

Manuscript Number: ATMENV-D-14-01268R2

Title: Multi-wavelength optical determination of Black and Brown Carbon in atmospheric aerosols

Article Type: Research Paper

Keywords: carbonaceous aerosol, light absorption, source apportionment

Corresponding Author: Dr. Dario Massabò, Ph.D.

Corresponding Author's Institution: University of Genoa

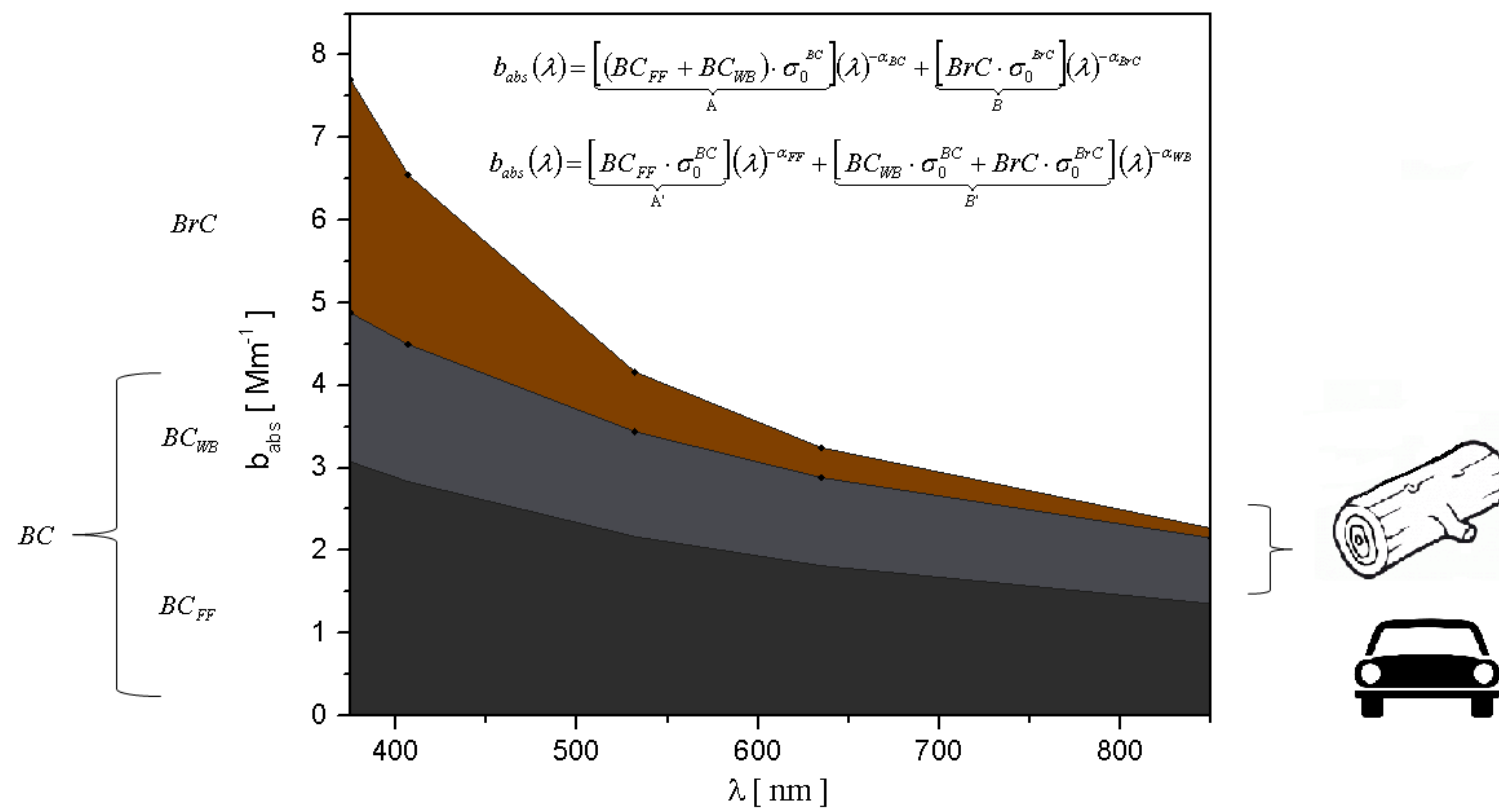
First Author: Dario Massabò, Ph.D.

Order of Authors: Dario Massabò, Ph.D.; Lorenzo Caponi; Vera Bernardoni; Maria Chiara Bove; Paolo Brotto; Giulia Calzolai; Federico Cassola; Massimo Chiari; Mariaelena Fedi; Paola Fermo; Martina Giannoni; Franco Lucarelli; Silvia Nava; Andrea Piazzalunga; Gianluigi Valli; Roberta Vecchi; Paolo Prati

Abstract: In this paper, a new way to apportion the absorption coefficient (babs) of carbonaceous atmospheric aerosols starting from a multi-wavelength optical analysis is shown. This methodology can disentangle and quantify the contribution to total absorption of equivalent black carbon (EBC) emitted by wood burning (EBCWB) and fossil fuel (EBCFF) as well as brown carbon (BrC) due to incomplete combustion. The method uses the information gathered at five different wavelengths in a renewed and upgraded version of the approach usually referred to as Aethalometer model. Moreover, we present the results of an apportionment study of carbonaceous aerosol sources performed in a rural area and in a coastal city, both located in the North-West of Italy. Results obtained by the proposed approach are validated against independent measurements of levoglucosan and radiocarbon. At the rural site the EBCWB and EBCFF relative contributions are about 40% and 60% in winter and 15% and 85% in summer, respectively. At the coastal urban site, EBCWB and EBCFF are about 15% and 85% during fall. The OC contribution to the wood burning source at the rural site results approximately 50% in winter and 10% in summer and about 15% at the coastal urban site in fall. The new methodology also provides a direct measurement of the absorption Ångström exponent of BrC (α_{BrC}) which resulted $\alpha_{\text{BrC}} = 3.95 \pm 0.20$.

Research highlights

- Aerosol light absorption at several λ s due to Black and Brown Carbon is measured
- The value of Ångström exponent of Brown Carbon is directly extracted by raw data
- The new apportionment procedure disentangles fossil and wood burning contributions
- Equivalent Black Carbon and Organic Carbon are separately apportioned
- The procedure is validated against independent Levoglucosan and ^{14}C determination



Multi-wavelength optical determination of black and brown carbon in atmospheric aerosols

D. Massabò^{1,*}, L. Caponi¹, V. Bernardoni², M. C. Bove¹, P. Brotto¹, G. Calzolai³, F. Cassola¹, M. Chiari⁴, M. E. Fedi⁴, P. Fermo⁵, M. Giannoni⁴, F. Lucarelli^{3,4}, S. Nava⁴, A. Piazzalunga⁶, G. Valli², R. Vecchi², P. Prati¹

1: Dept. of Physics, University of Genoa & INFN Via Dodecaneso 33, 16146, Genova, Italy

2: Dept. of Physics, Università degli Studi di Milano & INFN Via Celoria 16, 20133, Milano, Italy

3: Dept. of Physics, University of Florence, Via Sansone 1, 50019, Sesto Fiorentino (FI), Italy

4: INFN – Section of Florence, Via Sansone 1, 50019, Sesto Fiorentino (FI), Italy

5: Dept. of Chemistry, Università degli Studi di Milano, Via Golgi 19, 20133, Milano, Italy

6: Dept. of Environmental and Territorial Sciences, Università degli Studi di Milano-Bicocca, Piazza della Scienza 1, 20122 Milan, Italy

Abstract

In this paper, a new way to apportion the absorption coefficient (b_{abs}) of carbonaceous atmospheric aerosols starting from a multi-wavelength optical analysis is shown. This methodology can disentangle and quantify the contribution to total absorption of equivalent black carbon (EBC) emitted by wood burning (EBC_{WB}) and fossil fuel (EBC_{FF}) as well as brown carbon (BrC) due to incomplete combustion. The method uses the information gathered at five different wavelengths in a renewed and upgraded version of the approach usually referred to as Aethalometer model. Moreover, we present the results of an apportionment study of carbonaceous aerosol sources performed in a rural area and in a coastal city, both located in the North-West of Italy. Results obtained by the proposed approach are validated against independent measurements of levoglucosan and radiocarbon. At the rural site the EBC_{WB} and EBC_{FF} relative contributions are about 40% and 60% in winter and 15% and 85% in summer, respectively. At the coastal urban site, EBC_{WB} and EBC_{FF} are about 15% and 85% during fall. The OC contribution to the wood burning source at the rural site results approximately 50% in winter and 10% in summer and about 15% at the coastal urban site in fall. The new methodology also provides a direct measurement of the absorption Ångström exponent of BrC (α_{BrC}) which resulted $\alpha_{\text{BrC}} = 3.95 \pm 0.20$.

Keywords: carbonaceous aerosol, light absorption, source apportionment

* Corresponding Author: massabo@ge.infn.it

Introduction

36

37 Carbonaceous aerosols play an important role in environmental issues like air quality, human
38 health and global climate change. Although the classification of carbonaceous aerosol
39 components is still under debate (Pöschl, 2003), total carbon (TC) is generally divided in
40 black carbon (BC), organic carbon (OC) and carbonate carbon (CC).

41 Amongst atmospheric aerosols, BC is considered the most efficient light-absorber in the
42 visible spectrum (Bond et al., 2013; and reference therein) with a weak dependence on
43 wavelength (λ) (Moosmüller et al., 2009). Another light-absorbing component of
44 carbonaceous aerosols is the so-called brown carbon (BrC) (Andreae & Gelencsér, 2006;
45 Pöschl, 2003), the fraction of organic carbon with increased absorbance in the blue and
46 ultraviolet (UV) region of the solar spectrum (Moosmüller et al., 2011). Carbonaceous light-
47 absorbing particles are typically emitted by incomplete combustion of fossil fuels related to
48 traffic, industrial processes and domestic heating as well as by biomass burning.

49 It is worthy to note that beyond carbonaceous aerosols, also other aerosol components show
50 strong light-absorbing properties like iron oxides in mineral dust particles (Linke et al.,
51 2006).

52 The spectral dependence of the aerosol absorption coefficient (b_{abs}) is generally described by
53 the power-law relationship $b_{\text{abs}}(\lambda) \propto \lambda^{-\text{AAE}}$ where the AAE is the Ångström absorption
54 exponent (Moosmüller et al., 2011). In literature works AAE has been shown to be sensitive
55 to aerosol chemical composition but also to particle size and morphology (e.g. Kirchstetter et
56 al., 2004; Lewis et al. 2008; Utry et al., 2014). In a large number of cases, it has been
57 exploited as a chemically selective parameter useful to identify the aerosol origin and
58 apportion sources for different carbonaceous aerosols (Sandradewi et al., 2008; Ajtai et al.,
59 2010; Favez et al., 2010; Flowers et al., 2010; Filep et al. 2013; Utry et al., 2013);
60 nevertheless, Utry et al. (2014) claim that the assessment of aerosol microphysical properties
61 is needed to retrieve more accurate results on the aerosol absorption properties. AAE values
62 around 1 have been reported for BC and up to 9.5 for BrC (Lack & Langridge, 2013).

63 An advantage of the AAE determination in aerosol samples by multi- λ techniques is the
64 possibility of performing on-line source apportionment studies as done by many authors in
65 recent years to evaluate woodsmoke and traffic contributions adopting the so-called
66 Aethalometer model (Sandradewi et al., 2008; Favez et al., 2010) thus avoiding time
67 consuming laboratory analyses.

68 At the state of the art the measurement of light absorption is still challenging (Andreae, 2001;
69 Moosmüller et al., 2011), notwithstanding filter-based on-line techniques (e.g. the
70 Aethalometer; the Particle Soot Absorption Photometer; the Multi Angle Absorption
71 Photometer, among others) are widespread but – with the exception of the Aethalometer –
72 multi- λ analysis is generally not implemented. There are some important drawbacks to be
73 addressed in order to get reliable values from these filter-based instruments as it is well
74 known that they are affected by measurement and sampling artifacts (e.g. effects due to
75 multiple scatterings, to particle shadowing due to filter loading, absorption of organics; Bond
76 et al., 1999; Collaud Coen et al., 2010; Vecchi et al., 2014; among others). Although not very
77 widespread yet, photacoustic spectroscopy operated at multi- λ (Lewis et al., 2008; Ajtai et al.,
78 2010; Flowers et al., 2010) is currently the only method capable to overcome the above
79 mentioned drawbacks in absorption measurements.

80 At the University of Genoa a Multi-Wavelength Absorbance Analyzer (MWAA) has been
81 recently developed (Massabò et al., 2013) basing on the single- λ Multi Angle Absorption
82 Photometer concept (MAAP, Petzold & Schönlinner, 2004; Petzold et al., 2005). Such
83 instrumentation measures both transmitted and scattered light in the forward and back
84 hemispheres thus reducing the cross-sensitivity to aerosol scattering components and filter
85 loading effects (Müller et al., 2011). This approach does not need a posteriori data corrections
86 necessary when attenuation measurements only are performed (e.g. Collaud Coen et al.,
87 2010): such corrections are typically composition dependent and prevent real-time accurate
88 source apportionment.

89 In this work, we present a new apportionment methodology together with an original data
90 reduction approach developed using an up-graded version of the MWAA serving reliable b_{abs}
91 data at different wavelengths. From the direct apportionment of BC and BrC spectral
92 absorption properties, the contributions of fossil fuels (FF) and wood burning (WB) to the
93 carbonaceous aerosols concentration can be disentangled.

94

95 **1. The Multi-Wavelength Absorbance Analyzer (MWAA)**

96 *Set-up*

97 A detailed description of the original MWAA set-up is given in Massabò et al. (2013) and in
98 the following only major changes and upgrades will be reported.

99 The MWAA is basically composed by light emitting sources, an automatized sample-
100 changer, and 4 low-noise UV-enhanced photodiodes. In the original configuration three low-
101 power laser diodes ($\lambda = 407, 635, 850\text{nm}$) were displaced on a slide and manually aligned

102 thanks to mechanical benchmarks. In the new configuration, two laser diodes with $\lambda =$
103 375nm and 532nm (World Star Tech) have been added. A motorized stage has been added to
104 interchange the laser sources thus improving the system stability and reproducibility and
105 facilitating the analysis of many samples.

106 The 5- λ b_{abs} measurements are exploited in the model proposed in this paper to retrieve more
107 accurate results (see §3). In particular, measurement at UV wavelength is useful because the
108 absorption properties of atmospheric aerosols at this λ are generally poorly known and brown
109 carbon is expected to strongly absorb in this range (Andreae & Gelencsér, 2006; Kirchstetter
110 & Thatcher, 2012).

111 In the final configuration, the MWAA can perform 5-wavelength analysis of 16
112 filters/samples per session in less than 90 minutes measuring each filter in 64 different
113 points, each $\sim 1 \text{ mm}^2$ wide.

114

115 *Calculation of the aerosol absorption coefficient*

116 To derive the b_{abs} at each measured λ , the MWAA partially follows the approach reported by
117 Petzold & Schönlinner (2004) and implemented in the MAAP. From the measurement of the
118 light transmitted and scattered at fixed angles, the light angular distributions in the forward
119 and in the back hemispheres are retrieved using analytical functions. Once the light
120 distribution is obtained in both hemispheres (for details on the MWAA see Massabò et al.,
121 2013), a radiative transfer model taking into account the multiple scattering effects occurring
122 within the particle-filter system is applied (Hänel, 1987 and 1994). The model gives the two
123 parameters needed to calculate the sample absorbance (ABS – the fraction of light absorbed
124 by the loaded filter), i.e. the total optical thickness (τ) and the aerosol-filter layer single
125 scattering albedo (SSA). These parameters are linked to ABS through the relationship $\text{ABS} =$
126 $\tau (1-\text{SSA})$. Finally, b_{abs} is given by $b_{\text{abs}} = \text{ABS} \cdot \frac{A}{V}$, where A is the active surface filter area
127 and V is the volume of sampled air.

128

129 **2. Field campaigns and laboratory analyses**

130 *Samples collection*

131 PM10 aerosol samples were collected at two different locations in Liguria (Italy): a regional
132 background and an urban background site. The regional background monitoring site was
133 placed in a small village (Propata, 44°33'52.93''N, 9°11'05.57''E, 970 m a.s.l., population
134 160 inhabitants) in the Ligurian Appennines where wood burning is expected to be a major

135 aerosol source especially during wintertime as it is used for both domestic heating and
136 cooking. Due to its peculiar position, Propata can be occasionally impacted by pollution
137 advection from the Po valley as well as from the coastal area at South. The urban background
138 monitoring site was 2 km far from the Genoa city centre (44°24'08.93''N, 8°58'18.17''E, 60
139 m a.s.l., population 600,000 inhabitants). The sampling site was located on the terrace of the
140 Physics Department and it can be considered as representative of a maritime urban
141 background station because not directly influenced by local pollution sources.

142 In Propata the sampling covered different periods: February - July 2013 and November 2013
143 - January 2014. 48-hour PM₁₀ samples (120 in total) were collected on quartz-fibre filters
144 (Pall, 2500QAO-UP, 47 mm diameter) using a low-volume sampler (38.3 l min⁻¹ by TCR
145 Tecora, Italy). Additional PM₁₀ samples were collected for radiocarbon analysis on quartz-
146 fibre filters (Pall, QAT-UP, 150 mm diameter) using a high-volume sampler (500 l min⁻¹).
147 This sampling was carried out in March-April 2013. Each sampling lasted about 6 days and 4
148 samples were collected overall.

149 In Genoa 24-hour PM₁₀ was sampled for two weeks (October 31 - November 13 2013) by a
150 low-volume sampler (38.3 l min⁻¹ by TCR Tecora, Italy). Aerosol particles were collected on
151 quartz-fibre filters (Pall, 2500QAO-UP, 47 mm diameter).

152 The quartz-fibre filters were never heat-treated before sampling; any possible contamination
153 was assessed in each batch before sampling (the maximum OC contamination was 1.7 ± 0.3
154 $\mu\text{g cm}^{-2}$). Field blank filters were used to monitor any possible further contaminations.
155 Moreover, in this work we decided to neglect the possible effect of sampling artefacts due to
156 organics on light absorption measurements when using quartz-fibre filters as shown by
157 Vecchi et al. (2014) because it has been considered here not to alter the approach described
158 in §3.

159 *Laboratory analyses*

160 Filters were weighed before and after the sampling in an air-conditioned room. After
161 weighing, low-volume samples were analyzed by MWAA to retrieve b_{abs} at five different
162 wavelengths. EC and OC were determined on one punch (1.5 cm²) of the quartz-fibre filter
163 by a Thermal Optical Transmittance (TOT) instrument (Sunset Lab Inc.) using the
164 EUSAAR_2 protocol (Cavalli et al., 2010). In addition, levoglucosan – a well known marker
165 for wood burning (Simoneit et al., 1999) – was determined by High Performance Anion
166 Exchange Chromatography coupled with Pulsed Amperometric Detection on a portion of the
167 same quartz-fibre filter (Piazzalunga et al., 2010).

168 TC radiocarbon analyses on high-volume samples were performed at the Accelerator Mass
169 Spectrometry (AMS) facility of the INFN-LABEC laboratory of Florence (Italy) (see Fedi et
170 al., 2013 for details). Sample preparation for AMS analysis was carried out using a sample
171 preparation line suitably set up for aerosol samples (Calzolari et al., 2011), following the TC
172 sample preparation procedure described in Bernardoni et al. (2013).

173 The measured $^{14}\text{C}/^{12}\text{C}$ was corrected for the background signal and then for isotopic
174 fractionation according to $^{13}\text{C}/^{12}\text{C}$ measured in the accelerator. Data were normalized to the
175 isotopic ratio obtained for the oxalic acid II standard NIST 4990C and the results were
176 expressed as fraction of modern carbon (f_m) in the sample. It is noteworthy that nuclear
177 weapon tests in the '50s and '60s led to the increase of ^{14}C to ^{12}C ratio in biologic material.
178 Thus, f_m must be corrected for reference values of the ^{14}C content for the sampling period
179 ($f_{m,\text{mod}}$) to correctly apportion modern sources. In this work, $f_{m,\text{mod}}$ was set to 1.08, assuming
180 equivalent contributions by wood combustion and biogenic sources and using 1.116 and
181 1.036 as representative f_m values for the wood burnt and the biogenic material, respectively
182 (Zotter et al., 2014). The relative contribution from modern (non-fossil) sources was

183 estimated as $f_{NF} = \frac{f_m}{f_{m,\text{mod}}}$.

184

185 **3. Multi-wavelength analysis**

186 *Optical apportionment*

187 In this work, a new source apportionment model (MWAA approach) based on the
188 measurement of b_{abs} at five wavelengths was developed. The MWAA approach exploits the
189 information provided by the 5- λ measurements to obtain directly the BrC AAE (α_{BrC}) and the
190 BrC absorption coefficient ($b_{\text{abs}}^{\text{BrC}}$) at each measured λ . These are innovative features
191 compared to the Aethalometer model (Sandradewi et al., 2008; Favez et al., 2010) which uses
192 measurements at two λ only (even with the 7- λ instrument) and provides the total
193 contribution to b_{abs} due to fossil fuels ($b_{\text{abs,FF}}$) and wood burning ($b_{\text{abs,WB}}$) without information
194 on the species (i.e. BC or BrC) responsible for such contributions.

195 The minimization algorithm used in the MWAA approach joined to TOT measurements
196 allows also to apportion the contributions of FF and WB to the EC (EC_{FF} and EC_{WB} ,
197 respectively) and, although with some further assumptions, to the OC (OC_{FF} and OC_{WB} ,
198 respectively).

199 The MWAA approach starts from two different decompositions of $b_{\text{abs}}(\lambda)$.

200 In the first case, $b_{\text{abs}}(\lambda)$ is assumed to be the sum of the absorption coefficients of BC
 201 ($b_{\text{abs}}^{\text{BC}}(\lambda)$, regardless of its FF or WB origin), and BrC ($b_{\text{abs}}^{\text{BrC}}(\lambda)$), as follows

$$203 \quad b_{\text{abs}}(\lambda) = b_{\text{abs}}^{\text{BC}}(\lambda) + b_{\text{abs}}^{\text{BrC}}(\lambda) \quad (1)$$

204
 205 Furthermore, $b_{\text{abs}}^{\text{BC}}(\lambda)$ and $b_{\text{abs}}^{\text{BrC}}(\lambda)$ are assumed to be λ -dependent following the general
 206 relationship $b_{\text{abs}}(\lambda) \propto \lambda^{-\alpha}$, where α is different for BC and BrC (α_{BC} and α_{BrC} , respectively).

207 Thus, eq.(1) can be written as:

$$209 \quad \frac{b_{\text{abs}}^{\text{BC}}(\lambda_1)}{b_{\text{abs}}^{\text{BC}}(\lambda_{\text{ref}})} = \left(\frac{\lambda_1}{\lambda_{\text{ref}}} \right)^{-\alpha_{\text{BC}}} \quad \frac{b_{\text{abs}}^{\text{BrC}}(\lambda_1)}{b_{\text{abs}}^{\text{BrC}}(\lambda_{\text{ref}})} = \left(\frac{\lambda_1}{\lambda_{\text{ref}}} \right)^{-\alpha_{\text{BrC}}} \quad (2)$$

210
 211 where λ_{ref} can be arbitrarily chosen.

212 In the second case, the decomposition approach is the same as in the Aethalometer model:

$$213 \quad b_{\text{abs}}(\lambda) = b_{\text{abs,FF}}(\lambda) + b_{\text{abs,WB}}(\lambda) \quad (3)$$

214 where $b_{\text{abs,FF}}$ and $b_{\text{abs,WB}}$ are the contributions from FF and WB to the total b_{abs} . This
 215 decomposition assumes that FF and WB are the only sources of light absorbing species at the
 216 sampling site. In this case, it is assumed a λ -dependence of b_{abs} related to the source of the
 217 absorbing aerosol as follows:

$$219 \quad \frac{b_{\text{abs,FF}}(\lambda)}{b_{\text{abs,FF}}(\lambda_{\text{ref}})} = \left(\frac{\lambda}{\lambda_{\text{ref}}} \right)^{-\alpha_{\text{FF}}} \quad \frac{b_{\text{abs,WB}}(\lambda)}{b_{\text{abs,WB}}(\lambda_{\text{ref}})} = \left(\frac{\lambda}{\lambda_{\text{ref}}} \right)^{-\alpha_{\text{WB}}} \quad (4)$$

220
 221 where α_{FF} and α_{WB} are the AAE representative for FF and WB aerosol and λ_{ref} can be
 222 arbitrarily chosen.

223
 224 Moreover, the MWAA approach is based on the following assumptions:

- 225 a) Wood burning is the only source of BrC;
- 226 b) BC_{FF} and BC_{WB} have the same AAE (α_{BC}), disregarding the emission source;
- 227 c) BC and BrC have different spectral dependences, i.e. different AAE (α_{BC} and α_{BrC} ,
 228 respectively);
- 229 d) Fossil fuels are assumed not to contribute to BrC, thus AAE for FF aerosol is assumed to
 230 be the one for BC ($\alpha_{\text{FF}} = \alpha_{\text{BC}}$). α_{FF} value was set as explained in the following;

231 e) α_{WB} was set to a fixed value as explained in the following.

232

233 Equations (1) and (2) can be joined and rewritten as:

$$\begin{aligned}
 234 \quad b_{abs}(\lambda) &= \left[(BC_{FF} + BC_{WB}) \cdot MAC_{\lambda_{ref1}}^{BC} \right] \left(\frac{\lambda}{\lambda_{ref1}} \right)^{-\alpha_{BC}} + \left[BrC \cdot MAC_{\lambda_{ref2}}^{BrC} \right] \left(\frac{\lambda}{\lambda_{ref2}} \right)^{-\alpha_{BrC}} = \\
 &= \underbrace{\left[(BC_{FF} + BC_{WB}) \cdot \frac{MAC_{\lambda_{ref1}}^{BC}}{\lambda_{ref1}^{-\alpha_{BC}}} \right]}_A \lambda^{-\alpha_{BC}} + \underbrace{\left[BrC \cdot \frac{MAC_{\lambda_{ref2}}^{BrC}}{\lambda_{ref2}^{-\alpha_{BrC}}} \right]}_B \lambda^{-\alpha_{BrC}} = \quad (5) \\
 &= \underbrace{\left[(BC_{FF} + BC_{WB}) \cdot \sigma_0^{BC} \right]}_A \lambda^{-\alpha_{BC}} + \underbrace{\left[BrC \cdot \sigma_0^{BrC} \right]}_B \lambda^{-\alpha_{BrC}}
 \end{aligned}$$

235 Where:

236 • $MAC_{\lambda_{ref1}}^{BC}$ and $MAC_{\lambda_{ref2}}^{BrC}$ are the mass specific absorption coefficients (in $[m^2 g^{-1}]$) at
 237 arbitrarily chosen reference wavelengths (λ_{ref1} and λ_{ref2}) for BC and BrC, respectively;

238 it is noteworthy that $\sigma_0^{BC} = \frac{MAC_{\lambda_{ref1}}^{BC}}{\lambda_{ref1}^{-\alpha_{BC}}}$ and $\sigma_0^{BrC} = \frac{MAC_{\lambda_{ref2}}^{BrC}}{\lambda_{ref2}^{-\alpha_{BrC}}}$ only depend on BC and

239 BrC properties, respectively.

240 • BC_{FF} and BC_{WB} are the concentrations of BC emitted by FF and WB, respectively;

241 • BrC is the concentration of brown carbon. The indication of the BrC source is omitted
 242 as we assume it comes only from WB.

243 With the same formalism introduced for eq.(5), eq.(3) and (4) are joined and rewritten as:

244

$$245 \quad b_{abs}(\lambda) = \underbrace{\left[BC_{FF} \cdot \sigma_0^{BC} \right]}_{A'} \lambda^{-\alpha_{FF}} + \underbrace{\left[BC_{WB} \cdot \sigma_0^{BC} + BrC \cdot \sigma_0^{BrC} \right]}_{B'} \lambda^{-\alpha_{WB}} \quad (6)$$

246

247 Many works (Bond & Bergstrom, 2006; Moosmüller et al., 2011; and references therein)

248 show that aerosol produced by fossil fuel combustion has a typical value of $\alpha_{FF} \approx 1.0$,

249 considering laboratory as well as field experiments, thus meaning an inverse proportionality

250 between b_{abs} and λ . This further supports our assumption that $\alpha_{FF} = \alpha_{BC}$, that means that OC_{FF}

251 is expected not to contribute to $b_{abs,FF}$ (i.e. no BrC is emitted by FF combustion). Indeed, in

252 samples where negligible WB contribution is expected because of low levoglucosan

253 concentration, a $\lambda^{-\alpha_{exp}}$ function with $0.9 < \alpha_{exp} < 1.1$ suitably fits the $5-\lambda$ determined b_{abs} .

254 Considering α_{BC} , it is widely accepted that the absorption cross-section for “pure” BC in the

255 atmosphere varies as λ^{-1} , i.e. the imaginary part of the refractive index does not depend on λ
 256 (Bond & Bergström, 2006; Lack & Langridge, 2013; and references therein). Taking into
 257 account the previous considerations, in this study we fixed $\alpha_{FF} = 1.0$ (see further discussion at
 258 §4).

259 In the case of woodsmoke, α_{WB} values are generally in the range of 0.9-2.2 (Harrison et al.,
 260 2013; and therein cited literature) depending on several factors as air mass aging and the type
 261 of wood burnt as well as the specific wavelength range where the α_{WB} values are calculated.
 262 In this work, we fixed $\alpha_{WB} = 1.8$ which gave a fairly good agreement with independent ^{14}C
 263 measurements (see discussion at §4). Once α_{BC} , α_{FF} , and α_{WB} have been set, the system is
 264 numerically solved by fitting the two eq.(5) and (6) separately and using a MINUIT χ^2
 265 minimization routine (James, 1978) home-written as a C⁺⁺ program (ROOT package; Brun &
 266 Rademakers, 1997). The minimization program fits the 5- λ b_{abs} measurements performed by
 267 MWAA following eq.(5) and eq.(6) to obtain A, B, A', B' and α_{BrC} for each sample. It is
 268 noteworthy the multi- λ measurements allow to run the proposed model and provide an
 269 accurate fitting. The mean α_{BrC} values extracted for the Propata and Genoa datasets are $3.89 \pm$
 270 0.18 and 4.02 ± 0.19 , respectively (quoted uncertainties are the standard deviation of the two
 271 distributions). Values of α_{BrC} up to 9.5 have been reported for wavelength pairs 400 and
 272 700nm (Lack & Langridge, 2013; and references therein). The values obtained in this work
 273 are in good agreement with the findings of Yang et al. (2009) who reported $\alpha_{BrC} = 3.5$ (for
 274 wavelength pairs 470 and 660nm).

275 Considering eq.(5) and (6), the following relations can be derived:

$$276 \quad \begin{cases} A - A' = BC_{WB} \sigma_0^{BC} \\ A' = BC_{FF} \sigma_0^{BC} \\ B = BrC \sigma_0^{BC} \end{cases}$$

277 and the corresponding λ dependences lead to:

$$278 \quad \begin{cases} b_{\text{abs},WB}^{BC}(\lambda) = BC_{WB} \sigma_0^{BC} \lambda^{-\alpha_{BC}} = (A - A') \lambda^{-\alpha_{BC}} \\ b_{\text{abs},FF}^{BC}(\lambda) = BC_{FF} \sigma_0^{BC} \lambda^{-\alpha_{BC}} = A' \lambda^{-\alpha_{BC}} \\ b_{\text{abs}}^{BrC}(\lambda) = BrC \sigma_0^{BrC} \lambda^{-\alpha_{BrC}} = B \lambda^{-\alpha_{BrC}} \end{cases} \quad (7)$$

279 thus the source-dependent (FF or WB) light absorption contributions to BC and BrC
 280 ($b_{\text{abs},FF}^{BC}(\lambda)$, $b_{\text{abs},WB}^{BC}(\lambda)$, and $b_{\text{abs}}^{BrC}(\lambda)$) can be obtained from the results of the minimization
 281 algorithm. Please note that BrC is assumed to be emitted only by WB, thus the source
 282 indication is omitted in b_{abs}^{BrC} .

283 In Figure 1, the mean $b_{abs}(\lambda)$ source apportionment at the rural site of Propata is shown.
 284 Moreover, the optical apportionment for the Propata winter campaign is shown as an
 285 example in Figure 2 at $\lambda = 850\text{nm}$ and $\lambda = 375\text{nm}$. As expected, the Figures show that b_{abs}^{BrC}
 286 is very low at the infrared (IR) wavelength whereas it explains up to 50% of the total light
 287 absorption in the case of UV. Although $b_{abs}^{BrC}(850\text{nm})$ is generally low, it varies greatly from
 288 one day to another reaching values up to 11% of the total $b_{abs}(850\text{nm})$.

289

290 *Mass apportionment: equivalent black carbon (EBC)*

291 The MWAA approach described so far quantifies the three main contributors to total b_{abs}
 292 ($b_{abs,FF}^{BC}$, $b_{abs,WB}^{BC}$, and b_{abs}^{BrC}) at five different λ , so that:

293

$$294 \quad b_{abs}(\lambda) = b_{abs,FF}^{BC}(\lambda) + b_{abs,WB}^{BC}(\lambda) + b_{abs}^{BrC}(\lambda) \quad (8)$$

295

296 Moreover, following the approach in eq.(5) a wavelength-dependent mass-absorption cross-
 297 section for BC can be introduced as $MAC^{BC}(\lambda) = \sigma_0^{BC} \lambda^{-\alpha_{BC}} = \sigma_0^{BC} \lambda^{-1}$. This allows the
 298 evaluation of the equivalent black carbon in atmosphere (EBC) as the sum of EBC from FF
 299 and WB (EBC_{FF} and EBC_{WB} , respectively):

$$300 \quad EBC = EBC_{FF} + EBC_{WB} = \frac{b_{abs,FF}^{BC}(\lambda) + b_{abs,WB}^{BC}(\lambda)}{MAC^{BC}(\lambda)} \quad (9)$$

301 With the further assumption of equivalence between the EBC in atmosphere and the EC
 302 determined by thermal-optical analysis ($EBC_{FF} = EC_{FF}$ and $EBC_{WB} = EC_{WB}$), the following
 303 relationships hold at every λ :

$$304 \quad \begin{aligned} EC_{FF} &= EC - EC_{WB} = EC \left(1 - \frac{EC_{WB}}{EC} \right) = EC \left(1 - \frac{b_{abs,WB}^{BC}}{b_{abs}^{BC}} \right) = \\ &= EC \left(1 - \frac{b_{abs,WB}^{BC}}{b_{abs} - b_{abs}^{BrC}} \right) = EC \left(\frac{b_{abs} - b_{abs}^{BrC} - b_{abs,WB}^{BC}}{b_{abs} - b_{abs}^{BrC}} \right) \end{aligned} \quad (10)$$

305

306 Focusing on the IR range where BrC contribution to total b_{abs} is minimized (thus reducing
 307 uncertainties on the denominator evaluation) we obtain:

308

$$309 \quad EC_{FF} = EC \frac{b_{abs,FF}^{BC}(850\text{ nm})}{b_{abs}(850\text{ nm}) - b_{abs}^{BrC}(850\text{ nm})} \quad (11)$$

310

$$EC_{WB} = EC \frac{b_{abs,WB}^{BC}(850 \text{ nm})}{b_{abs}(850 \text{ nm}) - b_{abs}^{BrC}(850 \text{ nm})} \quad (12)$$

312

313 Figure 3 shows the EC apportionment deduced by eq.(11) and (12) at Propata and Genoa.
 314 EC_{WB} is higher in winter while it becomes small or almost zero during summertime. The
 315 resulting $MAC^{BC}(850nm)$ is $(6.57 \pm 0.13) \text{ m}^2 \text{ g}^{-1}$ ($R^2 = 0.87$) and $(6.40 \pm 0.10) \text{ m}^2 \text{ g}^{-1}$ ($R^2 =$
 316 0.98) for Propata and Genoa, respectively. As the two values are comparable within the
 317 experimental uncertainties, the mean value is calculated $\langle MAC^{BC}(850nm) \rangle = 6.5 \pm 0.10 \text{ m}^2$
 318 g^{-1} . However, it is worthy to note that generally the MAC^{BC} is an apparent, site-specific value
 319 including ambient factors as reported in several works (Bond & Bergstrom, 2006).

320

321 *Mass apportionment: organic carbon*

322 The OC source apportionment is less straightforward as not light-absorbing carbon as well as
 323 non-combustion components (OC_{NC}) such as spores, pollen, etc. can contribute to OC. The
 324 OC values determined by the TOT analysis can be thus expressed as:

$$OC = OC_{FF} + OC_{WB} + OC_{NC} \quad (13)$$

326 In the following, all the biogenic compounds are considered as not optically active and they
 327 are summed up in the OC_{NC} term. Moreover, BrC is assumed to be produced only by the WB
 328 source; this is actually confirmed by the inter-comparison discussed in §4 (see also Zheng et
 329 al., 2013).

330 To perform the OC apportionment, it is assumed a linear relationship between BC_{FF} and
 331 OC_{FF} as well as between BrC and OC_{WB} (i.e. BC_{FF} and BrC are used as tracers for FF and
 332 WB sources, respectively). Moreover, the linear relationship between BC_{FF} and BrC and their
 333 absorption coefficients allows re-writing eq.(13) as:

334

$$OC = \underbrace{k_1 \cdot b_{abs,FF}^{BC}(850nm)}_{OC_{FF}} + \underbrace{k_2 \cdot b_{abs}^{BrC}(407nm)}_{OC_{WB}} + OC_{NC} \quad (14)$$

336 Where:

- 337 • OC is the organic carbon concentration in [$\mu\text{g m}^{-3}$] measured by TOT analysis;
- 338 • $b_{abs,FF}^{BC}(850nm)$ in [Mm^{-1}] is the contribution to the BC_{FF} absorption @ $\lambda = 850nm$;
- 339 • $b_{abs}^{BrC}(407nm)$ in [Mm^{-1}] is the contribution to the BrC absorption @ $\lambda = 407nm$;
- 340 • k_1 is a constant coefficient in [g m^{-2}] related to the $MAC^{BC}(850nm)$ and to OC_{FF}/BC_{FF} ;
- 341 • k_2 is a constant coefficient in [g m^{-2}] related to the BrC MAC @ $\lambda = 407nm$ and to
- 342 OC_{WB}/BrC ;

343 • OC_{NC} is expressed in [$\mu\text{g m}^{-3}$].

344 The absorption coefficient determined at $\lambda = 850\text{nm}$ and 407nm were chosen as starting
345 points for OC_{FF} and OC_{WB} determination. The choice of $\lambda = 407\text{nm}$ as reference for the
346 evaluation of BrC contribution is related to the unavailability of the laser diode with $\lambda =$
347 375nm during the first winter campaign.

348 With our approach, both k_1 and k_2 are directly determined by the experimental optical data.
349 In samples where the α_{exp} is close to 1.0 (i.e. the measured $b_{\text{abs}}(\lambda)$ approximately follows λ^{-1})
350 $b_{\text{abs}}^{\text{BrC}}$ is negligible, thus eq.(14) reduces to:

351

$$352 \quad OC \approx k_1 \cdot b_{\text{abs},FF}^{\text{BrC}}(850\text{nm}) + OC_{NC} \quad \text{when} \quad \alpha_{\text{exp}} \approx 1 \quad (15)$$

353

354 The k_1 parameter is then determined by a regression study of OC vs. $b_{\text{abs},FF}^{\text{BrC}}$ restricted to the
355 samples with $\alpha_{\text{exp}} \approx 1$.

356 Such analysis on the rural site dataset is shown in Figure 4a and gives $k_1 = 0.52 \pm 0.06$ ($R^2 =$
357 0.79). Once determined k_1 - and thus the OC_{FF} contribution for each sample - k_2 can be
358 calculated by performing another linear regression involving the remaining part of the
359 dataset:

$$360 \quad OC - OC_{FF} = k_2 \cdot b_{\text{abs}}^{\text{BrC}}(407 \text{ nm}) + OC_{NC} \quad (16)$$

361

362 Whereas k_1 remains nearly constant during the whole campaign (data in Figure 4a span all
363 over the year), the value of k_2 is season-dependent: in the cold period of 2013 (approximately
364 February-March 2013), the regression study (Figure 4b, open squares) gives $k_2 = 0.36 \pm 0.03$
365 ($R^2 = 0.91$) while in the warm period (approximately between May and July 2013, Figure 4c)
366 k_2 is 1.53 ± 0.23 ($R^2 = 0.88$). In the last part of the campaign (between November 2013 and
367 February 2014) k_2 is again very close to the value found in the cold period of 2013 (Figure
368 4b, full triangles) with a value of 0.43 ± 0.02 ($R^2 = 0.97$). In this analysis, transition days
369 between cold and warm periods are not taken into account. Once determined k_1 and k_2 , OC_{FF}
370 and OC_{WB} are calculated for each sample and OC_{NC} is obtained by eq.(13). In Propata, OC_{WB}
371 concentration values are typically high during wintertime and especially during late fall 2013
372 (Figure 5a). The OC_{FF} fraction is similar all over the year with a percentage increase in early
373 springtime. OC_{NC} concentration values are mostly negligible during the cold periods while
374 they increase to a mean $\langle OC_{NC} \rangle = (0.44 \pm 0.10) \mu\text{g m}^{-3}$ during late spring and summer. In
375 Genoa, the OC_{FF} fraction is dominant during the whole campaign accounting for about 70%

376 of total OC (Figure 5b). In Table 1 a summary of the apportionment results for the two sites
377 is reported.

378

379 **4. Comparison with independent techniques**

380 The reliability of the optical mass apportionment was checked by independent levoglucosan
381 and ^{14}C measurements. The comparison between parallel determinations (please note that
382 radiocarbon measurements have been performed on 4 samples only, see §2) is shown in
383 Figure 6a.

384 A very good agreement between f_{NF} by radiocarbon measurements and levoglucosan/TC is
385 found (slope = 4.38, $R^2 = 0.99$). The intercept value (55%) can be probably attributed to local
386 background secondary organic aerosol, as the possible biogenic contribution seems to be
387 negligible according to the optical apportionment (Figure 5a).

388 The best agreement between the results obtained by our optical approach and radiocarbon
389 measurements is found for $\alpha_{\text{WB}} = 1.8$ (Figure 6b); this value is similar to $\alpha_{\text{WB}} = 1.86$ found by
390 Sandradewi et al. (2008).

391 Literature works (Favez et al., 2010; Sandradewi et al., 2008; Harrison et al., 2013) showed
392 the sensitivity of the Aethalometer model results to the a priori setting of α_{FF} and α_{WB} . In this
393 work, few trials changing the values assumed for α_{BC} , α_{FF} , and α_{WB} were carried out; among
394 them, the one most affecting the regression parameters in Figure 6b is α_{WB} . Indeed, varying
395 this parameter by ± 0.1 a change in the slope of about $\pm 10\%$ and in the intercept by $\pm 8\%$ is
396 observed.

397 Furthermore, for the low volume samples the reliability of the optical mass apportionment
398 was verified versus the independent determination of the levoglucosan concentration. In
399 Figure 7a the following results are reported: $\text{OC}_{\text{WB}} = (5.31 \pm 0.15) \cdot \text{levoglucosan}$ ($R^2 = 0.90$)
400 at the rural background site in Propata and $\text{OC}_{\text{WB}} = (7.10 \pm 0.41) \cdot \text{levoglucosan}$ ($R^2 = 0.89$) at
401 the urban background site in Genoa. Especially in Propata, the determined OC_{WB} to
402 levoglucosan concentration ratio is in very good agreement with results by Piazzalunga et al.
403 (2011), Bernardoni et al. (2011), and Favez et al., (2010) who used $\text{OM}_{\text{WB}}/\text{levoglucosan} \sim$
404 10.8 and $\text{OM}_{\text{WB}}/\text{OC}_{\text{WB}} = 1.8$. A good correlation with levoglucosan was found for EC_{WB} too;
405 the regression study (Figure 7b) gave $\text{EC}_{\text{WB}} = (1.10 \pm 0.05) \cdot \text{levoglucosan}$ ($R^2 = 0.85$) with no
406 significant differences between the two datasets. On the contrary, the levoglucosan
407 concentration values do not show any correlation with the EC_{FF} and OC_{FF} values.

408 We can therefore conclude that, at the present level of knowledge, $\alpha_{\text{WB}} = 1.8 \pm 0.1$ is a
409 reliable figure for future uses of the optical apportionment methodology.

410

411 **5. Conclusions**

412 The MWAA can measure offline the aerosol absorption coefficient at 5 wavelengths ranging
413 from IR to UV. In this way, a new apportionment methodology based on the so-called
414 Aethalometer model but fully exploiting the multi-wavelength approach, can be applied to
415 monitor the variability of WB and FF contributions to aerosol optical absorption and mass
416 concentration.

417 The model results depend on the choice of α_{BC} , α_{FF} , and α_{WB} values in eq.(4) and (5): in the
418 present study, they have been fixed according to the recent literature and validated against
419 independent techniques ($\alpha_{\text{BC}} = \alpha_{\text{FF}} = 1.0$; $\alpha_{\text{WB}} = 1.8$). Although these values in principle can
420 be site-dependent and modified in future works, the data reduction approach here proposed
421 remains valid. The optical apportionment is based on the assumption that the absorbing
422 species in aerosols are related to FF and WB only. In case of significant dust intrusions, this
423 might be not true and could lead to inaccurate source apportionment. The possible optical
424 activity of biogenic compounds is also neglected in this work but this issue merits a further
425 investigation. Despite of the mentioned limitations, in this paper we introduce a new
426 methodology which produces sounding results both at a rural site and in a large coastal city in
427 Italy.

428

429 **6. Acknowledgements**

430 This work was partially financed by the National Institute of Nuclear Physics (INFN) in the
431 frame of the MANIA experiment and by the Amministrazione Provinciale di Genova. The
432 authors acknowledge Vincenzo Ariola, Franco Parodi, and Francesco Saffioti (INFN-
433 Genova) for technical support in the MWAA development and. Maria Teresa Zannetti and
434 Federico Manni (Amministrazione Provinciale di Genova) for the collaboration during
435 sampling.

436

437

438 **References**

439 Ajtai, T., Filep, Á., Schnaiter, M., Linke, C., Vragel, M., Bozóki, Z., Szabó, G., Leisner, T.,
440 2010. A novel multi-wavelength photoacoustic spectrometer for the measurement of the

441 UV-vis-NIR spectral absorption coefficient of atmospheric aerosols. *Journal of Aerosol*
442 *Science*, 41, 1020-1029.

443 Andreae, M.O., 2001. The dark side of aerosols. *Nature*, 409, 671-672.

444 Andreae, M.O., & Gelencsér, A., 2006. Black Carbon or Brown Carbon? The nature of light-
445 absorbing carbonaceous aerosol. *Atmospheric Chemistry and Physics*, 6, 3131-3148.

446 Bernardoni, V., Vecchi, R., Valli, G., Piazzalunga, A., Fermo, P. (2011). PM10 source
447 apportionment in Milan (Italy) using time-resolved data. *Science of the Total Environment*,
448 409, 4788-4795.

449 Bernardoni, V., Calzolari, G., Chiari, M., Fedi, M., Lucarelli, F., Nava, S., Piazzalunga, A.,
450 Riccobono, F., Taccetti, F., Valli, G., Vecchi, R., 2013. Radiocarbon analysis on organic
451 and elemental carbon in aerosol samples and source apportionment at an urban site in
452 Northern Italy. *Journal of Aerosol Science*, 56, 88-99.

453 Bond, T.C., Anderson, T.L., Campbell, D., 1999. Calibration and intercomparison of filter-
454 based measurements of visible light absorption by aerosols. *Aerosol Science and*
455 *Technology* 30, 582-600.

456 Bond, T.C., & Bergstrom, R.W., 2006. Light absorption by carbonaceous particles: an
457 investigative review. *Aerosol Science and Technology*, 40, 27-67.

458 Bond, T.C., Doherty, S.J., Fahey, D.W., Forster, P.M., Berntsen, T., DeAngelo, B.J., Flanner,
459 M.G., Ghan, S., Kärcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P.K., Sarofim, M.C.,
460 Schultz, M.G., Schulz, M., Venkataraman, C., Zhang, H., Zhang, S., Bellouin, N.,
461 Guttikunda, S.K., Hopke, P.K., Jacobson, M.Z., Kaiser, J.W., Klimont, Z., Lohmann, U.,
462 Schwarz, J.P., Shindell, D., Storelvmo, T., Warren, S.G., Zender, C.S., 2013. Bounding the
463 role of black carbon in the climate system: A scientific assessment. *Journal of Geophysical*
464 *Research: Atmospheres*, 118, 5380-5552.

465 Brun, R., & Rademakers, F., 1997. ROOT-An Object Oriented Data Analysis Framework.
466 *Nuclear Instruments and Methods in Physics Research Section A: Accelerators,*
467 *Spectrometers, Detectors and Associated Equipment*, 389, 81-86.

468 Calzolari G., Bernardoni V., Chiari M., Fedi M., Lucarelli F., Nava S., Riccobono F., Taccetti
469 F., Valli G., Vecchi R., 2011. The new sample preparation line for radiocarbon
470 measurements on atmospheric aerosol at LABEC. *Nuclear Instruments and Methods in*
471 *Physics Research Section B: Beam Interactions with Materials and Atoms*, 269, 203-208.

472 Cavalli, F., Putaud, J.P., Viana, M., Yttri, K.E., Gemberg, J., 2010. Toward a Standardized
473 Thermal-Optical Protocol for Measuring Atmospheric Organic and Elemental Carbon: The
474 EUSAAR Protocol. *Atmospheric Measurement Techniques*, 3, 79-89.

475 Collaud Coen, M., Weingartner, E., Apituley, A., Ceburnis, D., Fierz-Schmidhauser, R.,
476 Flentje, H., Henzing, J.S., Jennings, S.G., Moerman, M., Petzold, A., Schmid, O.,
477 Baltensperger, U., 2010. Minimizing light absorption measurement artifacts of the
478 Aethalometer: evaluation of five correction algorithms. *Atmospheric Measurement and*
479 *Techniques*, 3, 457-474.

480 Favez, O., El Haddad, I., Piot, C., Boreave, A., Abidi, E., Marchand, N., Jaffrezo, J.L.,
481 Besombes, J.L., Personnaz, M.B., Sciare, J., Wortham, H., Geroge, C., D'Anna, B., 2010.
482 Inter-comparison of source apportionment models for the estimation of wood burning
483 aerosols during wintertime in an Alpine city (Grenoble, France). *Atmospheric Chemistry*
484 *and Physics*, 10, 5295-5314.

485 Fedi M.E., Bernardoni V., Caforio L., Calzolari G., Carraresi L., Manetti M., Taccetti F.,
486 Mandò P.A., 2013. Status of sample combustion and graphitization lines at INFN-LABEC,
487 Florence, *Radiocarbon*, 55, 657-664.

488 Filep, Á., Ajtai, T., Utry, N., Pintér, M.D., Nyilas, T., Takács, S., Máté, Z., Gelencsér, A.,
489 Hoffer, A., Schnaiter, M., Bozóki, Z., Szabó, G., 2013. Absorption Spectrum of Ambient
490 Aerosol and Its Correlation with Size Distribution in Specific Atmospheric Conditions
491 after a Red Mud Accident. *Aerosol and Air Quality Research*, 13, 49-59.

492 Flowers, B.A., Dubey, M.K., Mazzoleni, C., Stone, E.A., Schauer, J.J., Kim, S.-W., Yoon,
493 S.C., 2010. Optical-chemical-microphysical relationships and closure studies for mixed
494 carbonaceous aerosols observed at Jeju Island; 3-laser photoacoustic spectrometer, particle
495 sizing, and filter analysis. *Atmospheric Chemistry and Physics*, 10, 10387-10398.

496 Hänel, G., 1987. Radiation budget of the boundary layer: Part II. Simultaneous measurement
497 of mean solar volume absorption and extinction coefficients of particles. *Beiträge zur*
498 *Physik der Atmosphäre*, 60, 241-247.

499 Hänel, G., 1994. Optical properties of atmospheric particles: Complete parameter sets
500 obtained through polar photometry and an improved inversion technique. *Applied Optics*,
501 33, 7187-7199.

502 Harrison, R.M., Beddows, D.C.S., Jones A.M., Calvo A., Alves C., Pio C., 2013. An
503 evaluation of some issues regarding the use of aethalometers to measure woodsmoke
504 concentrations. *Atmospheric Environment*, 80, 540-548.

505 James, F., 1978. MINUIT, A package of programs to minimise a function of n variables,
506 compute the covariance matrix, and find the true errors. Program library code D507,
507 CERN, 1978.

508 Kirchstetter, T.W., Novakok, T., Hobbs, P.V., 2004. Evidence that the spectral dependence of
509 light absorption by aerosols is affected by organic carbon. *Journal of Geophysical*
510 *Research*, 109, D21208.

511 Kirchstetter, T.W. & Thatcher, T.L., 2012. Contribution of organic carbon to wood smoke
512 particulate matter absorption of solar radiation. *Atmospheric Chemistry and Physics*, 12,
513 6067-6072.

514 Lack, D.A. & Langridge, J.M., 2013. On the attribution of black and brown carbon light
515 absorption using the Ångström exponent. *Atmospheric Chemistry and Physics*, 13, 10535-
516 10543.

517 Lewis, K., Arnott, W.P., Moosmüller, H., Wold, C.E., 2008. Strong spectral variation of
518 biomass smoke light absorption and single scattering albedo observed with a novel dual-
519 wavelength photoacoustic instrument. *Journal of Geophysical Research*, 113, D16203.

520 Linke, C., Möhler, O., Veres, A., Mohácsi, Á., Bozóki, Z., Szabó, G., Schnaiter M., 2006.
521 Optical Properties and mineralogical composition of different Saharan mineral dust
522 samples: a laboratory study. *Atmospheric Chemistry and Physics*, 6, 3315-3323.

523 Massabò, D., Bernardoni, V., Bove, M.C., Brunengo, A., Cuccia, E., Piazzalunga, A., Prati,
524 P., Valli, G., Vecchi, R., 2013. A multi-wavelength optical set-up for the characterization
525 of carbonaceous particulate matter. *Journal of Aerosol Science*, 60, 34-46.

526 Moosmüller, H., Chakrabarty, R.K., Arnott, W.P., 2009. Aerosol light absorption and its
527 measurement: A review. *Journal of Quantitative Spectroscopy and Radiative Transfer*, 110,
528 844-878.

529 Moosmüller, H., Chakrabarty, R.K., Ehlers, K.M., Arnott, W.P., 2011. Absorption Ångström
530 coefficient, brown carbon, and aerosols: basic concepts, bulk matter, and spherical
531 particles. *Atmospheric Chemistry and Physics*, 11, 1217-1225.

532 Müller, T., Henzing, J. S., de Leeuw, G., Wiedensohler, A., Alastuey, A., Angelov, H.,
533 Bizjak, M., Collaud Coen, M., Engström, J. E., Gruening, C., Hillamo, R., Hoffer, A.,
534 Imre, K., Ivanow, P., Jennings, G., Sun, J. Y., Kalivitis, N., Karlsson, H., Komppula, M.,
535 Laj, P., Li, S.-M., Lunder, C., Marinoni, A., Martins dos Santos, S., Moerman, M., Nowak,
536 A., Ogren, J. A., Petzold, A., Pichon, J. M., Rodriguez, S., Sharma, S., Sheridan, P. J.,
537 Teinilä, K., Tuch, T., Viana, M., Virkkula, A, Weingartner, E., Wilhelm, R, Wang, Y.Q.,
538 2011. Characterization and intercomparison of aerosol absorption photometers: result of
539 two intercomparison workshops. *Atmospheric Measurement Techniques*, 4, 245–268.

540 Petzold, A., & Schöllner, M., 2004. Multi-angle absorption photometry - a new method for
541 the measurement of aerosol light absorption and atmospheric black carbon. *Journal of*
542 *Aerosol Science*, 35, 421-441.

543 Petzold, A., Schloesser, H., Sheridan, P.J., Arnott, W.P., Ogren, J.A., Virkkula, A., 2005.
544 Evaluation of Multiangle Absorption Photometry for Measuring Aerosol Light Absorption.
545 *Aerosol Science and Technology*, 39, Issue 1.

546 Piazzalunga, A., Fermo, P., Bernardoni, V., Vecchi, R., Valli, G., De Gregorio M.A., 2010. A
547 simplified method for levoglucosan quantification in wintertime atmospheric particulate
548 matter by high performance anion-exchange chromatography coupled with pulsed
549 amperometric detection. *International Journal of Environmental Analytical Chemistry*, 90,
550 934-947.

551 Piazzalunga, A., Belis, C., Bernardoni, V., Cazzuli, O., Fermo, P., Valli, G., Vecchi, R.,
552 2011. Estimates of wood burning contribution to PM by the macro-tracer method using
553 tailored emission factors. *Atmospheric Environment*, 45, 6642-6649.

554 Pöschl U., 2003. Aerosol particle analysis: challenges and progress. *Analytical and*
555 *Bioanalytical Chemistry*, 375, 3032

556 Sandradewi, J., Prevot, A.H., Szidat, S., Perron, N., Rami Alfarra, M., Lanz, V., Weingartner,
557 E., Baltensperger, U., 2008. Using Aerosol Light Absorption Measurements for the
558 Quantitative Determination of Wood Burning and Traffic emission Contributions to
559 Particulate Matter. *Environmental Science & Technology*, 42, 3316-3323.

560 Simoneit, B.R.T., Schauer, J.J., Nolte, C.G., Oros, D.R., Elias, V.O., Fraser, M.P., Rogge,
561 W.F., Cass, G.R., 1999. Levoglucosan, a tracer for cellulose in biomass burning and
562 atmospheric particles. *Atmospheric Environment*, 33, 173-182.

563 Utry, N., Ajtai, T., Filep, Á., Dániel P.,M., Hoffer, A., Bozoki, Z., Szabó, G., 2013. Mass
564 specific optical absorption coefficient of HULIS aerosol measured by a four-wavelength
565 photoacoustic spectrometer at NIR, VIS and UV wavelengths. *Atmospheric Environment*,
566 69, 321-324.

567 Utry, N., Ajtai, T., Filep, Á., Pintér, M., Török, Zs., Bozóki, Z., Szabó, G., 2014. Correlations
568 between absorption Angström exponent (AAE) of wintertime ambient urban aerosol and its
569 physical and chemical properties. *Atmospheric Environment*, 91, 52-59.

570 Vecchi, R., Bernardoni, V., Paganelli, C., Valli, G., 2014. A filter-based light-absorption
571 measurement with polar photometer: Effects of sampling artefacts from organic carbon.
572 *Journal of Aerosol Science*, 70, 15-25.

573 Yang, M., Howell, S.G., Zhuang, J., Huebert, B.J., 2009. Attribution of aerosol light
574 absorption to black carbon, brown carbon, and dust in China - interpretations of
575 atmospheric measurements during EAST-AIRE. *Atmospheric Chemistry and Physics*, 9,
576 2035-2050.

577 Zheng, G., He, K., Duan, F., Cheng, Y., Ma, Y., 2013. Measurement of humic-like
578 substances in aerosols: A review. *Environmental Pollution*, 181, 301-314.

579 Zotter, P., El-Haddad, I., Zhang, Y., Hayes, P.L., Zhang, X., Lin, Y-H., Wacker, L., Schnelle-
580 Kreis, J., Abbaszade, G., Zimmermann, R., Surratt, J.D., Weber, R., Jimenez, J.L., Szidat,
581 S., Baltensperger, U., Prévôt, A.S.H, 2014. Diurnal cycle of fossil and non fossil carbon
582 using radiocarbon analyses during CalNex. *Journal of Geophysical Research:*
583 *Atmospheres*, 119, 6818-6835.

584

585 **FIGURE CAPTIONS**

586

587 **Figure 1:** Mean BC_{FF} , BC_{WB} and BrC light absorption coefficients as a function of
588 wavelength at Propata during winter 2014.

589

590 **Figure 2:** Optical apportionment obtained at the rural site during wintertime 2014: (a) $\lambda =$
591 850 nm (IR) and (b) $\lambda = 375$ nm (UV).

592

593 **Figure 3:** EC apportionment at the rural site (a) and the urban background site (b).

594

595 **Figure 4:** Rural site dataset: (a) OC versus $b_{abs,FF}^{BC}(850\text{ nm})$ only for samples with $\alpha_{exp} \approx 1$;
596 (b) $OC - OC_{FF}$ vs. $b_{abs}^{BrC}(407\text{ nm})$ open squares refer to wintertime 2013 and full triangles to
597 wintertime 2014; (c) $OC - OC_{FF}$ vs. $b_{abs}^{BrC}(407\text{ nm})$ for samples collected in the warm
598 period.

599

600 **Figure 5:** OC apportionment for the rural site (a) and the urban background site (b).

601

602 **Figure 6:** Rural site high volume samples: (a) f_{NF} vs. levoglucosan/TC; (b) f_{NF} vs. $TC - TC_{FF}$
603 obtained by the optical approach ($TC_{FF} = EC_{FF} + OC_{FF}$).

604

605 **Figure 7:** (a) OC_{WB} vs. levoglucosan and (b) EC_{WB} vs. levoglucosan. In the case of OC_{WB}
606 two different OC_{WB} to levoglucosan ratios were found for Propata (open squares) and Genoa
607 (full triangles).

608

609 **Table 1:** Average EC and OC apportionment at the two sites. Values are given as percentages
610 of total measured EC and OC.

Table1[Click here to download Table: Table_1.docx](#)

Site	$\langle EC_{FF} \rangle$	$\langle EC_{WB} \rangle$	$\langle OC_{FF} \rangle$	$\langle OC_{WB} \rangle$	$\langle OC_{NC} \rangle$
Propata, winter 2013	68 ± 15	32 ± 11	63 ± 7	35 ± 10	7 ± 12
Propata, summer 2013	87 ± 19	13 ± 8	53 ± 4	11 ± 6	36 ± 12
Propata, winter 2014	47 ± 9	53 ± 9	38 ± 5	61 ± 5	4 ± 11
Genoa, fall 2013	84 ± 11	16 ± 7	67 ± 5	15 ± 5	19 ± 9

Table 1

Figure1

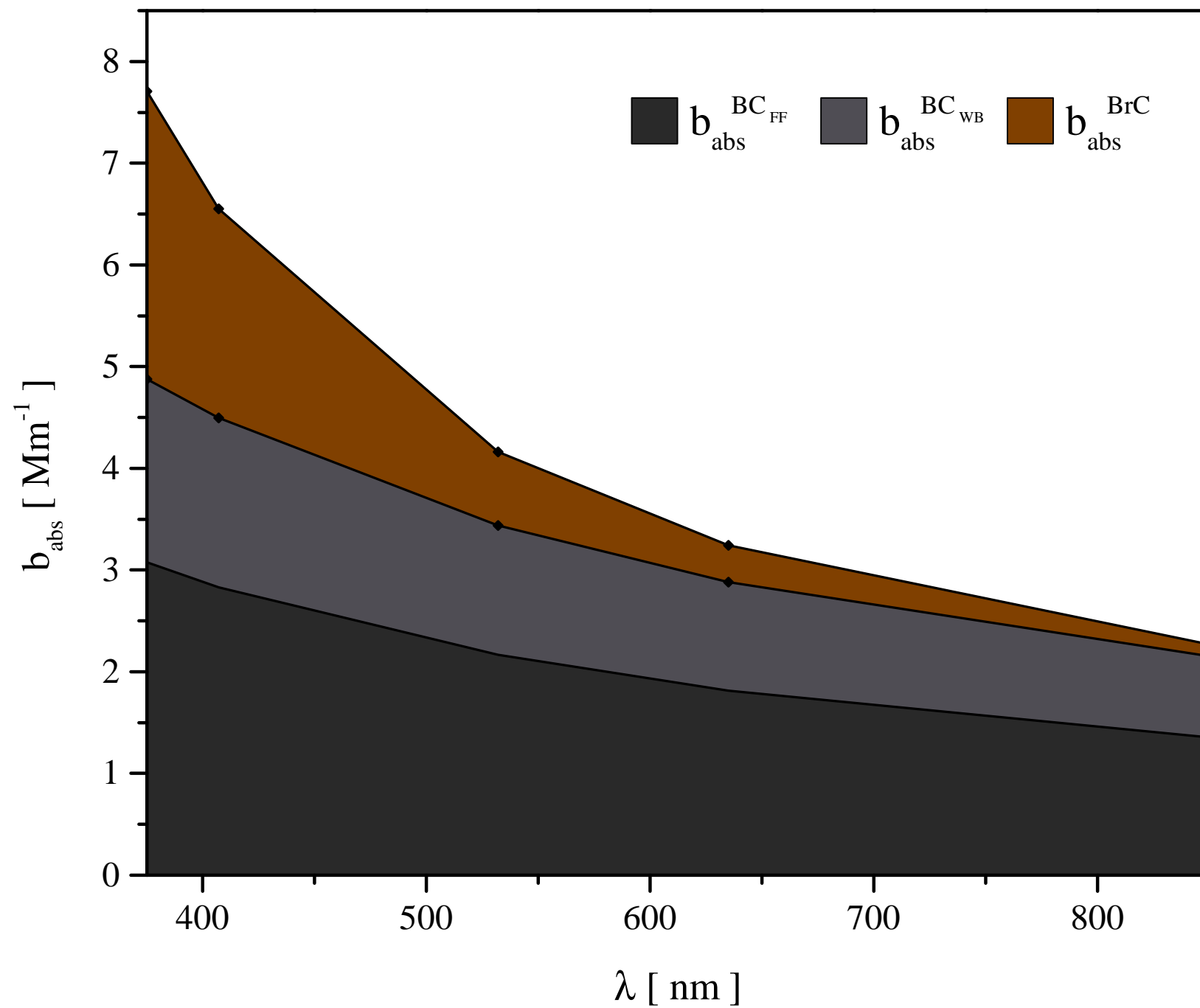


Figure2a

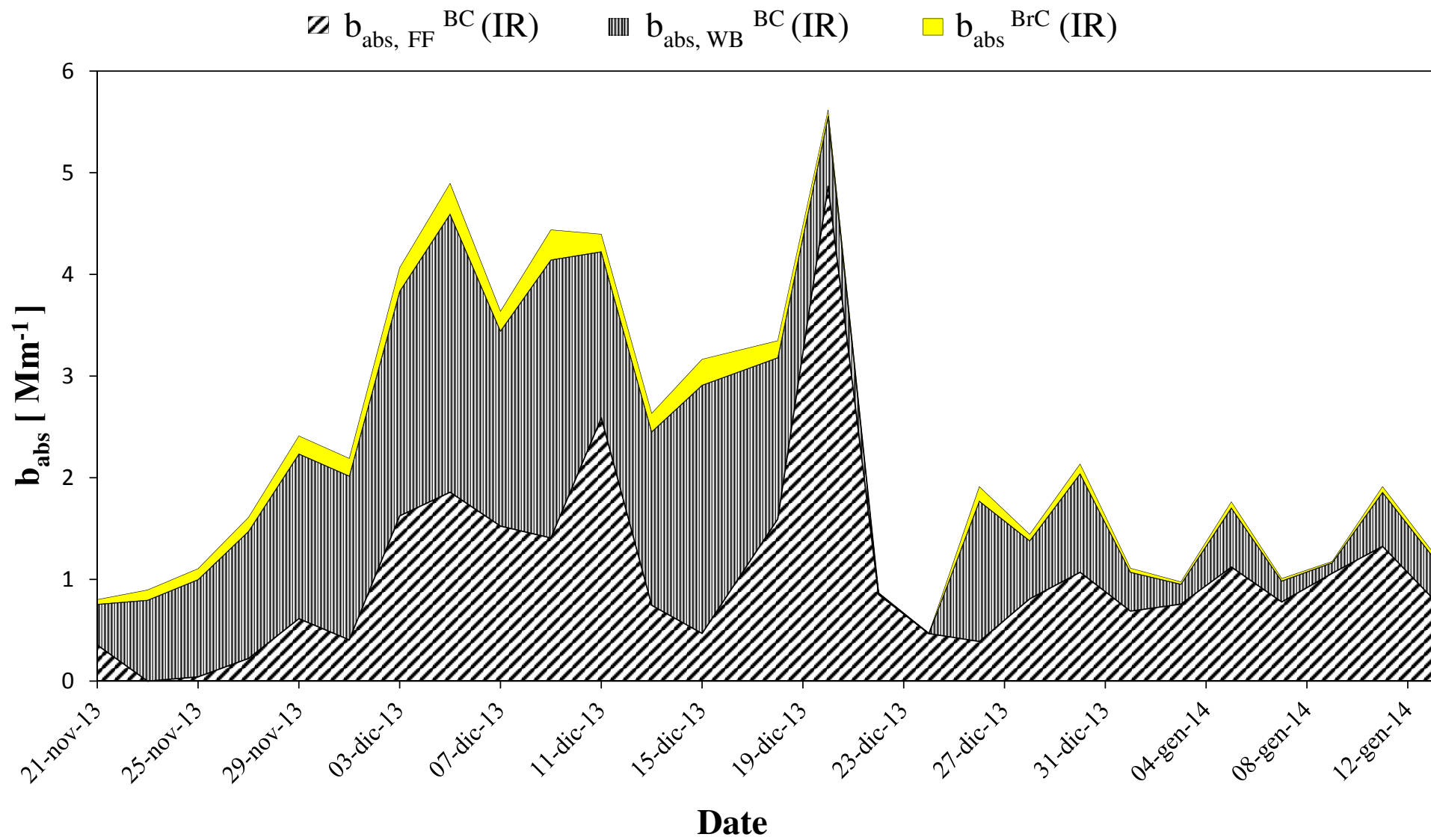


Figure2b

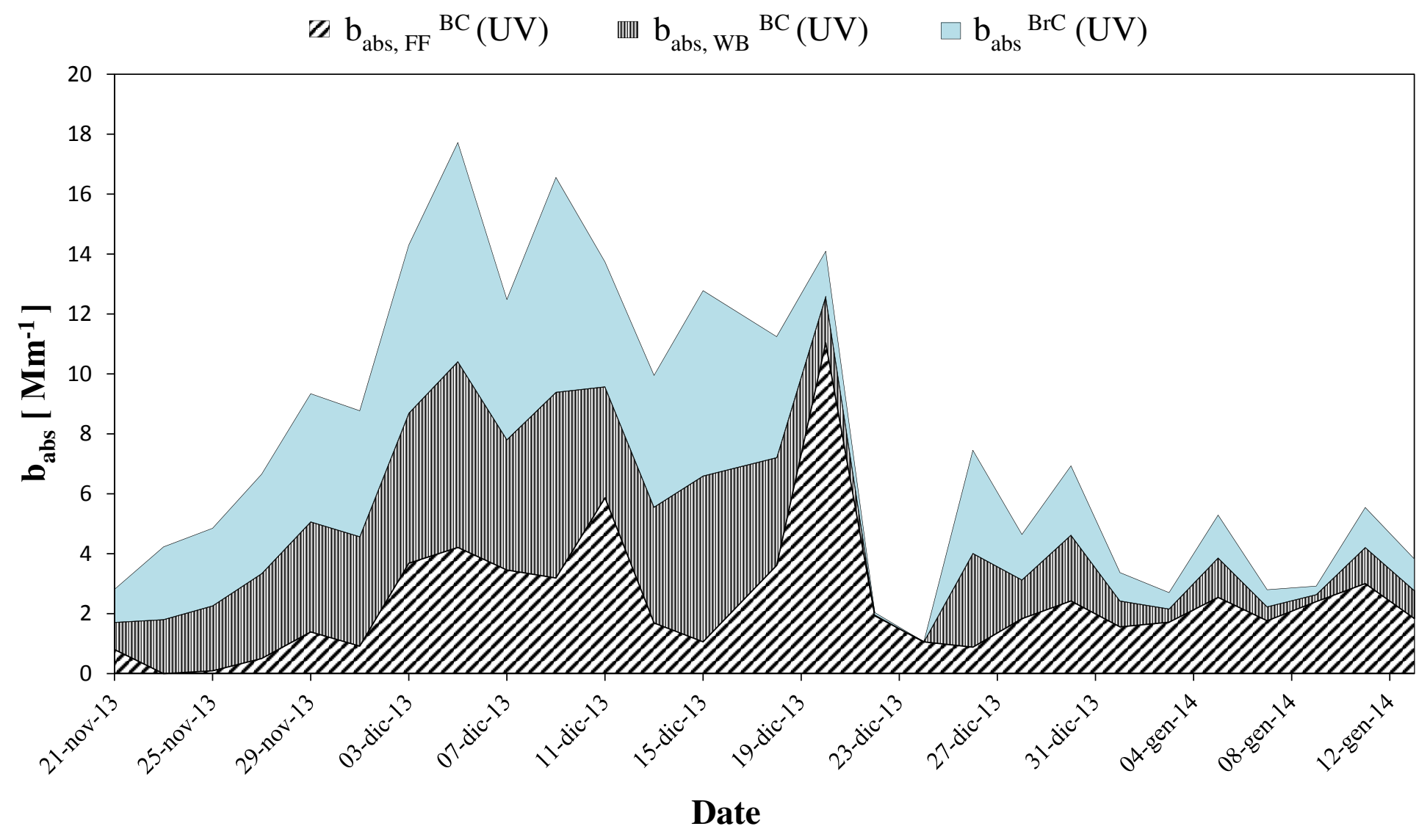


Figure3a

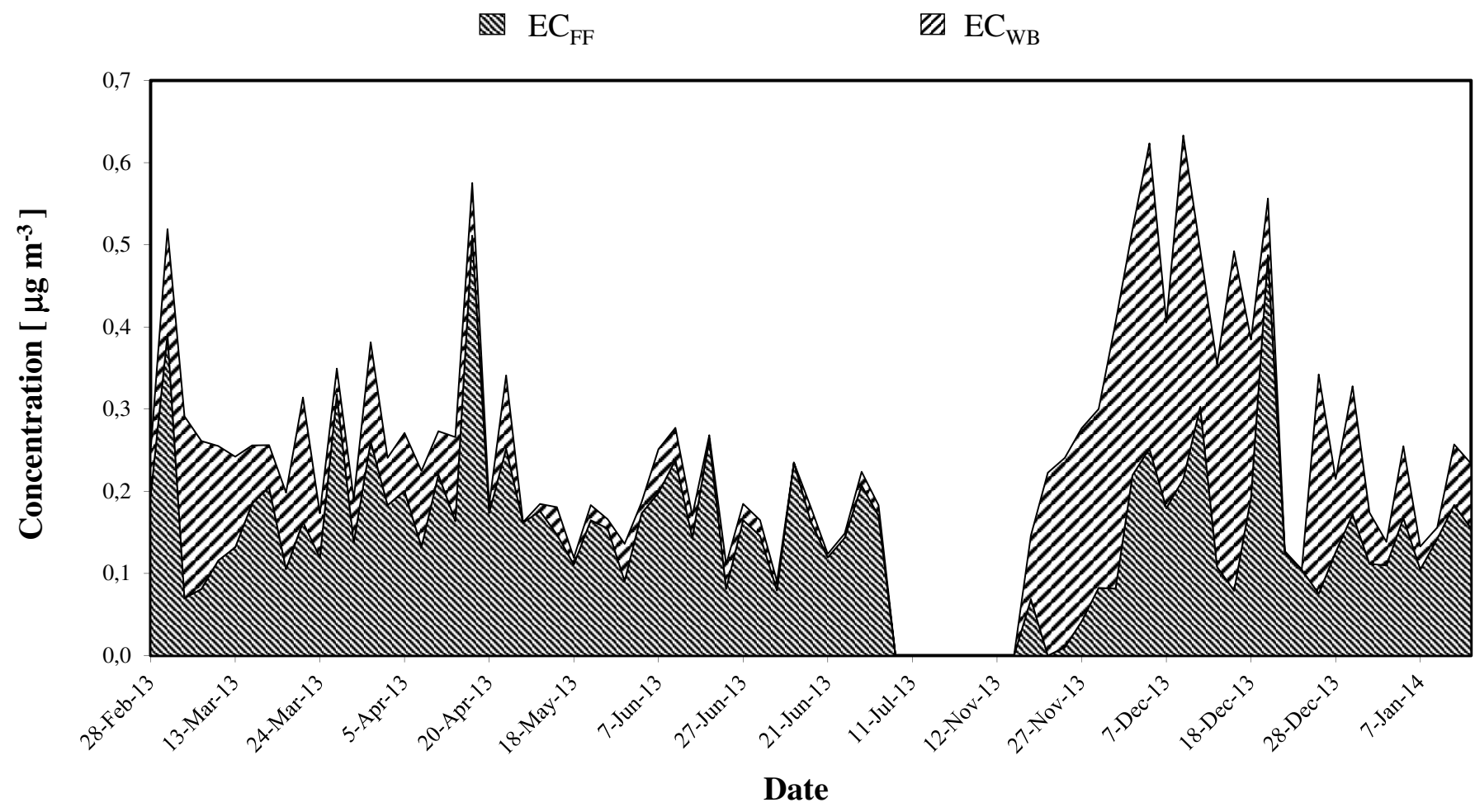


Figure3b

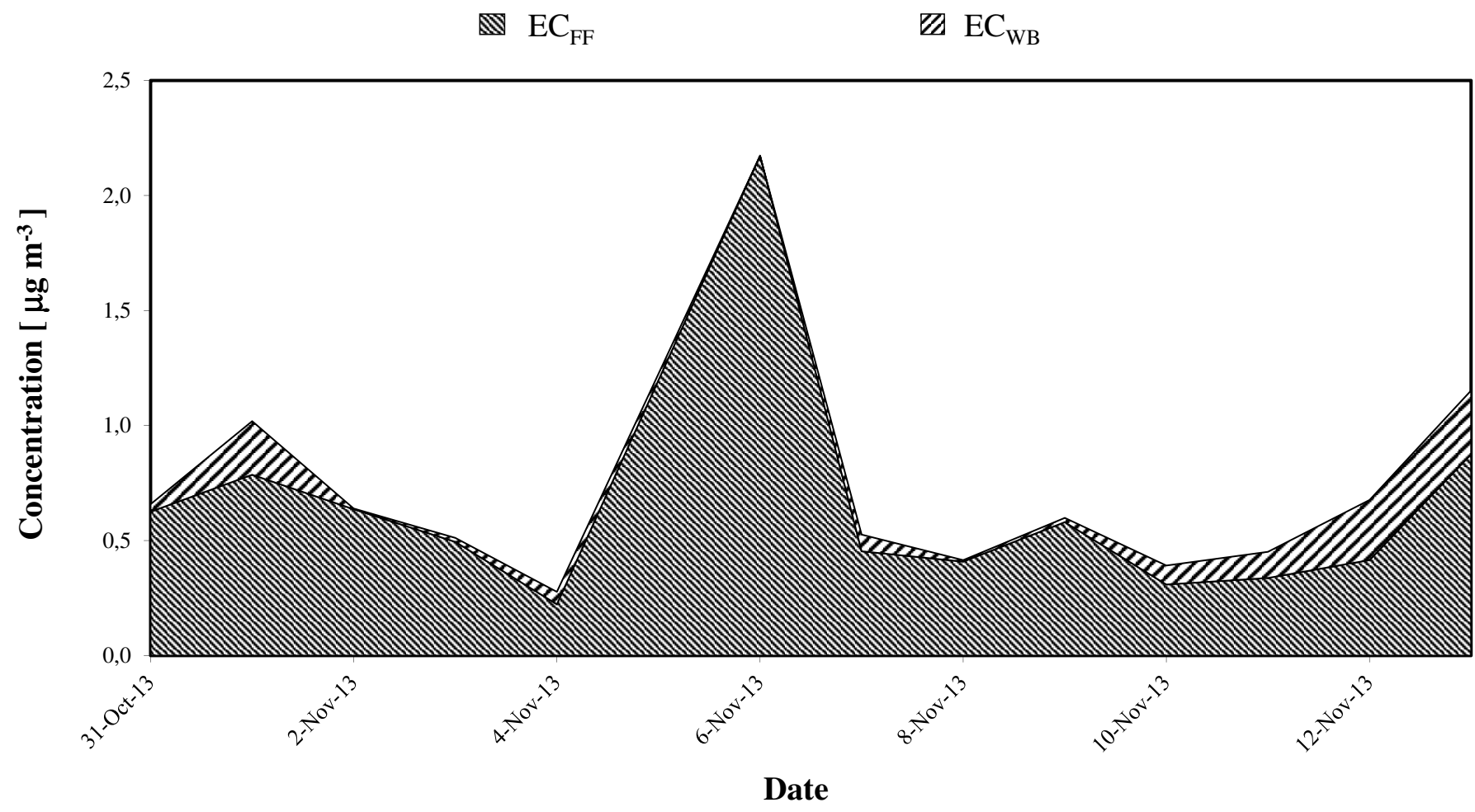


Figure4a

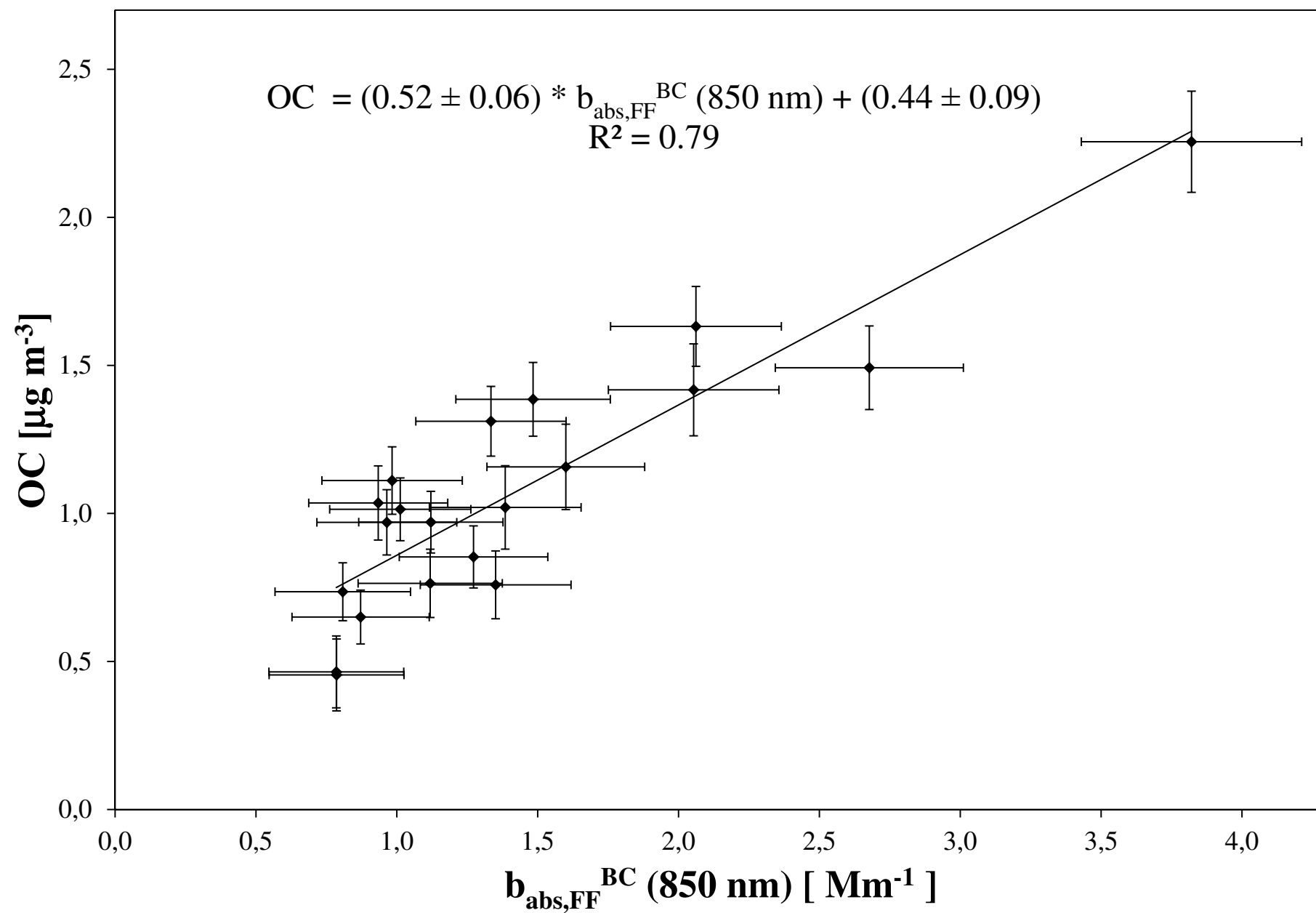


Figure4b

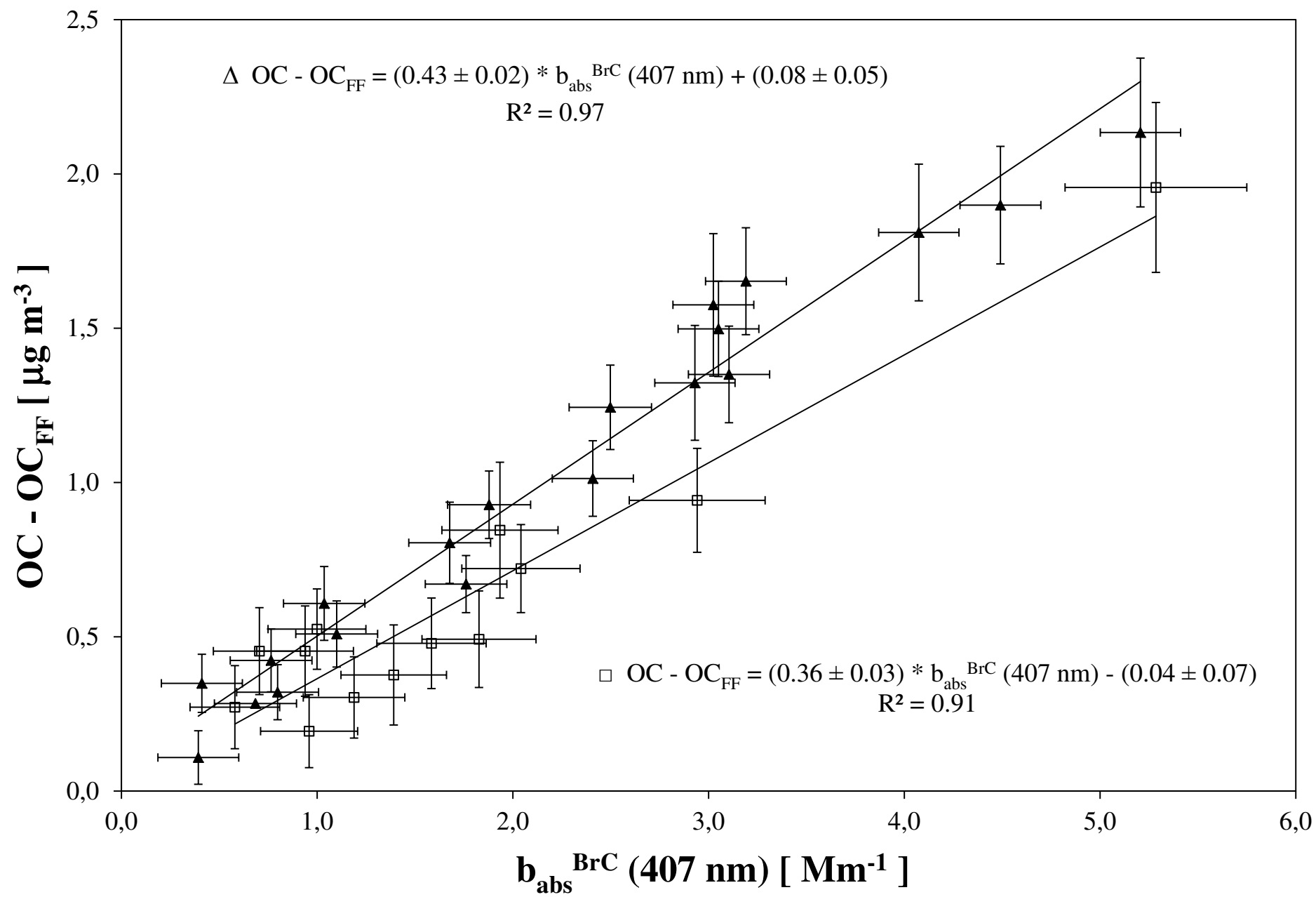


Figure4c

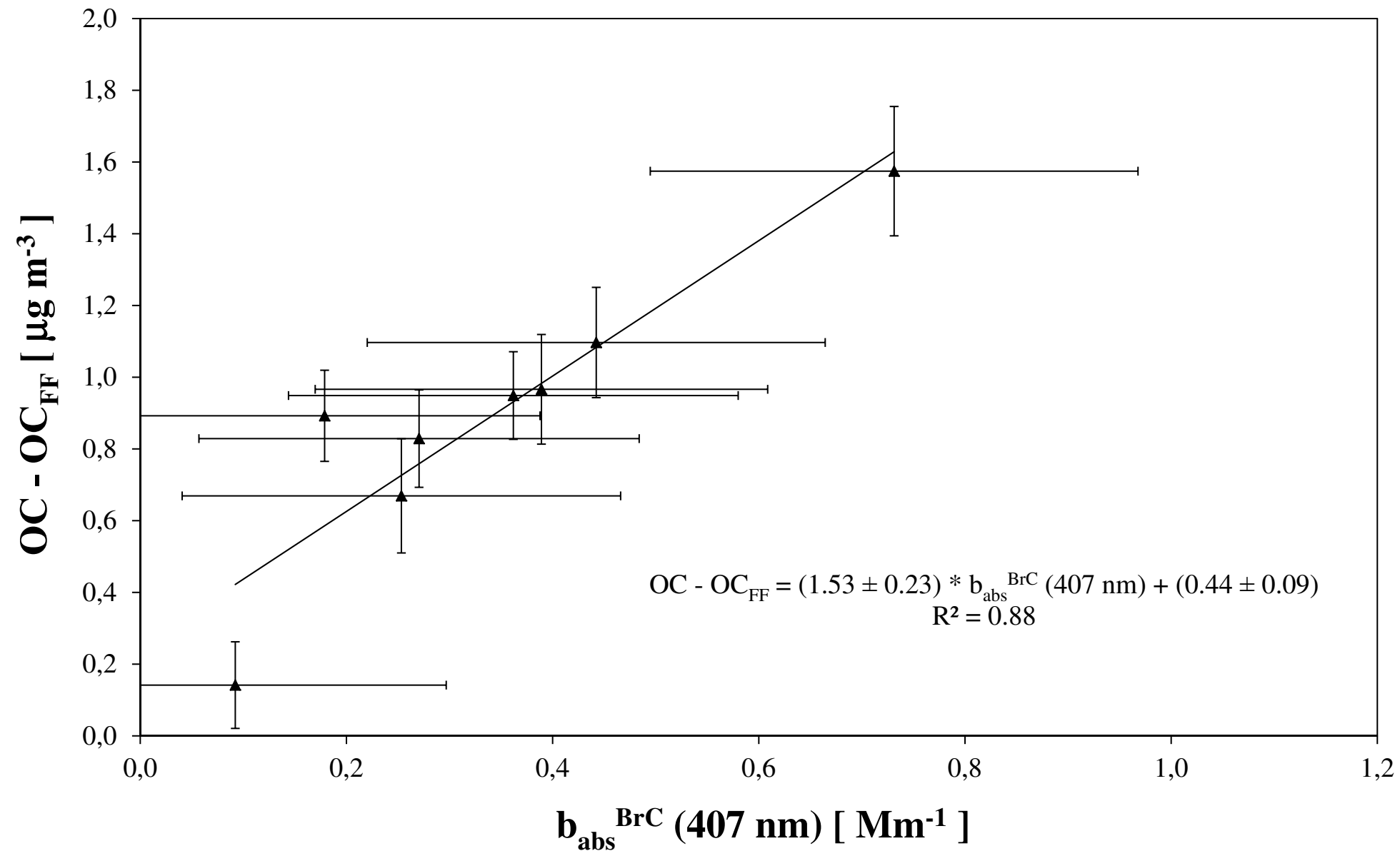


Figure5a

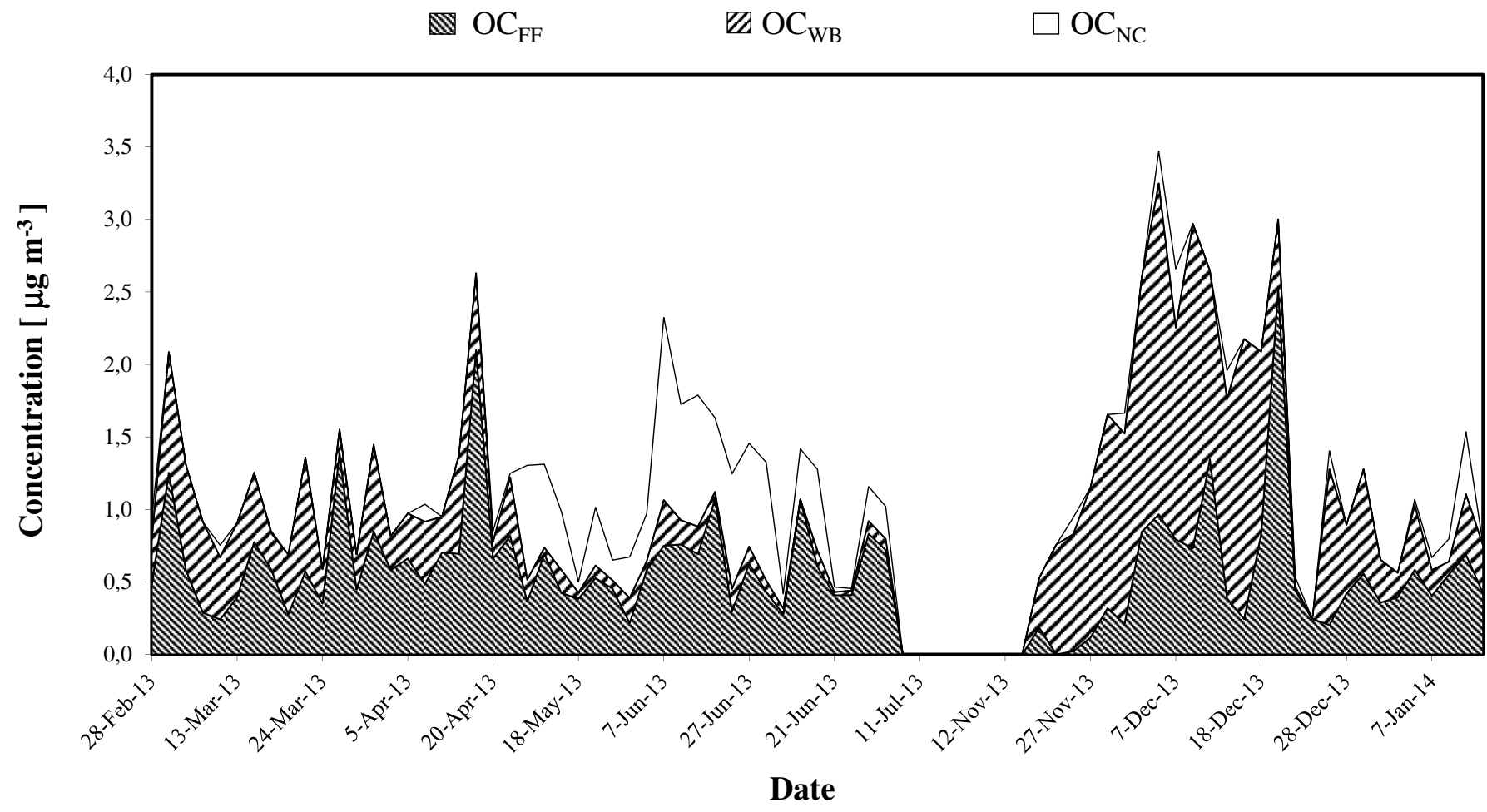


Figure5b

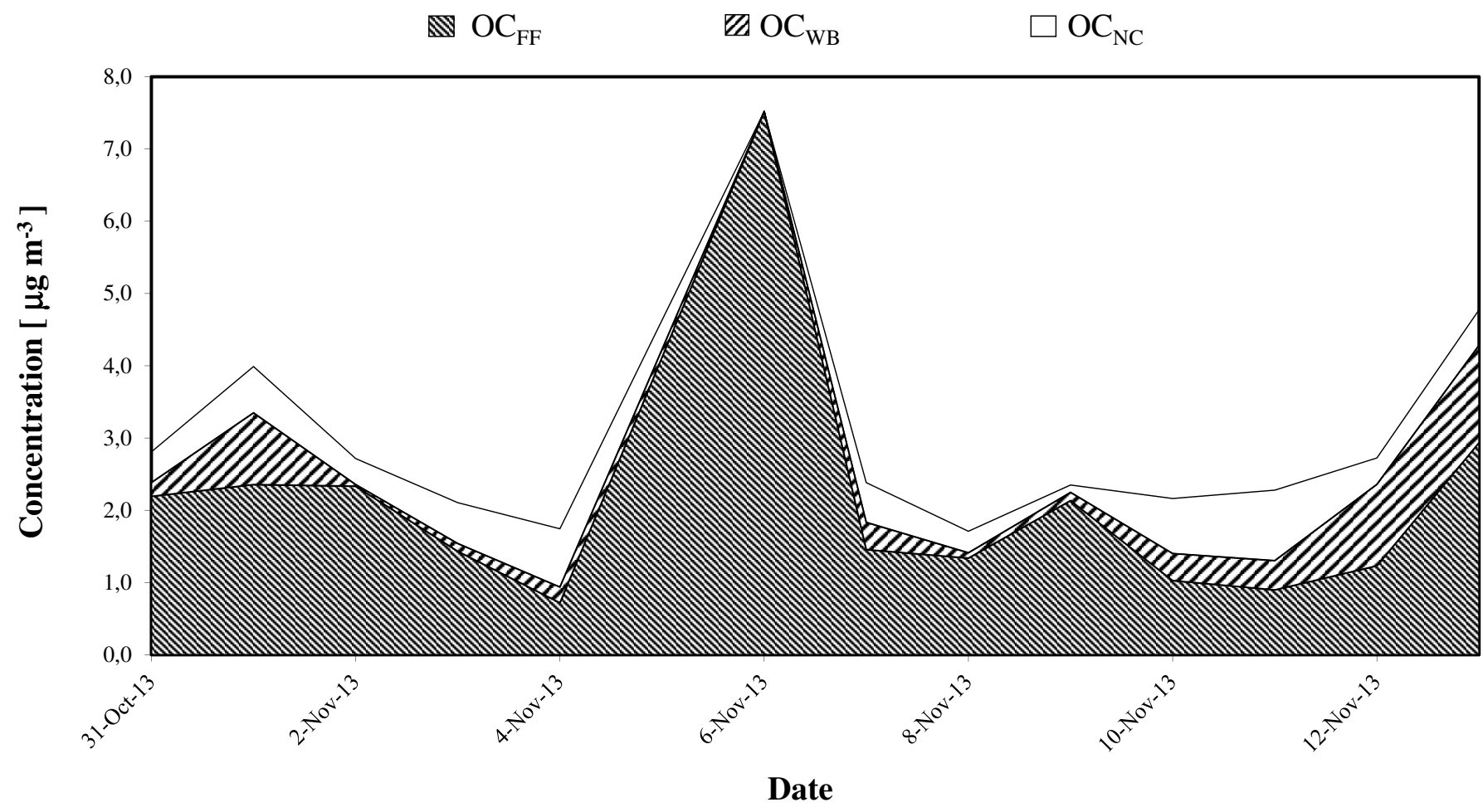


Figure6a

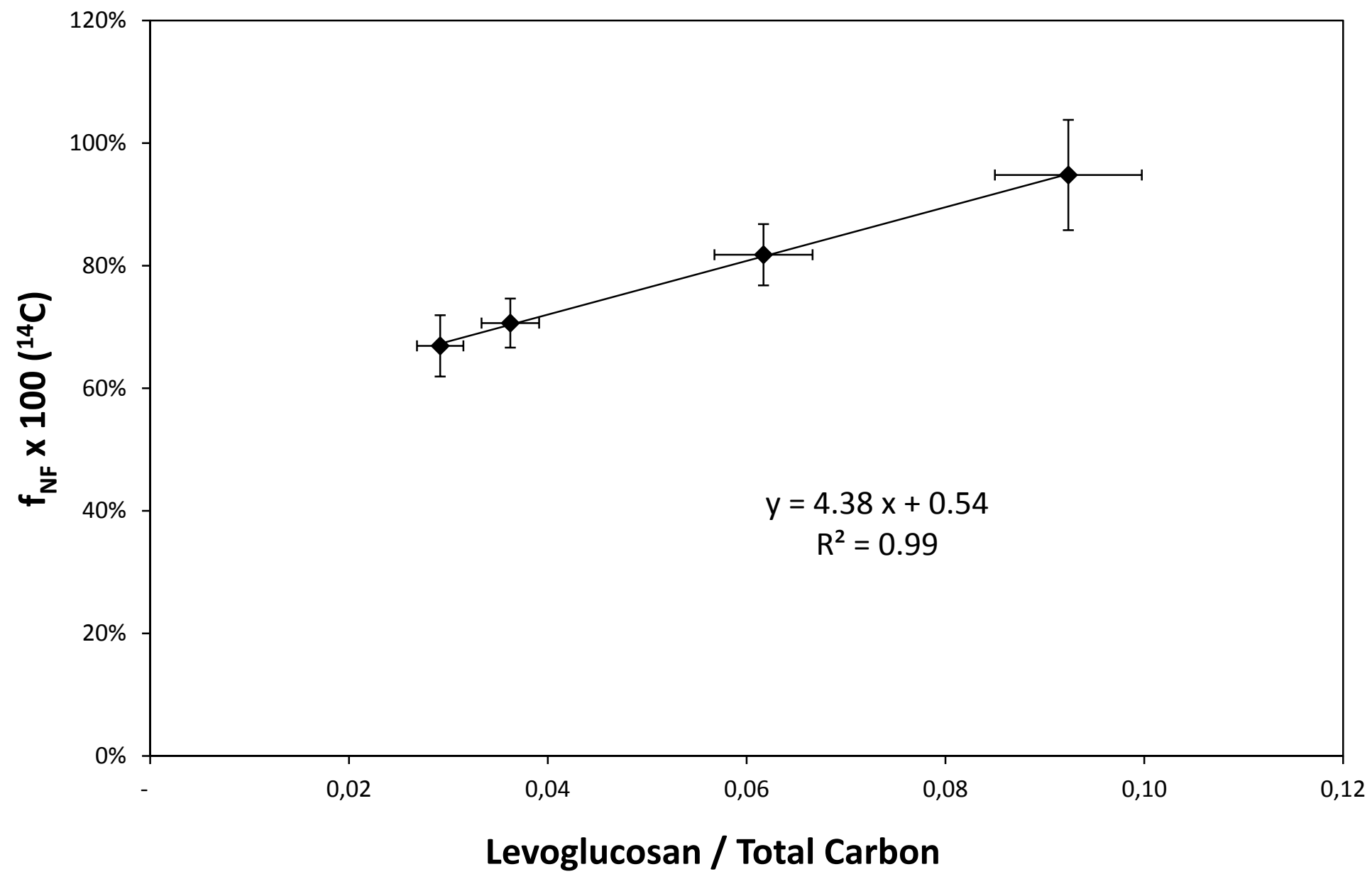


Figure6b

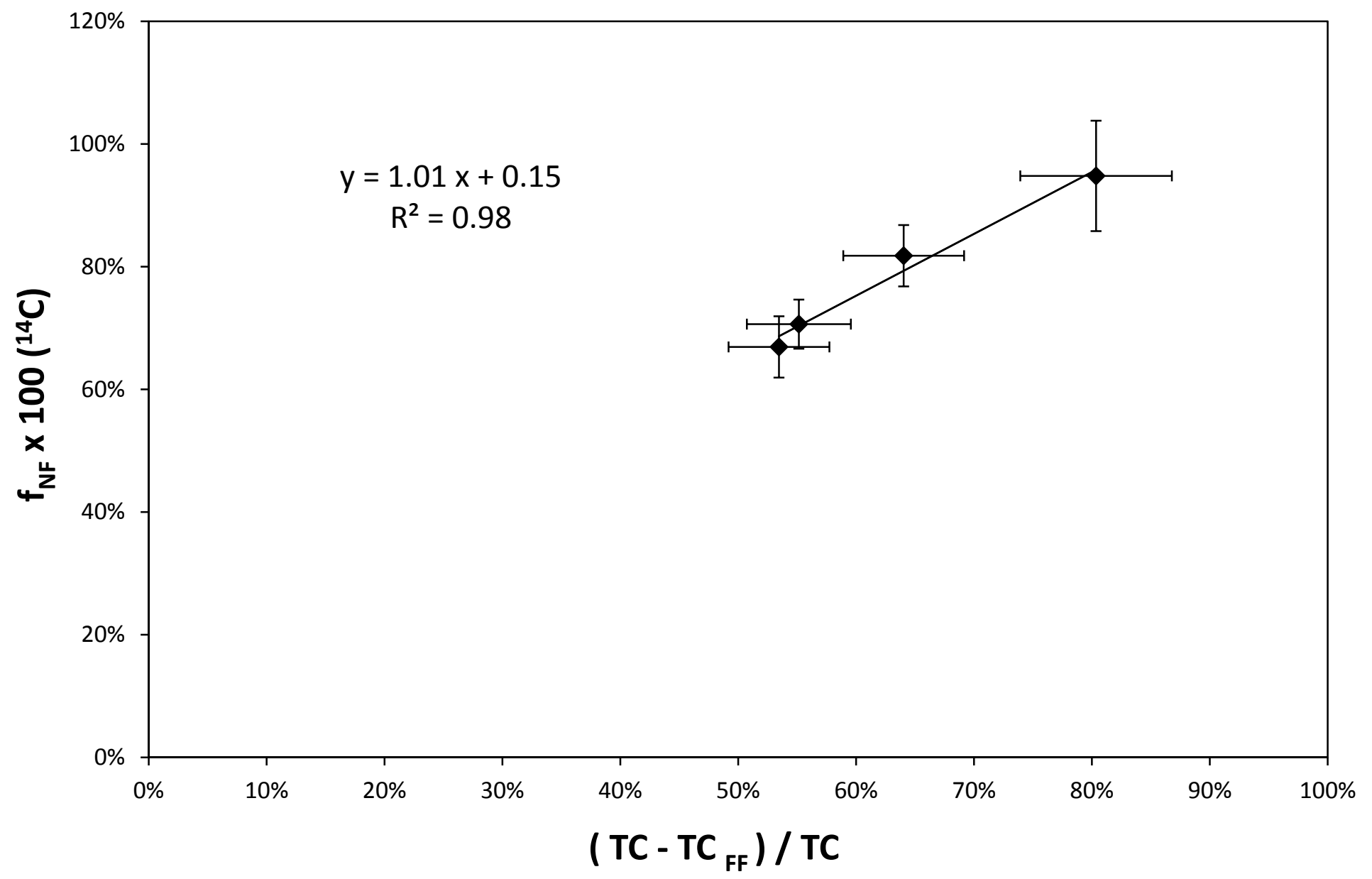


Figure7a

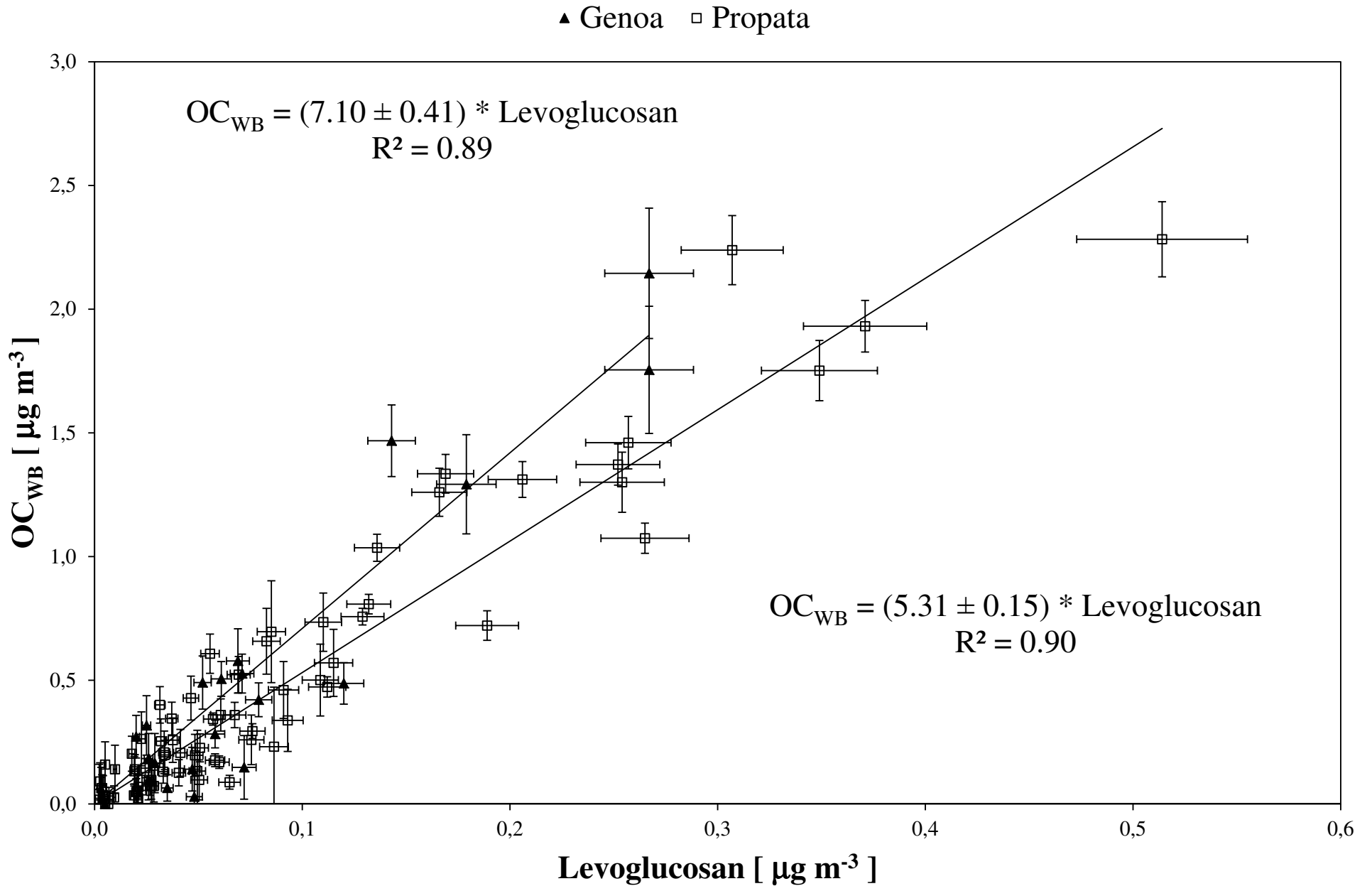


Figure7b

▲ Genoa □ Propata

