2^2 p(t) + \mathcal{O}((nh)^{-1}), \quad (8)$$

157 where $\mu_2(K)$ is the second moment of K , $p''(t)$ is the second derivative of
 158 the estimated function, and the symbol $\|\cdot\|_2^2$ is the L_2 -norm operator, i.e.,
 159 $\|p(t)\|_2^2 = \int p(t)^2 dt$. Note that while the bias error is quadratic in h , the
 160 variance error is inversely proportional to the product of the number of particles
 161 and the kernel bandwidth. Thus, large h values will reduce the variance
 162 (noise), while increasing the bias (oversmoothing).

163 An optimal value of h can be obtained by minimizing (6). Should the
 164 derivative of $MISE$ with respect to h be zero, in the limit when $h \rightarrow 0$ and
 165 $(nh) \rightarrow \infty$, and recalling that $\int p(t)dt = 1$, then [51]

$$h_0 = \left(\frac{\|K\|_2^2}{n\|p''(t)\|_2^2\mu_2(K)} \right)^{1/5}. \quad (9)$$

166 Equation (9) shows that h_0 is inversely proportional to the number of
 167 particles $n^{1/5}$. Thus, the larger the number of particles the smaller the band-
 168 width size should be. Moreover, h_0 depends on the second derivative of the
 169 density function $p''(t)$, i.e., on the shape of the density function in itself, which
 170 is unknown. To deal with the latter issue, several methods have been pro-
 171 posed in the literature [52, 53]. Among them, a common one is the so-called
 172 *plug-in* method, which has been applied for instance by Fernandez-Garcia
 173 and Sanchez-Vila [18] to reconstruct reactions rates from particle distribu-
 174 tions. This method, which is described in detail in Appendix A, is used in
 175 the applications shown in following sections.

176 The global bandwidth (GB) method is obtained by simply substituting
 177 the optimal bandwidth h_0 with the kernel density estimator (3), that is

$$p(t) \approx \hat{p}_0(t) = \frac{1}{n} \sum_{i=1}^n \frac{1}{h_0} K \left(\frac{t - t_i}{h_0} \right), \quad (10)$$

178 3.2. Method performance

179 The performance of the GB method is shown in Figure 1, which com-
 180 pares GB reconstructed BTCs obtained from particle tracking simulations
 181 with corresponding reference solutions. Transport simulations consider the
 182 one-dimensional homogeneous multirate mass transfer (MRMT) model [35],
 183 which simulates advection-dispersion solute transport with rate-limited mass
 184 transfer. In this model, heavily-tailed BTCs result from the diffusion of so-
 185 lute mass previously stored in low permeability areas with almost immobile
 186 water (back-diffusion). The details of the model are given in Appendix B.

187 To our purposes, the phenomena embedded in the model is not important.
188 The model simply provides a way to generate both analytical solutions and
189 particle tracking solutions of heavily-tailed BTCs with similar shapes as those
190 observed in the field. Thus, in principle, the methods presented here can be
191 equally used to reconstruct heavily-tailed breakthrough curves associated to
192 different phenomena.

193 Analytical solutions of the multirate mass transfer model were generated
194 with the STAMMT-L code [54] using a power-law distribution of mass trans-
195 fer coefficients. Table 2 summarizes the parameters used to generate the
196 reference solutions. These solutions were contrasted against reconstructed
197 BTCs obtained from randomly distributed particles, of different population
198 sizes, generated using a well-known particle tracking code, RW3D [25]. For
199 completeness, the numerical algorithm of the random walk particle tracking
200 method used here is also presented in Appendix C. We note that the chosen
201 reference solutions are limited to power-law behavior and thereby other dis-
202 tribution shapes (e.g., multi-modal BTCs) are not specifically studied here.

203 Four dimensionless parameters control the general shape of the BTCs:
204 the Peclet number (Pe), defined as the ratio of the rate of advection to the
205 rate of dispersion; the Damköhler number (Da), defined as the ratio of the
206 rate of mass-transfer to the rate of advection; the field capacity (β), defined
207 as the volume ratio of immobile to mobile water; and the exponent of the
208 power-law distribution of mass transfer rates (k). For a given Pe , when Da
209 is small, BTCs exhibit very peaked shapes with long tails of the form $p \sim t^k$.
210 In this case, the field capacity β determines the area of the BTC underneath
211 the tail. As Da increases, mass transfer occurs more quickly and the process
212 approaches equilibrium, which leads to BTCs with Gaussian-like behavior.
213 In this case, the solute travel distance is retarded with respect to groundwater
214 by a factor of $R = 1 + \beta$. Here, since the focus is on heavily-tailed BTCs,
215 numerical simulations considered a constant Pe with varying Da , k and β
216 values.

217 Two different numbers of particles were used to generate the distributions
218 shown in Figure 1. The solid thick blue lines refer to the reconstructed BTCs
219 obtained for $n = 10^4$, while the solid thin red lines illustrates those obtained
220 for $n = 10^5$. We note that the GB method closely matches both the early
221 arrival and the peak of the reference BTCs. This occurs for both particles
222 datasets, and for any degree of symmetry of the reference distributions (i.e.,
223 any k - Da pairs). However, it performs poorly at late times (after the peak).
224 Starting from an estimated concentration close to $1/n$, the numerical solution

225 oscillates as the estimation is performed based on a limited number of parti-
 226 cles. Eventually, it tends to a fictitious asymptotic plateau at late times; in
 227 the case that the estimation is performed using only one particle, the prob-
 228 ability of the plateau is $\widehat{p}_0(t) = (nh_0\sqrt{2\pi})^{-1}$, using a Gaussian kernel (Table
 229 1). This effect is more pronounced for heavily-tailed distributions obtained
 230 with small Da and number of injected particles n . Nevertheless, fluctuations
 231 are less noticeable in symmetric BTCs with small tailing obtained for rela-
 232 tively large Da . This result suggests that the performance of the GB method
 233 is best for symmetric distribution of particle travel times. Unfortunately, this
 234 is usually not the case in subsurface hydrology applications.

235 4. Adaptive bandwidth method

236 4.1. Mathematical statement

237 Adaptive bandwidth (AB) methods are considered to be an efficient man-
 238 ner to deal with long-tailed distributions [26]. Long-tailed distributions sug-
 239 gest that the mass extends over a wider range of time values. Intuitively,
 240 as the number of particles used to estimate their density diminishes, the
 241 smoothing parameter h should increase to allow a smoother estimation in
 242 regions where data are lacking. This explains why errors appeared to be
 243 more pronounced at the BTCs tails in Figure 1. To partially overcome this
 244 problem, a common procedure consists in deterministically varying h as a
 245 function of the local data density. This can be formally written as

$$p(t) \approx \widehat{p}_1(t) = \frac{1}{n} \sum_{i=1}^n \frac{1}{h_1(t_i)} K\left(\frac{t-t_i}{h_1(t_i)}\right). \quad (11)$$

246 A common approach to find $h_1(t_i)$ is suggested by Silverman [26] and
 247 it is tested here against the reference distributions previously shown. This
 248 approach considers that h_i is a function of a normalized local data density
 249 measure λ_i , which in turn is a function of $\widehat{p}_0(t)$ (defined in (10)). That is

$$h_1(t_i) = h_0\lambda_i, \quad \lambda_i = \left(\frac{\widehat{p}_0(t_i)}{G}\right)^{-\alpha}, \quad G = \exp\left(\frac{1}{n} \sum_{i=1}^n \ln(\widehat{p}_0(t_i))\right), \quad (12)$$

250 where α is a sensitivity parameter fulfilling $0 \leq \alpha \leq 1$. Note that h_1 is no
 251 longer a constant value but encouraged by the first estimate to change from

252 particle to particle. Note that the new density estimate does not follow a
 253 second optimization process but relies on the first estimated value h_0 . Choos-
 254 ing a particular value of α determines how important the shape of the first
 255 guess is with respect to the the second estimation. Although there is not a
 256 well-defined criterion to obtain a prior estimate of α , it is worth mentioning
 257 that Abramson [55] found a theoretical response given by $\alpha = 0.5$.

258 *4.2. Method Performance*

259 The comparison between the reference solutions and the particle tracking
 260 BTCs obtained by the AB method is shown in Figure 2 for $\alpha = 0.5$. The AB
 261 method performs reasonably well for heavily-tailed BTCs (low Da). However,
 262 an underestimation of peak values and departure from the analytical solution
 263 at late time is clearly visible for increasing Da (higher symmetry). The
 264 reason is related to the adaptive h . Let us focus for instance on the pair ($k =$
 265 $3, Da = 2.71$). In this case the kernel estimation is erroneously overweighting
 266 the late-time concentrations, significantly sacrificing the early portions of the
 267 BTCs. A surprising effect is that the AB method seems to work better for a
 268 small number of particles n in the case of low Da (more asymmetric curves).

269 Quantitatively, the adequacy of the fit can be determined by the relative
 270 local ϵ error between the estimated and the reference (analytical) density
 271 functions, which is generally written as

$$\epsilon(t) = \left| \frac{\widehat{p}(t) - p(t)}{p(t)} \right|. \quad (13)$$

272 The results are shown in Figure 3, where the black lines with circles refer
 273 to the GB method and the dashed lines depict the errors of the AB method
 274 for different α values. Note that the scales are logarithmic to emphasize
 275 that ϵ can span over several orders of magnitude. The GB method shows
 276 minimum error around the concentration peaks (where particle densities are
 277 higher) but performs poorly at late times. This is especially true for heavily-
 278 tailed BTCs with low Da . The opposite is also true for the AB method. In
 279 this case, the reconstruction of tailing in heavily-tailed BTCs with $k = 1$
 280 displays relatively small errors.

281 In short, the AB method is capable of improving the quality of highly
 282 skewed BTCs with low Da (with respect to the GB method), but their cor-
 283 responding local errors are in general dramatically large for other cases. Es-
 284 pecially, it causes a general oversmoothing of the concentration peaks posing
 285 serious concerns about the reliability of the reconstructed BTCs, as it can

286 give misleading solute transport parameters with a catastrophic impact on
287 decision-making processes (e.g., in risk assessment or aquifer remediation
288 activities). A solution to this problem is presented in the following section.

289 5. A universal adaptive method

290 5.1. Mathematical statement

291 The analysis of the two KDE methods described so far suggests that
292 the choice of the reconstruction method depends on the type and goal of
293 the problem. While the GB method can properly estimate the early arrival
294 and peak of concentrations, the AB method can better reproduce the tail of
295 heavily-skewed BTCs.

296 Unfortunately, in most hydrogeological applications the actual structure
297 of the BTC is not known *a priori* and the selection of the kernel estimator
298 is uncertain. During tracer tests performed in heterogeneous aquifers, for
299 instance, the shape of BTCs measured after injecting solutes from different
300 positions is highly variable (e.g., [56]) and can be drastically different from
301 ideal parametric distributions. Dimensionless transport parameters control-
302 ling the shapes of the BTCs (such as Pe , Da and β factors) are obtained
303 *a posteriori* by model fitting (e.g., [38]) and cannot be used to select the
304 method *a priori*.

305 Therefore, the aim is to look for a common and general procedure inde-
306 pendent from the final shape of the BTCs. In this case, we proposed a new
307 KDE approach that combines the strength of both methods (GB and AB).
308 In sum, as extracted from previous analysis, the bandwidth h_i should be sim-
309 ilar to h_0 in (9) at the rising limb of the BTCs and around the concentration
310 peak, but it also should increase at late times to smooth out the solution
311 where particle information is scarce, similarly to h_1 in (12). This means find-
312 ing a new equation that could incorporate this behavior of h_i as a function
313 of time. A suitable and appealing function displaying these characteristics
314 is the experimental cumulative density function, $P(t)$. A quick and reliable
315 manner to obtain $P(t)$ is by sorting in ascending order the vector of arrival
316 times ($t_1 < t_2 < \dots < t_n$) so that

$$P(t_i) \approx \frac{i - C}{n}, \quad i = 1, \dots, n, \quad (14)$$

317 where C gives the functional form of P for a datum (e.g., for a triangle
318 shape, $C = 0.5$; for a box shape, $C = 0$). $P(t_i)$ always takes values comprised

319 between $[0, 1]$, being lower around the early time and increasing in the regions
 320 where less mass is used to estimate the BTC for larger times. The new
 321 estimator for the normalized BTC can be this written as

$$p(t) \approx \widehat{p}_2(t) = \frac{1}{n} \sum_{i=1}^n \frac{1}{h_2(t_i)} K\left(\frac{t - t_i}{h_2(t_i)}\right). \quad (15)$$

322 where the new adaptive bandwidth method is

$$h_2(t_i) = (1 - P(t_i))h_0 + P(t_i)h_1(t_i). \quad (16)$$

323 We call this method universal adaptive bandwidth (UAB). The executable
 324 file of the code along with an example of its application is available on the
 325 GHS-UPC software web page at <http://www.h2ogeo.upc.es/software/quern/index.htm>.

326 5.2. Method Performance

327 Recalling that h_1 is sensitive to the α shape parameter (see (12)), we test
 328 this method using three different and increasingly values of α , ranging from
 329 $\alpha = 0.25$ to $\alpha = 1.0$ ($\alpha = 0$ corresponds to the GB method). The results are
 330 plotted in Figures 4, 5 and 6, respectively. We note from these figures that
 331 the UAB method substantially improves the estimation obtained for each
 332 combination of parameters, and works satisfactorily well for both heavily-
 333 tailed and less skewed BTCs.

334 More specifically, results from this numerical sensitivity analysis suggest
 335 that not only the oversmoothing effects around the concentration peak dis-
 336 appear (especially by using $\alpha = 0.5$ and $\alpha = 1$) but tailing is also properly
 337 captured, matching well with the reference analytical solutions. This is an
 338 effect of the shape of the new universal adaptive bandwidth, $h_2(t_i)$. Figure 7
 339 compares the behavior of the latter function with the one of the AB method,
 340 $h_1(t_i)$. Both functions are normalized by h_0 and are functions of time, when
 341 $k = 2$, $\beta = 1$ and $Da = 0.009$. Note that the AB method produces ex-
 342 tremely large normalized bandwidth values at early stages, giving raise to
 343 the oversmoothing effect on the reconstructed BTCs. On the contrary, the
 344 UAB method keeps the values close to h_0 at early times, allowing a good
 345 reconstruction of the early behavior and peak of the BTCs. At late times,
 346 both solutions approach the same values, so that the accuracy of the adap-
 347 tive method in that region is maintained. Table 3 summarizes the different
 348 bandwidth solutions adopted in this analysis.

349 The critical question now is which α parameter to choose for the UAB
 350 method. The goal is to find a general parameter capable of reconstructing
 351 the entire BTC at once. From a quantitative perspective, the ϵ criterion
 352 helps discerning which α provides the lowest local error and a more accurate
 353 local estimation for each reference case. Figure 8 shows that at shorter times,
 354 all errors collapse to those of the GB method, while at larger times the UAB
 355 method is always the best solution. In general, results show that $\alpha = 0.5$ is
 356 the most reliable value for a universal and robust estimation:

- 357 • When $\alpha = 0.25$ (Figure 4), the UAB method provides very few benefits
 358 at late times, where the estimated BTCs still oscillate. In this case, the
 359 bandwidth h_2 in (16) is still controlled by the global bandwidth esti-
 360 mate h_0 rather than from its adaptive counterpart h_1 and thereby the
 361 reconstruction process gives results similar to those of the GB method.
- 362 • When $\alpha = 1.0$ (Figure 6), fluctuations almost disappear and local errors
 363 are in most cases small, but the solution departs from the reference
 364 distribution at large times for the most symmetric distributions (high
 365 Da). In this case, the bandwidth h_2 in (16) is mostly controlled by the
 366 adaptive bandwidth h_1 rather than from its global counterpart h_0 and
 367 thereby the reconstruction process gives results similar to those of the
 368 AB method.
- 369 • When $\alpha = 0.5$ (Figure 5), an excellent agreement is found for all cases.
 370 Even though residual fluctuations are still visible in BTCs with small
 371 Da and k , the solution is never biased (never departs from the refer-
 372 ence distribution) and a minimum expression of fluctuations is obtained
 373 compared to the solution provided by other α values. Note that, despite
 374 we have not proved this theoretically for the UAB method, a selection
 375 of $\alpha = 0.5$ is consistent with the theoretical analysis of the AB method
 376 provided by Abramson [55].

377 Another important aspect of the UAB method is in terms of computa-
 378 tion cost. Particle tracking simulations directly depends on the number of
 379 particles injected. So, the less particles injected the more efficient is the
 380 simulation. In this context, it is important to highlight that the proposed
 381 UAB method is capable of providing the same quality of results with less
 382 particles. A quick comparison of the reconstructed BTCs shown in Figures

383 1-6 illustrates that the use of the GB and AB methods with $n = 10^5$ per-
 384 form less accurately than the UAB method with $\alpha = 0.5$ and $n = 10^4$. To
 385 quantitatively examine this in more detail, the dependence of the root mean
 386 square error $RMSE$ on the number of particles injected during simulations
 387 was analyzed,

$$RMSE(n) = \sqrt{\frac{1}{n} \left(\sum_{i=1}^n (\epsilon(t_i))^2 \right)}. \quad (17)$$

388 Results are shown in Figure 9. The UAB method with α values ranging
 389 from 0.5 to 1.0 always gives the smallest overall error, independently from
 390 the final shape of the BTC. Most interestingly and remarkably important is
 391 that, for the same quality of results, i.e., equal $RMSE$, the number of parti-
 392 cles needed in the UAB method can be several orders of magnitude smaller
 393 than the GB method. On a computer equipped with Intel(R) Xeon(R) CPU
 394 with 2.80 GHz and 6 GB of RAM memory, the time needed to run one
 395 estimation with the UAB method is about 13 s using $n = 10^4$ and about
 396 780 s using $n = 10^5$. Thus, performing a large number of Monte Carlo
 397 simulations with complex chemical reactions could dramatically reduce the
 398 computational cost.

399 6. Conclusions

400 Particle tracking methods constitute an attractive method to simulate a
 401 large variety of solute transport phenomena. Yet, severe fluctuations arise
 402 when converting a limited number of particles into concentrations. In this pa-
 403 per, the advantages and limitations of reconstructing BTCs were examined
 404 using two broad classes of kernel density estimators: global and adaptive
 405 methods. Results have shown that each method is more oriented to estimate
 406 a specific portion of the BTC. While global methods have offered a valid
 407 approach to estimate early travel times and peak concentrations, they still
 408 showed important fluctuations at the tails where fewer particles exist. In con-
 409 trast, adaptive methods appeared to be capable of improving the tail of the
 410 BTCs while oversmoothing both early travel times and peak concentrations.

411 Based on this, a new method was proposed using the cumulative distri-
 412 bution of particles to locally adapt the smoothing parameter to the particle
 413 density at any given location. Thus, the smoothing parameter is progres-
 414 sively increased at the tails to allow a smoother estimation in regions where

415 particles are lacking. By doing this over a wide range of heavily-tailed BTCs,
416 the proposed method has showed a more complete reconstruction of all parts
417 of the BTCs at once, including the tail. Furthermore, the proposed method
418 only needs one parameter α that slightly depends on the shape of the BTCs.
419 A universal value of $\alpha \approx 0.5$ is shown to work well in all the tested cases.

420 The computational efficiency of the method was evaluated and contrasted
421 against other KDE methods. Results showed that the proposed method can
422 largely reduced CPU time during particle tracking simulations. Thus, it is
423 shown that to get the same quality of results the number of particles needed
424 by the new method may be several orders of magnitude smaller than with
425 the rest of methods. Among the advantages of the method is the fact that
426 that Monte Carlo transport simulations of complex chemical systems may
427 now be feasible at a minimum computational cost.

428 **Acknowledgments**

429 This work has been supported by the Spanish Ministry of Science and In-
430 novation through the projects Consolider-Ingenio 2010 CSD2009-00065 and
431 FEAR CGL2012-38120. DP acknowledges the funding provided by the Span-
432 ish Ministry of Education through the FPU-MED 2009 Scholarship program.

Figure 1: Comparison between reference BTCs obtained from analytical solutions and those estimated using the global bandwidth (GB) method. Simulations are run using different numbers of particle datasets. Parameters k indicate the slope of the BTCs, β is the capacity coefficient and Da is the Damköhler number.

Figure 2: Comparison between reference BTCs obtained from analytical solutions and those estimated using the adaptive bandwidth (AB) method.

Figure 3: Absolute relative errors (ϵ) between the benchmark and estimated BTCs using global bandwidth (GB) method and adaptive bandwidth (AB) methods. The dotted line indicates the time of the concentrations peak (T_P) of the corresponding BTCs.

Figure 4: Comparison between reference BTCs obtained from analytical solutions and those estimated using the universal adaptive method (UAB) with $\alpha = 0.25$.

Figure 5: Comparison between reference BTCs obtained from analytical solutions and those estimated with using the universal adaptive method (UAB) with $\alpha = 0.5$.

Figure 6: Comparison between reference BTCs obtained from analytical solutions and those estimated using the universal adaptive method (UAB) with $\alpha = 1$.

Figure 7: Shape of the bandwidths for AB and UAB methods to reproduce case $k = 2$, $\beta = 1$ using $n = 10^4$. Log-scale emphasizes the fact that the weights for the universal adaptive method (UAB) are closer to $h_j(t_i)/h_0 = 1$ ($j=1,2$ depending on the method) at early time and around the peak, while they tend to the adaptive method (AB) values as time elapses, for each α value. T_P indicates the temporal position of the density peak

Figure 8: Absolute relative errors (ϵ) between the benchmark and estimated BTCs using global bandwidth (GB) method and the universal adaptive method (UAB).

Figure 9: Root mean square error ($RMSE$) of the different methods as a function of the number of particles used to estimate the BTCs.

Table 1: Examples of kernel functions (from Silverman [26])

Type of kernel	$K(\tau)$
Biweight	$\begin{cases} \frac{3}{4\sqrt{5}}(1 - \frac{\tau^2}{5}) & \text{for } \tau < 1 \\ 0 & \text{otherwise} \end{cases}$
Epanechnikov	$\begin{cases} \frac{15}{16}(1 - \tau^2)^2 & \text{for } \tau < \sqrt{5} \\ 0 & \text{otherwise} \end{cases}$
Gaussian	$(2\pi)^{-1/2} \exp(\frac{-\tau^2}{2})$
Rectangular	$\begin{cases} \frac{1}{2} & \text{for } \tau < 1 \\ 0 & \text{otherwise} \end{cases}$
Triangular	$\begin{cases} 1 - \tau & \text{for } \tau < 1 \\ 0 & \text{otherwise} \end{cases}$

Table 2: Parameters for analytical and numerical solutions: v = pore-scale velocity; a =longitudinal dispersivity; L =domain length; k =distribution power; β =capacity coefficient; w = mass transfer rates; Pe =Peclet Number; Da =Damköhler numbers.

Variable	Value [dimensions]
v	1 [LT ⁻¹]
a	0.05 [L]
L	5 [L]
Pe	100 [-]
k	1; 2; 3 [-]
β	1; 5 [-]
w_{min}	10 ⁻⁴ [T ⁻¹]
w_{max}	1 [T ⁻¹]
Da ($k=1, \beta = 1$)	0.002 [-]
Da ($k=1, \beta = 5$)	0.005 [-]
Da ($k=2, \beta = 1$)	0.009 [-]
Da ($k=2, \beta = 5$)	0.023 [-]
Da ($k=3, \beta = 1$)	1.08 [-]
Da ($k=3, \beta = 5$)	2.71 [-]

Table 3: Bandwidth functions used in the three kernel density models. See the text for the definition of each parameter.

Method	Bandwidth function
Fixed bandwidth (GB)	h_0
Globally adaptive bandwidth (AB)	$h_0(\widehat{p}_0(t_i)/G)^{-\alpha}$
Universally adaptive bandwidth (UAB)	$(1 - P(t_i))h_0 + P(t_i)(\widehat{p}_0(t_i)/G)^{-\alpha}$

433 **Appendix A. The plug-in method**

434 The optimal bandwidth involves the unknown function p'' which needs
 435 to be further estimated. Here, we used the plug-in method presented by
 436 Engel [52], who employed another kernel estimate of p'' based on a different
 437 bandwidth b to estimate $\|p''\|_2^2$ by

$$\|p''\|_2^2 \approx R(b) = \frac{1}{n^2 b^5} \sum_{i,j} L_2 \left(\frac{X_{p,i} - X_{p,j}}{b} \right). \quad (\text{A.1})$$

438 The functional L_2 is the convolution of K'' by itself, $L_2 = K'' \circ K''$. The
 439 bandwidth b is chosen to be a linear function of h_0 with the form $b(h_0) =$
 440 $Cn^q h_0$, where C and q are constants. Briefly, the algorithm presented by
 441 [52] is as follows: (1) Select an initial bandwidth, $h_0^{(0)} = Bn^{-r}$; (2) Iterate
 442 $i = 1, \dots, I$ with $b^{(i-1)} = Cn^q h_0^{(i-1)}$ and $h_0^{(i-1)} = (\|K\|_2^2 / R(b^{(i-1)})) \mu_2^2(K) n^{1/5}$;
 443 and finally, (3) set the bandwidth b to $b^{(I)}$. Tuning parameters $\{C, q, r, B, I\}$
 444 required for the implementation of this algorithm are given in [52].

445 **Appendix B. The multirate mass transfer model**

446 The multi-rate mass transfer model [35] describes the porous media as
 447 a multiporosity system formed by a mobile domain, in which solute moves
 448 by advection and dispersion, and many immobile regions, in which water is
 449 mostly stagnant and solute transport is controlled by diffusion. This con-
 450 ceptualization allows one to represent a system that encompasses different
 451 mass transfer characteristics (different geometries, shapes, mineralogy,...).
 452 Provided that, initially, the total mass is only present in the mobile domain
 453 by the condition $p(t) = \delta(t)$, the governing equation of the MRMT can be
 454 formally written as

$$\frac{\partial p}{\partial t} + \beta \frac{\partial}{\partial t} \int_0^t g(t - \tau) p(\tau) d\tau = -v \frac{\partial p}{\partial x} + av \frac{\partial^2 p}{\partial x^2} \quad (\text{B.1})$$

455 where $g(t)$ is known as the memory function [41, 57], v is the velocity, β
 456 is the capacity coefficient (ratio between total immobile porosity and mobile
 457 porosity), and a is the local dispersivity. The second term of the equation is a
 458 source/sink term that represents the mass exchange from mobile to immobile
 459 regions. The memory function can be formally defined as the probability of
 460 residence times in the immobile region and relates to the probability density
 461 function of mass transfer rates $f(w)$ by

$$g(t) = \int_0^{\infty} w f(w) e^{-wt} dw \quad (\text{B.2})$$

462 Among the various probability density functions $f(w)$ described in the
 463 literature, we considered the truncated power-law function of the form [57],

$$f(w) = \mathcal{A}(k) w^{k-3}, \quad w_{min} < w < w_{max} \quad (\text{B.3})$$

464 where k is the parameter of the memory function, and $\mathcal{A}(k)$ is defined as

$$\mathcal{A}(k) = \begin{cases} (k-2)/(w_{max}^{k-2} - w_{min}^{k-2}) & \text{if } k \neq 2 \\ 1/\ln(w_{max}/w_{min}) & \text{if } k = 2 \end{cases} \quad (\text{B.4})$$

465 The shape of the resulting BTCs is controlled by k , β and two dimension-
 466 less numbers, the Peclet number (Pe), which relates advective and dispersive
 467 characteristic times, and the Damköhler number (Da), which relates advective
 468 and mass-transfer characteristic times. Both numbers can be written

$$Pe = \frac{L}{a}, \quad Da = \frac{(1+\beta)Lw_h}{v}, \quad (\text{B.5})$$

469 where L is the domain length, w_h is the harmonic mean of mass-transfer
 470 rates, and v is the groundwater velocity. If a power-law memory function
 471 is selected, w_h depends on w_{min} , w_{max} and k (this value was computed here
 472 numerically [54]). Note that for small Da values, the selection of a power-law
 473 memory function leads to BTCs with late-time behavior of the form $p \sim t^k$
 474 . The analytical solutions of the MRMT model can be found by numerical
 475 inversion of the Laplace transform of equation (B.1). A suitable code for this
 476 purpose is STAMMT-L [54], which is the one used here. Table 2 summarizes
 477 the parameters used to generate the reference solutions.

478 Appendix C. The random walk algorithm

479 This appendix describes the random walk algorithm used to simulate
 480 multirate mass transfer. The algorithm follows the method developed and
 481 implemented in the RW3D code by Salamon et al. [25]. The displacement
 482 of an i th particle X_i is given by a drift term that simulates advection and a
 483 superposed Brownian motion responsible for dispersion,

$$X_i(t + \Delta t) = X_i(t) + v\Delta t + \xi(t)\sqrt{2av\Delta t}, \quad (\text{C.1})$$

484 where v is the groundwater velocity, a is the dispersivity, Δt is the time
 485 step, and $\xi(t)$ is an independent (uncorrelated in space and time), normally
 486 distributed random variable with zero mean and unit variance. RW3D sim-
 487 ulates multirate mass transfer by also tracking in time the state of a par-
 488 ticle, which defines the domain (mobile/immobile) at which the particle is
 489 present at a given time within the multiporosity system [25]. To apply this
 490 with a continuous distribution of mass transfer rates, the probability den-
 491 sity function $f(w)$ must be first discretized into a set of N zones, so that
 492 $f(w) = \sum_{i=1}^N f_i \delta(w - w_i)$ and $\sum_i^N f_i = 1$. The individual probabilities f_i can
 493 be estimated as

$$f_i = \int_{w_i - \Delta w_i/2}^{w_i + \Delta w_i/2} f(w) dw, \quad (\text{C.2})$$

494 where w_i , $i = 1, \dots, N$ is the discrete space of mass transfer rates, and Δw_i
 495 is the interval size of the discretization. A specific formulation is found for
 496 the power-law density function

$$f_i = \frac{(w_i + \Delta w_i/2)^{k-2} - (w_i - \Delta w_i/2)^{k-2}}{w_{max}^{k-2} - w_{min}^{k-2}}, \quad k \neq 0, \quad (\text{C.3})$$

$$f_i = \frac{1}{\ln(w_{max}/w_{min})} \ln \left(\frac{2w_i + \Delta w_i}{2w_i - \Delta w_i} \right), \quad k = 2. \quad (\text{C.4})$$

497 The change from one state to another is easily determined using transition
 498 probabilities. The transition probability, $P_{ij}(\Delta t)$, that a particle presently
 499 in state j will be in state i at a time $t + \Delta t$ is given by

$$\mathbf{P}(\Delta t) = \begin{pmatrix} -\beta \sum_{j=1}^N w_j f_j & w_1 & \cdots & w_N \\ \beta w_1 f_1 & -w_1 & 0 & 0 \\ \vdots & 0 & \ddots & 0 \\ \beta w_N f_N & 0 & 0 & -w_N \end{pmatrix}. \quad (\text{C.5})$$

500 If a particle is in the mobile domain, then it is susceptible to advection
 501 and dispersion, otherwise the particle is not allowed to move. Having cal-
 502 culated the transition probabilities, numerical implementation into particle

503 tracking is done without difficulties. For each time step a uniform $[0, 1]$ ran-
504 dom number r is drawn for each particle and the new particle state is taken
505 as that integer s for which

$$\sum_{i=1}^{s-1} P_{ij}(\Delta t) < r \leq \sum_{i=1}^s P_{ij}(\Delta t). \quad (\text{C.6})$$

506 **References**

- 507 [1] S. W. Ahlstrom, H. P. Foote, R. C. Arnett, C. R. Cole, R. J. Serne,
508 Multi-component mass transport model: theory and numerical im-
509 plementation (discrete parcel random walk version)., Tech. rep., Rep.
510 BNWL-2127. Battelle Pacific Northwest Lab., Richland, Washington
511 (1977).
- 512 [2] W. Kinzelbach, The random walk method in pollutant transport simula-
513 tion. advances in analytical and numerical groundwater flow and quality
514 modelling., Tech. rep., In: Custodio, E., et al. (Eds.), NATO ASI Series
515 C, vol. 224, pp. 227-246. (1987).
- 516 [3] D. W. Pollock, Semianalytical computation of path lines for finite-
517 difference models, *Ground Water* 26 (1988) 743–750.
- 518 [4] T. A. Prickett, T. G. Naymik, C. G. Longquist, A random walk so-
519 lute transport model for selected groundwater quality evaluations., Tech.
520 rep., Illinois State Water Survey. Bulletin, vol. 65. 103 pp. (1981).
- 521 [5] B. Berkowitz, A. Cortis, M. Dentz, H. Scher, Modeling non-fickian trans-
522 port in geological formations as a continuous time random walk, *Reviews*
523 *of Geophysics* 44 (2006) RG2003. doi:10.1029/2005RG000178.
- 524 [6] Y. Zhang, E. M. LaBolle, K. Pohlmann, Monte carlo simu-
525 lation of superdiffusion and subdiffusion in macroscopically het-
526 erogeneous media, *Water Resources Research* 45 (2009) W10417.
527 doi:10.1029/2008WR007448.
- 528 [7] D. A. Benson, M. M. Meerschaert, A simple and efficient random walk
529 solution of multi-rate mobile/immobile mass transport equations, *Ad-
530 vances in Water Resources* 32 (2009) 532–539.
- 531 [8] F. Delay, J. Bodin, Time domain random walk method to simulate trans-
532 port by advection-dispersion and matrix diffusion in fracture networks,
533 *Geophys. Res. Lett.* 28 (21) (2001) 4051. doi:10.1029/2001GL013698.
- 534 [9] M. Dentz, B. Berkowitz, Transport behavior of a passive solute in con-
535 tinuous time random walks and multirate mass transfer, *Water Resour.
536 Res.* 39 (5) (2003) 1111–1131. doi:10.1029/2001WR001163.

- 537 [10] S. Painter, V. Cvetkovic, J. Mancillas, O. Pensado, Time domain parti-
538 cle tracking methods for simulating transport with retention and first-
539 order transformation, *Water Resources Research* 44 (2008) W01406.
540 doi:10.1029/2007WR005944.
- 541 [11] P. Salamon, D. Fernàndez-Garcia, J. Gomez-Hernandez, Modeling mass
542 transfer processes using random walk particle tracking, *Water Resour.*
543 *Res.* 42 (2006) W11417. doi:10.1029/2006WR004927.
- 544 [12] Y. W. Tsang, C. F. Tsang, A particle-tracking method for advective
545 transport in fractures with diffusion into finite matrix blocks, *Water*
546 *Resources Research* 37 (3) (2001) 831–835.
- 547 [13] D. Bolster, P. de Anna, D. A. Benson, A. M. Tartakovsky, Incomplete
548 mixing and reactions with fractional dispersion, *Advances in Water Re-*
549 *sources* 37 (2012) 86–93. doi:10.1016/j.advwatres.2011.11.005.
- 550 [14] A. M. Tartakovsky, D. M. Tartakovsky, P. Meakin, Stochastic langevin
551 model for flow and transport in porous media, *Physical Review Letters*
552 101 (2008) 044502.
- 553 [15] Y. Zhang, D. A. Benson, M. M. Meerschaert, E. M. LaBolle, H. P.
554 Scheffler, Random walk approximation of fractional-order multiscaling
555 anomalous diffusion, *Physical Review E* 74 (2006) 026706.
- 556 [16] D. A. Benson, M. M. Meerschaert, Simulation of chemical reaction
557 via particle tracking: Diffusion-limited versus thermodynamic rate-
558 limited regimes, *Water Resources Research* 44 (7) (2008) W12201.
559 doi:10.1029/2008WR007111.
- 560 [17] Y. Ederly, H. Scher, B. Berkowitz, Modeling bimolecular reactions and
561 transport in porous media, *Geophys. Res. Lett.* 36 (2009) L02407.
- 562 [18] D. Fernàndez-Garcia, X. Sanchez Vila, Optimal reconstruction of con-
563 centrations, gradients and reaction rates from particle distributions,
564 *Journal of Contaminant Hydrology* 120-121 (2011) 99–114.
- 565 [19] J. Palanichamy, T. Becker, M. Spiller, J. Kongeter, S. Mohan, Multi-
566 component reaction modelling using a stochastic algorithm, *Computing*
567 *and Visualization in Science* 12 (2) (2009) 51–61. doi:10.1007/s00791-
568 007-0080-y.

- 569 [20] A. F. B. Tompson, A. L. Schafer, R. W. Smith, Impacts of physical
570 and chemical heterogeneity on contaminant transport in sandy porous
571 medium, *Water Resources Research* 32 (4) (1996) 801–818.
- 572 [21] F. Boso, A. Bellin, M. Dumbser, Numerical simulations of solute trans-
573 port in highly heterogeneous formations: A comparison of alternative
574 numerical schemes, *Advances in Water Resources* 52 (2013) 178–189.
- 575 [22] G. Liu, C. Zheng, S. Gorelick, Limits of applicability of the advection-
576 dispersion model in aquifers containing connected high-conductivity
577 channels., *Water Resources Research* 40 (2004) W08308.
- 578 [23] P. Salamon, D. Fernàndez-Garcia, J. Gomez-Hernandez, A review and
579 numerical assessment of the randomwalk particle tracking method, *Jour-
580 nal of Contaminant Hydrology* 87 (1-3) (2006) 277–305.
- 581 [24] M. Riva, A. Guadagnini, D. Fernandez-Garcia, X. Sanchez-
582 Vila, T. Ptak, Relative importance of geostatistical and trans-
583 port models in describing heavily tailed breakthrough curves at the
584 Lauswiesen site, *Journal of Contaminant Hydrology* 101 (2008) 1–13.
585 doi:10.1016/j.jconhyd.2008.07.004.
- 586 [25] P. Salamon, D. Fernàndez-Garcia, J. Gomez-Hernandez, Modeling
587 tracer transport at the made site: The importance of heterogeneity, *Wa-
588 ter Resources Research* 43 (2007) W08404. doi:10.1029/2006WR005522.
- 589 [26] B. W. Silverman, *Density estimation for statistics and data analysis*,
590 Chapman & Hall CRC, 1986.
- 591 [27] R. Andricevic, V. Cvetkovic, Evaluation of risk from contaminants mi-
592 grating by groundwater, *Water Resources Research* 32 (3) (1996) 611–
593 621. doi:10.1029/95WR03530.
- 594 [28] R. Andricevic, Exposure concentration statistics in the subsur-
595 face transport, *Advances in Water Resources* 31 (2008) 714–725.
596 doi:10.1016/j.advwatres.2008.01.007.
- 597 [29] A. Bellin, Y. Rubin, On the use of peak concentration arrival times for
598 the inference of hydrogeological parameters, *Water Resources Research*
599 40 (7) (2004) 1–13. doi:10.1029/2003WR002179.

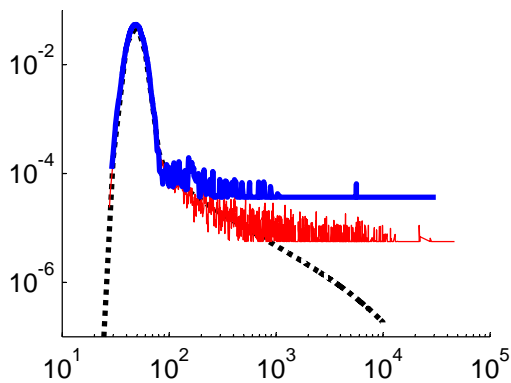
- 600 [30] S. Berglund, V. Cvetkovic, Pump and treat remediation of heteroge-
601 neous aquifers: Effects of rate-limited mass transfer, *Ground Water*
602 33 (4) (1995) 675–685. doi:10.1111/j.1745-6584.1995.tb00324.x.
- 603 [31] G. Dagan, V. Cvetkovic, A. M. Shapiro, A solute flux approach to trans-
604 port in heterogeneous formations: 1. the general framework, *Water Re-*
605 *sources Research* 28 (5) (1992) 1369–1376. doi:10.1029/91WR03086.
- 606 [32] A. M. Shapiro, V. Cvetkovic, Stochastic analysis of solute arrival time in
607 heterogeneous porous media, *Water Resources Research* 24 (10) (1988)
608 1711–1718. doi:10.1029/WR024i010p01711.
- 609 [33] D. Zimmerman et al., A comparison of seven geostatistically based in-
610 verse approaches to estimate transmissivities for modeling advective
611 transport by groundwater flow, *Water Resources Research* 34 (6) (1998)
612 1373–1413. doi:10.1029/98WR00003.
- 613 [34] D. Fernàndez-Garcia, J. J. Gomez-Hernandez, Impact of upscaling
614 on solute transport: Traveltimes, scale dependence of dispersivity,
615 and propagation of uncertainty, *Water Resources Research* 43 (2).
616 doi:10.1029/2005WR004727.
617 URL <http://dx.doi.org/10.1029/2005WR004727>
- 618 [35] R. Haggerty, S. Gorelick, Multiple-rate mass transfer for modeling diffu-
619 sion and surface reactions in media with pore-scale heterogeneity, *Water*
620 *Resour. Res.* 31 (10) (1995) 2383–2400.
- 621 [36] F. P. J. de Barros, Y. Rubin, A risk-driven approach for subsur-
622 face site characterization, *Water Resour. Res.* 58 (2008) W01414.
623 doi:10.1029/2007WR006081.
- 624 [37] D. Pedretti, D. Fernàndez-Garcia, D. Bolster, X. Sanchez-Vila, On the
625 formation of heavy-tailed breakthrough curves during convergent flow
626 tracer tests in three-dimensional heterogeneous sandy aquifers, *Water*
627 *Resources Research*.
- 628 [38] M. Willmann, J. Carrera, X. Sanchez-Vila, Transport upscal-
629 ing in heterogeneous aquifers: What physical parameters con-
630 trol memory functions?, *Water Resour. Res.* 44 (2008) W12437.
631 doi:10.1029/2007WR006531.

- 632 [39] D. Fernàndez-Garcia, G. Llerar-Meza, J. J. Gomez-Hernandez, Upscal-
633 ing transport with mass transfer models: Mean behavior and propaga-
634 tion of uncertainty, *wrr* 45 (2009) W10411. doi:10.1029/2009WR007764.
- 635 [40] A. Moench, Convergent radial dispersion: A laplace transform solution
636 for aquifer tracer testing, *Water Resour. Res.* 25 (3) (1989) 439–447.
- 637 [41] J. Carrera, X. Sanchez-Vila, I. Benet, A. Medina, G. Galarza,
638 J. Guimerà, On matrix diffusion: formulations, solution methods and
639 qualitative effects., *Hydrogeol. J.* 6 (1998) 178–190.
- 640 [42] P. Gouze, T. Le Borgne, R. Leprovost, G. Lods, T. Poidras, P. Pezard,
641 Non-Fickian dispersion in porous media: 1. Multiscale measurements
642 using single-well injection withdrawal tracer tests, *Water Resources Re-*
643 *search* 44 (2008) W06426. doi:10.1029/2007WR006278.
- 644 [43] M. Dentz, A. Castro, Effective transport dynamics in porous media with
645 heterogeneous retardation properties, *Geophys. Res. Lett.* 36 (2009)
646 L03403. doi:10.1029/2008GL036846.
- 647 [44] M. T. Van Genuchten, P. J. Wierenga, Mass transfer studies in sorbing
648 porous media i. analytical solutions, *Soil Sci. Soc. Am. J.* 40 (4) (1976)
649 473–480.
- 650 [45] J. J. Starn, Bagtzoglou, G. A. Robbins, Methods for simulating solute
651 breakthrough curves in pumping groundwater wells, *Computer and Geo-*
652 *sciences* 48 (2012) 244–255. doi:10.1016/j.cageo.2012.01.011.
- 653 [46] W. Hardle, *Smoothing Techniques with Implementation in S*, Springer-
654 Verlag, New York, 1990.
- 655 [47] M. Moroni, N. Kleinfelter, J. H. Cushman, Analysis of dispersion in
656 porous media via matched-index particle tracking velocimetry experi-
657 ments, *Advances in Water Resources* 30 (2007) 1–15.
- 658 [48] A. F. B. Tompson, L. W. Gelhar, Numerical-simulation of solute trans-
659 port in 3-dimensional, randomly heterogeneous porous-media, *Water*
660 *Resources Research* 26 (1990) 2541–2562.
- 661 [49] Y. Rubin, *Applied stochastic hydrogeology*, Oxford University Press,
662 USA, 2003.

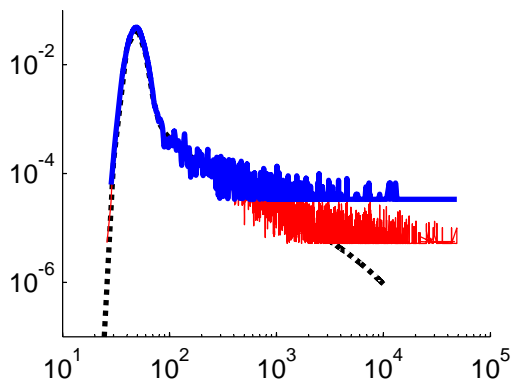
- 663 [50] M. Rosenblatt, Remarks on some nonparametric estimates of a density
664 function, *The Annals of Mathematical Statistics* 27 (3) (1956) pp. 832–
665 837.
- 666 [51] E. Parzen, On estimation of a probability density function and mode,
667 *Ann. Math. Statist.* 33 (1962) 1065–1076.
- 668 [52] J. Engel, E. Herrmann, T. Gasser, An iterative bandwidth selector
669 for kernel estimation of densities and their derivatives, *Nonparametric*
670 *Statistics* 4 (1994) 21–34.
- 671 [53] U. Park, J. S. Marron, Comparison of data-driven bandwidth selectors,
672 *Journal of the American Statistical Association* 85 (409) (1990) 66–72.
673 doi:10.1080/01621459.1990.10475307.
- 674 [54] R. Haggerty, P. C. Reeves, *Stamnt-1 1.0, formulation and user’s guide*,
675 technical report erms 520308, Tech. rep., Sandia National Laboratories,
676 Albuquerque, NM, USA (2002).
- 677 [55] I. S. Abramson, On bandwidth variation in kernel estimates-a square
678 root law, *The Annals of Statistics* 10 (4) (1982) 1217–1223.
- 679 [56] D. Fernández-García, T. H. Illangasekare, H. Rajaram, Conservative
680 and sorptive forced-gradient and uniform flow tracer tests in a three-
681 dimensional laboratory test aquifer, *Water Resources Research* 40 (2004)
682 W10103. doi:10.1029/2004WR003112.
- 683 [57] R. Haggerty, S. A. McKenna, L. C. Meigs, On the late-time behavior
684 of tracer test breakthrough curves, *Water Resour. Res.* 36 (12) (2000)
685 3467–3479.

Global bandwidth (GB) method

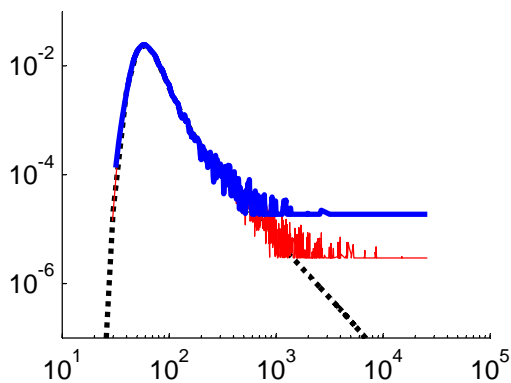
$k=1, \beta=1$ ($Da=0.002$)



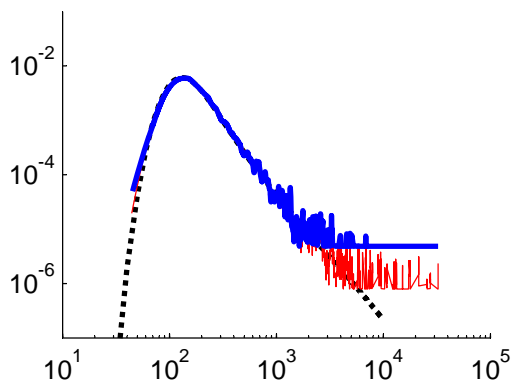
$k=1, \beta=5$ ($Da=0.005$)



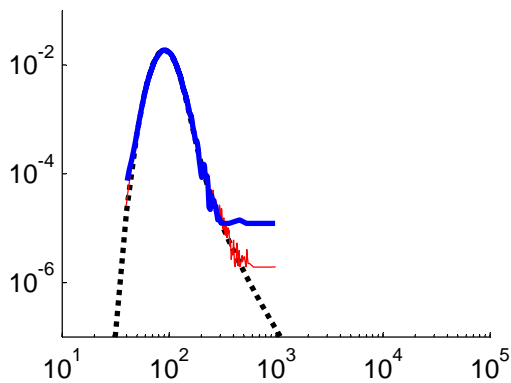
$k=2, \beta=1$ ($Da=0.009$)



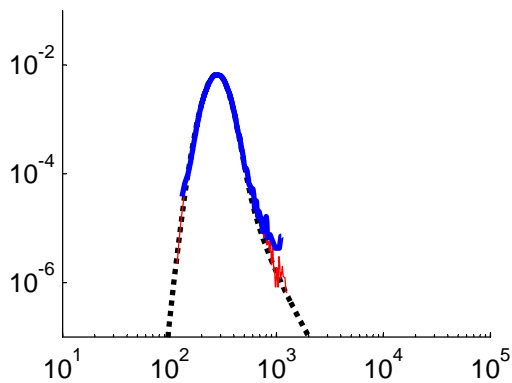
$k=2, \beta=5$ ($Da=0.023$)



$k=3, \beta=1$ ($Da=1.08$)



$k=3, \beta=5$ ($Da=2.71$)

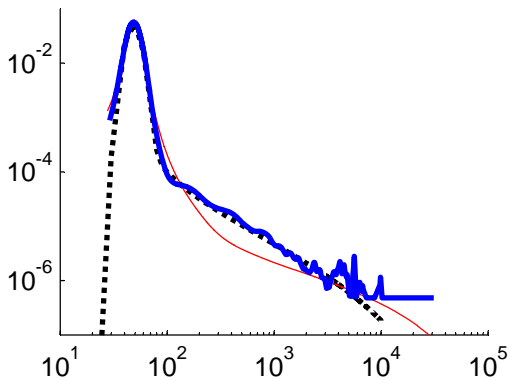


estimated density
time

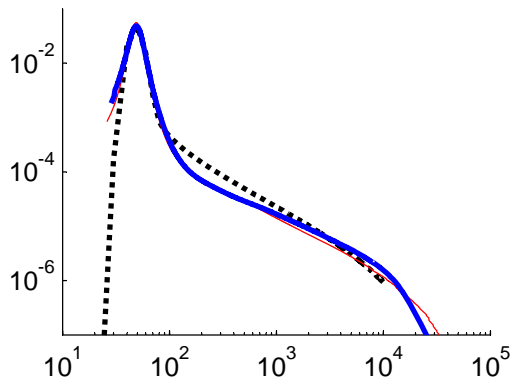
..... Analytical — $n=10^5$ — $n=10^4$

Adaptive bandwidth (AB) method - $\alpha=0.5$

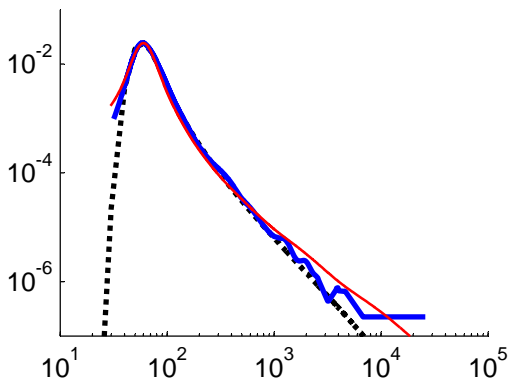
$k=1, \beta=1$ (Da=0.002)



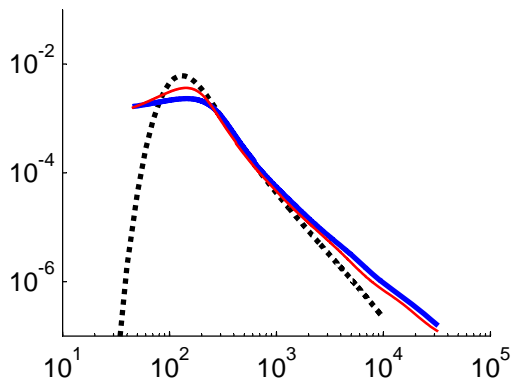
$k=1, \beta=5$ (Da=0.005)



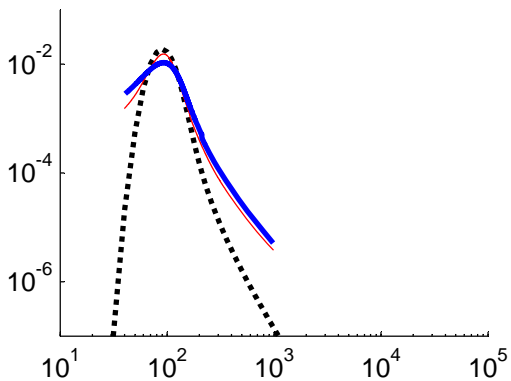
$k=2, \beta=1$ (Da=0.009)



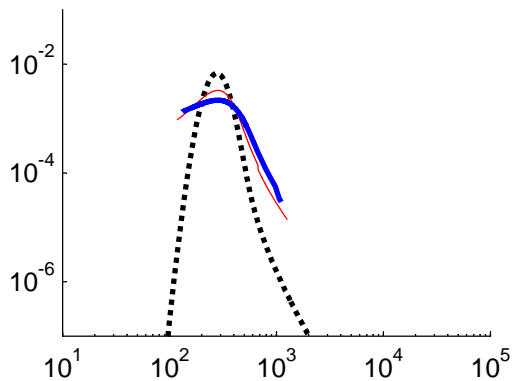
$k=2, \beta=5$ (Da=0.023)



$k=3, \beta=1$ (Da=1.08)



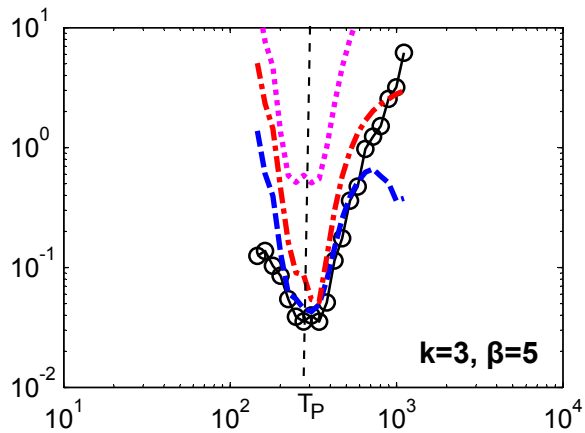
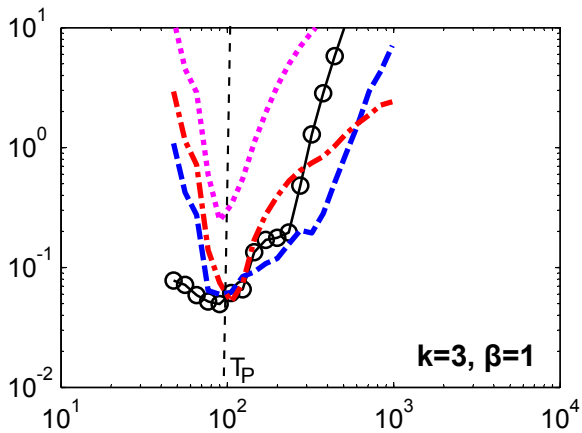
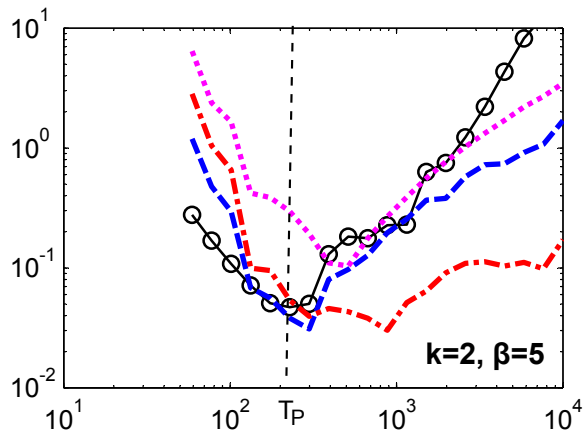
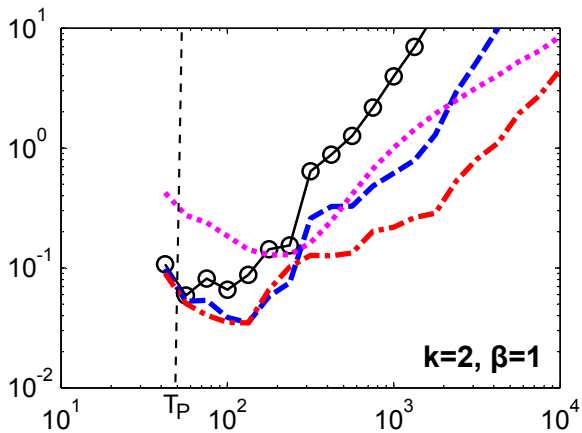
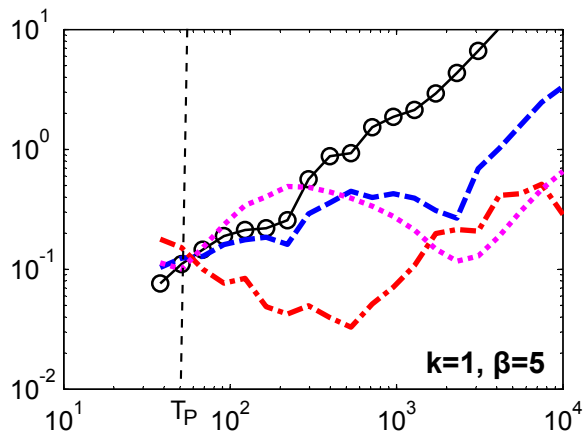
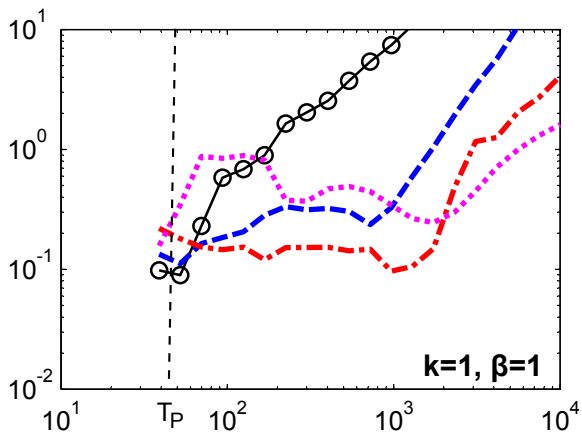
$k=3, \beta=5$ (Da=2.71)



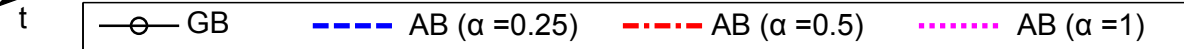
estimated density
time

..... Analytical — $n=10^5$ — $n=10^4$

RELATIVE ERRORS (GB and AB methods , $n=10^4$)



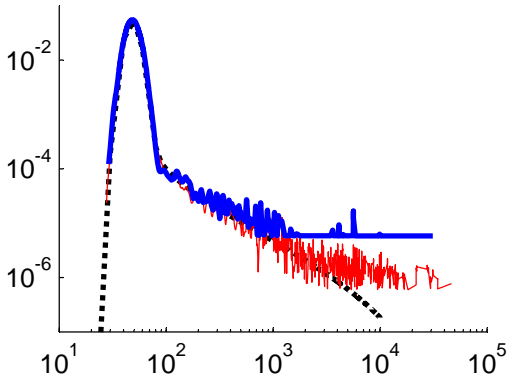
error (ϵ)



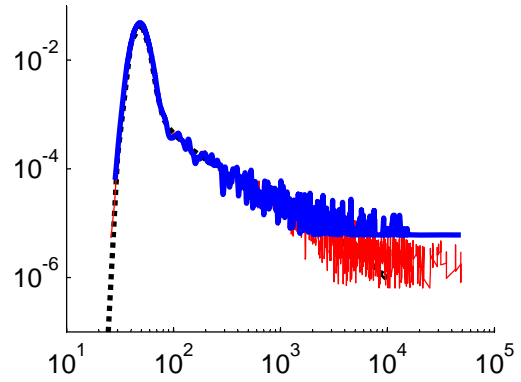
t

Universal adaptive bandwidth (UAB) method - $\alpha=0.25$

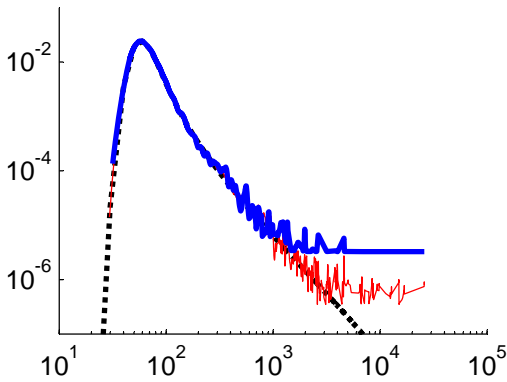
$k=1, \beta=1$ ($Da=0.002$)



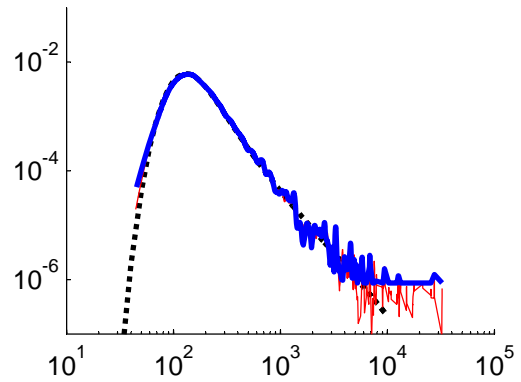
$k=1, \beta=5$ ($Da=0.005$)



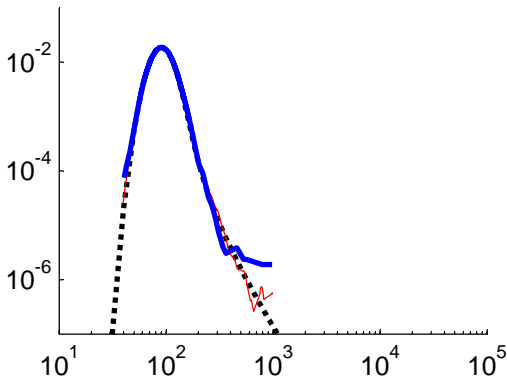
$k=2, \beta=1$ ($Da=0.009$)



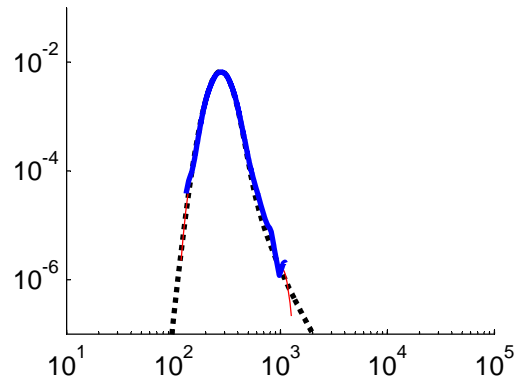
$k=2, \beta=5$ ($Da=0.023$)



$k=3, \beta=1$ ($Da=1.08$)



$k=3, \beta=5$ ($Da=2.71$)

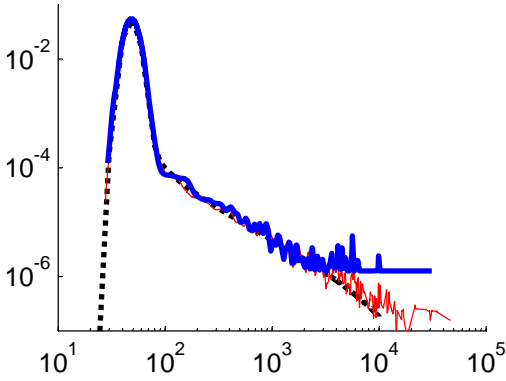


estimated density
time

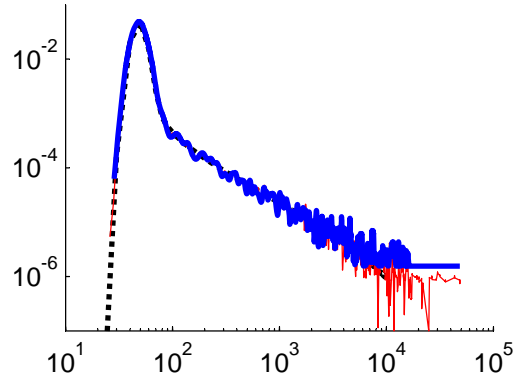
..... Analytical — n=10⁵ — n=10⁴

Universal adaptive bandwidth (UAB) method - $\alpha=0.5$

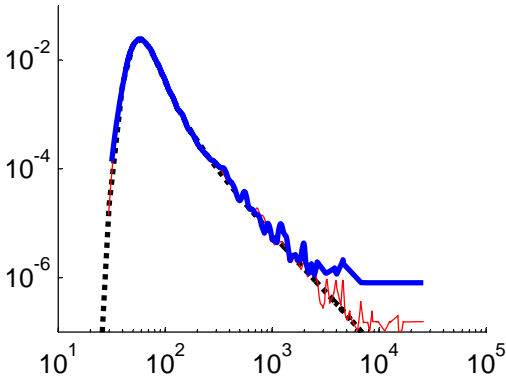
$k=1, \beta=1$ (Da=0.002)



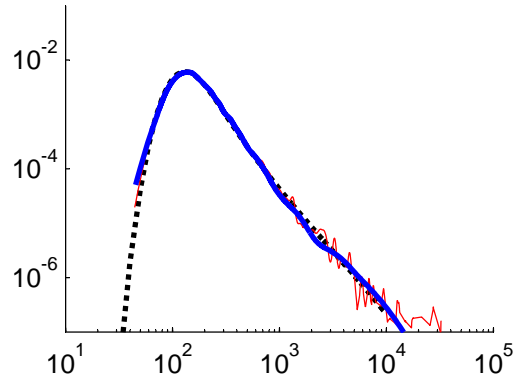
$k=1, \beta=5$ (Da=0.005)



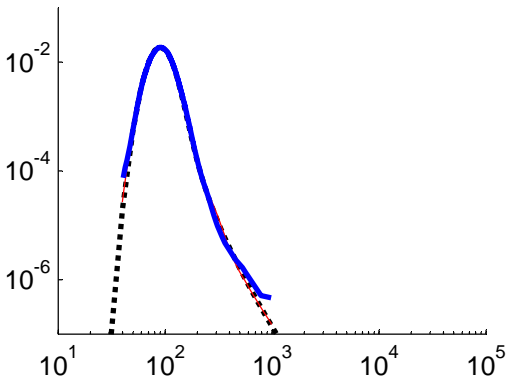
$k=2, \beta=1$ (Da=0.009)



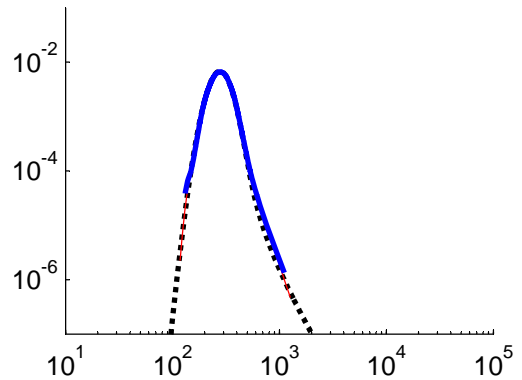
$k=2, \beta=5$ (Da=0.023)



$k=3, \beta=1$ (Da=1.08)



$k=3, \beta=5$ (Da=2.71)

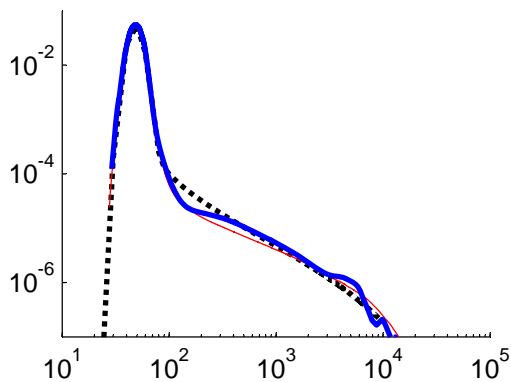


estimated density
time

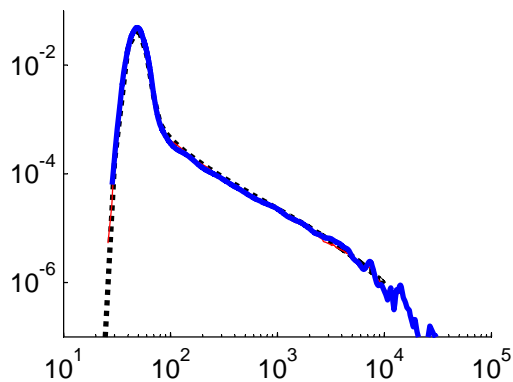
..... Analytical — $n=10^5$ — $n=10^4$

Universal adaptive bandwidth (UAB) method - $\alpha=1$

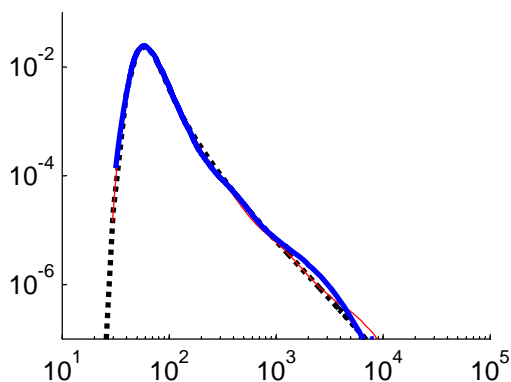
$k=1, \beta=1$ ($Da=0.002$)



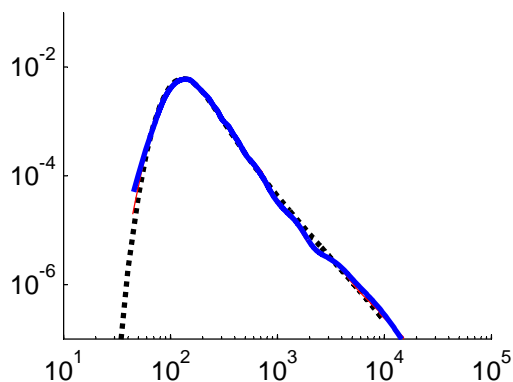
$k=1, \beta=5$ ($Da=0.005$)



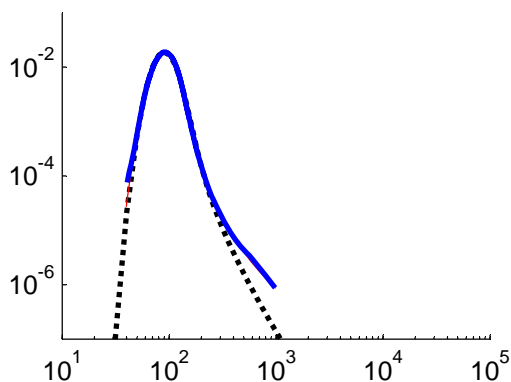
$k=2, \beta=1$ ($Da=0.009$)



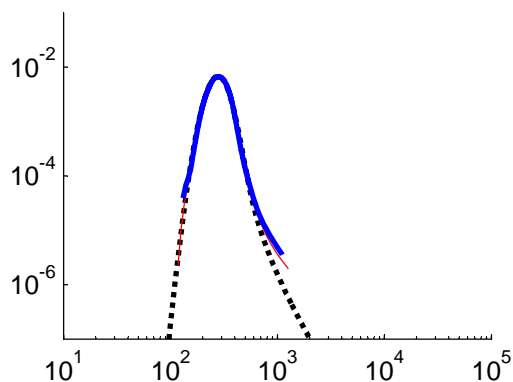
$k=2, \beta=5$ ($Da=0.023$)



$k=3, \beta=1$ ($Da=1.08$)



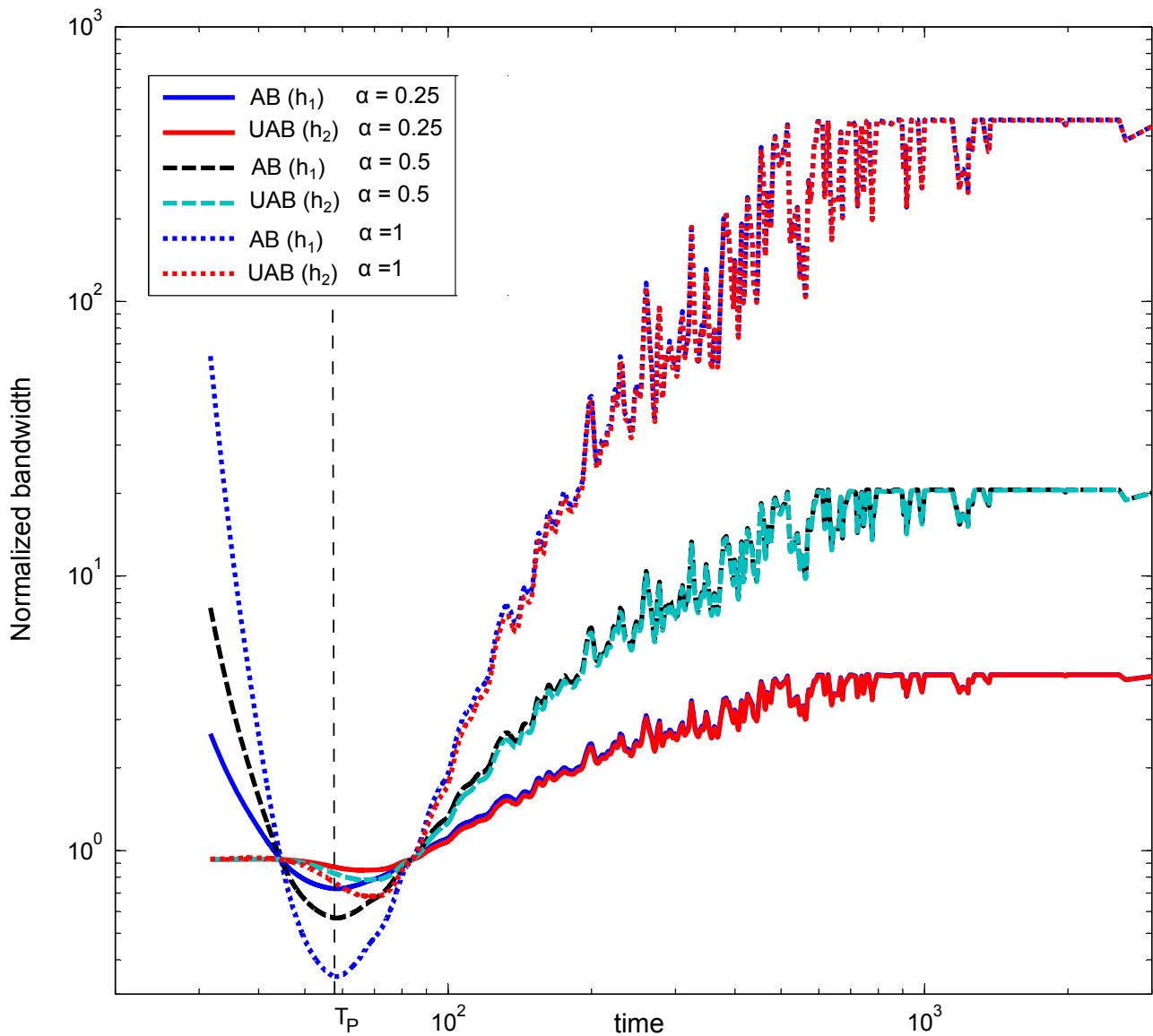
$k=3, \beta=5$ ($Da=2.71$)



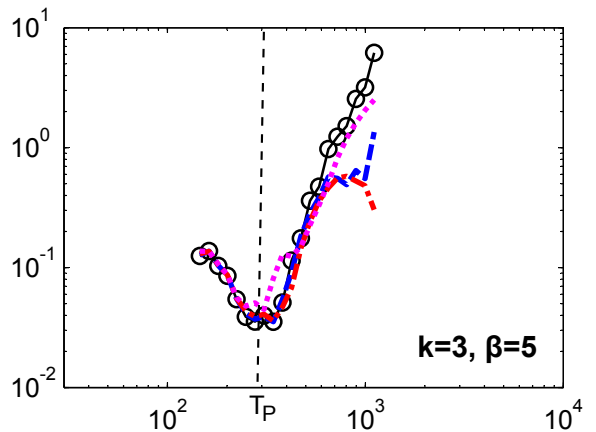
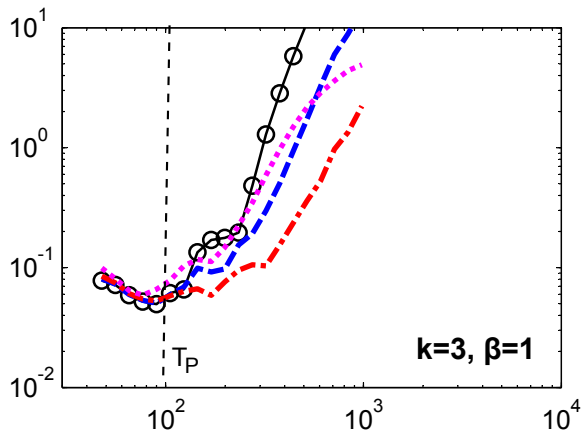
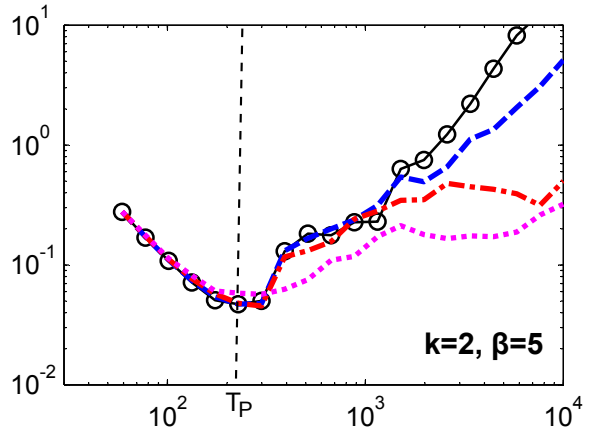
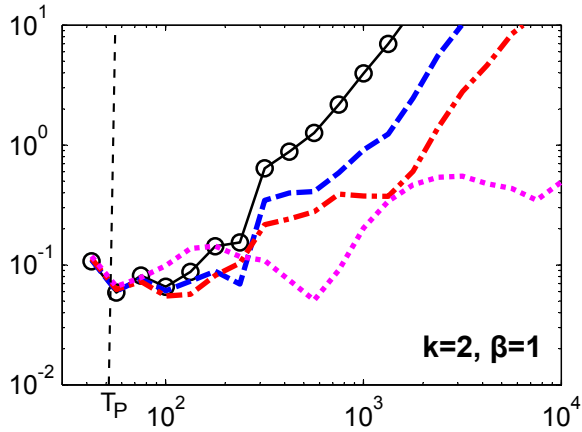
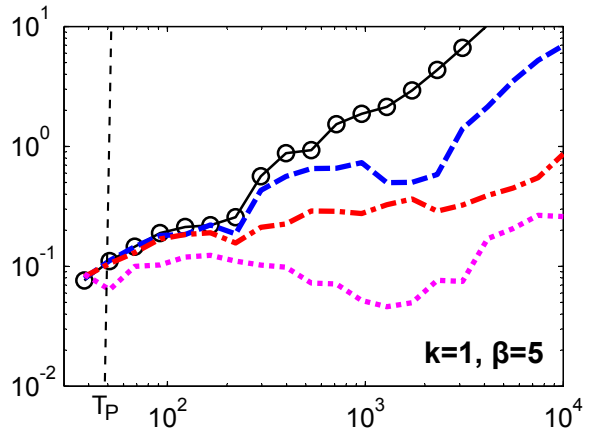
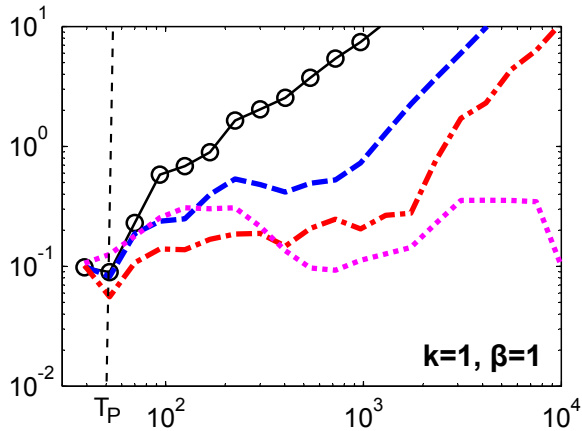
estimated density
time

..... Analytical — $n=10^5$ — $n=10^4$

Shape of the adaptive bandwidth function



RELATIVE ERRORS (GB and UAB methods , $n=10^4$)



error (ϵ)

t



Convergence of the methods

