

1 Article

2 Investigation of trace and critical elements (including 3 actinides) in flotation sulphide concentrates of 4 **Kassandra mines (Chalkidiki, Greece)**

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19 **Abstract:** Pyrite/As-pyrite/arsenopyrite (Py-(As)Py-AsPy), galena (PbS) and sphalerite (ZnS)
20 concentrates from the flotation plants of Olympiada and Stratoni (Kassandra mines, Chalkidiki, N.
21 Greece) have been investigated for their trace and critical element content, including actinides
22 associated to natural radioactivity. It has been revealed that except Pb, Zn, Ag, and Au, being
23 exploited by Hellas Gold S.A., there are also significant concentrations of Sb and Ga (Sb: >0.2 wt.%
24 in PbS concentrate; Ga: 25 ppm in ZnS concentrate), but no considerable contents of Bi, Co, V, and
25 REE. Concerning other elements, As is found in elevated concentrations (> 1 wt.% in
26 Py-(As)Py-AsPy Olympiada concentrate and almost 1 wt.% in Stratoni PbS and ZnS concentrates)
27 together with Cd (specifically in ZnS concentrate). However, it has been postulated, for first time in
28 the literature, that actinides occur in very low concentrations (U<2 ppm and Th<0.5 ppm in all
29 examined concentrates), eliminating the possibility of natural radioactivity in the Hellas Gold S.A.
30 products. The concentrations of the natural radionuclides (²³⁸U, ²³²Th and ⁴⁰K) are much lower
31 compared to commercial granitic rocks. Thus, the associated radioactive dose is insignificant and it
32 can be assumed that no risk concerning natural radionuclides contamination of surface and
33 underground waters is present.

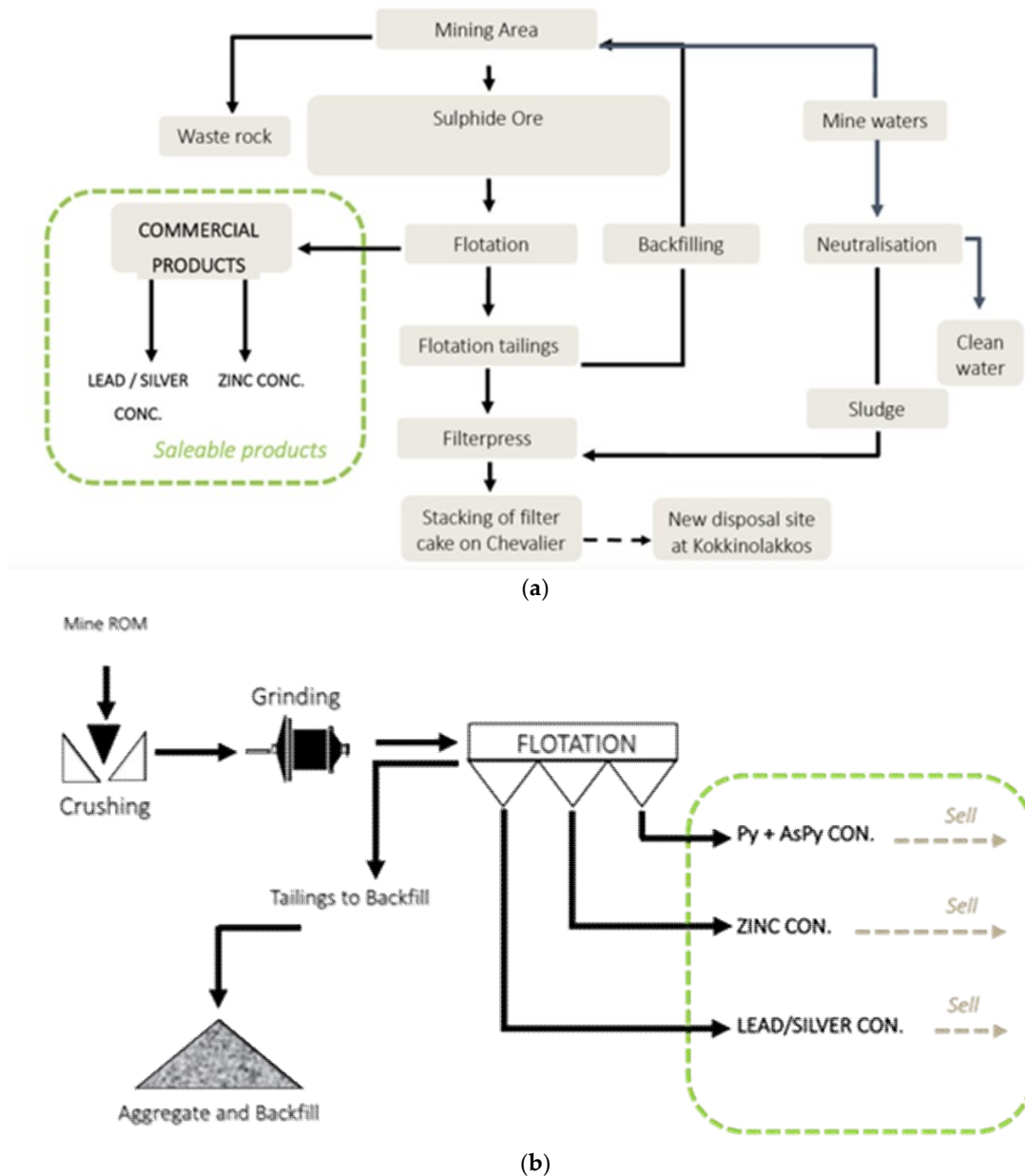
34 **Keywords:** Kassandra mines; Chalkidiki; sulphide ores; flotation concentrates; critical elements;
35 actinides
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37 1. Introduction

38 1.1 Kassandra mines flotation concentrates

39 The Kassandra mines are located in the Chalkidiki peninsula, Northern Greece. Presently
40 production is being held in Olympiada mine and Stratoni mines comprising of two deposits;
41 MademLakkos and MavresPetres. There is a long history of mining in the Kassandra area. It has
42 been estimated, from the volume of ancient slags, that about 1 million tons of ore were extracted
43 during the classical Greek period and that the Stratoni mine continued in production through the

44 Roman, Byzantine and Turkish periods. Currently, the Kassandra mines are operated by Hellenic
 45 Gold S.A. [1] and produce Pb (galena), ZnS (sphalerite) and Fe-As-S (pyrite/As-pyrite/arsenopyrite)
 46 concentrates in two flotation plants, constructed during 70s, at Stratoni and Olympiada areas (Figure
 47 1). Recently (2018), Olympiada mine has started again the production of concentrates containing Pb,
 48 Zn, Ag and Au.



49 **Figure 1.** The current flotation plant scheme in Stratoni(a) and Olympiada (b).

50 1.2 Kassandra deposits and previous work on mineralogy and geochemistry of the concentrates

51 The Kassandra mining district (Figure 2) contains porphyry Au-Cu and Au-Ag-Pb-Zn-Cu
 52 carbonate replacement deposits that are associated with Oligocene-Miocene intrusions emplaced
 53 into poly-deformed metamorphic basement rocks belonging to the Permo-Carboniferous to Late
 54 Jurassic Kerdilion unit and the Ordovician-Silurian Vertiskos unit. Regional extensional tectonics
 55 active since the middle Eocene resulted in the development of widespread normal and

56 transtensional faults, including the Stratoni fault zone that hosts carbonate replacement sulfide ore
57 bodies [2] . More particularly, Stratoni (MademLakkos, MavresPetres) and Olympiada are the two
58 main carbonate-replacement massive sulphidePb-Zn (Ag-Au) deposits of the district; they are
59 located on the footwall of the Tertiary Stratoni-Varvara fault. Both deposits are interpreted to form
60 the proximal and distal part of a fault-controlled exoskarn-type ore system triggered by nearby
61 small-scale intrusions close to the fault system [3] Sulphide mineralization occurs within
62 amphibolite grade metamorphic rocks of the Kerdylia assemblage. The assemblage represents
63 ametamorphosed marine sedimentary-volcanic sequence of probable Mesozoic or older age. Eocene
64 and Oligocene age granitic intrusions occur throughout the Kerdylia sequence, mainly as pegmatite
65 and granite dykes of several generations that range from syn- to post-metamorphic in age. The
66 sequence is affected by syn-peak metamorphic penetrative deformation that is manifested by
67 adominant, shallow dipping layer-parallel foliation. At least two other foliation-forming events
68 affect the sequence with progressively less strain, as well as significant late extensional faulting.

69 Previous workers [4] have interpreted the area to lie at the southwestern margin of the
70 Rhodope complex, and that the shallow dipping foliations which are present formed in response to
71 Tertiary unroofing of the Rhodope Complex as a metamorphic core complex. In such an
72 interpretation, the Stratoni Fault has been interpreted as the principal detachment fault forming the
73 southern major bounding structure between the Rhodope complex, represented locally by the
74 Kerdylia sequence, and the Vertiskos Formation to the south.

75 Other interpretations suggest that the fabrics are contractional and that the fault may
76 remobilize a major reverse structure that superimposed the Vertiskos sequence against the Kerdylia.
77 Geological relationships suggest that the metamorphic fabrics represent contractional rather than
78 extensional fabrics, and the Stratoni Fault as is currently manifested is dominantly a later, lower
79 greenschist grade extensional structure that is superimposed onto the amphibolite grade fabrics.

80 Mineralization at Olympiada and Stratoni (M. Lakkos-MavresPetres) is of carbonate
81 replacement. It occurs in association with a marble horizon. Mineralization is structurally late in
82 timing and is superimposed on the metamorphic fabrics in the area and in association with an
83 extensional, brittle to semi-brittle fault network that was likely active coevally with the ore-hosting
84 Stratoni Fault to the south.

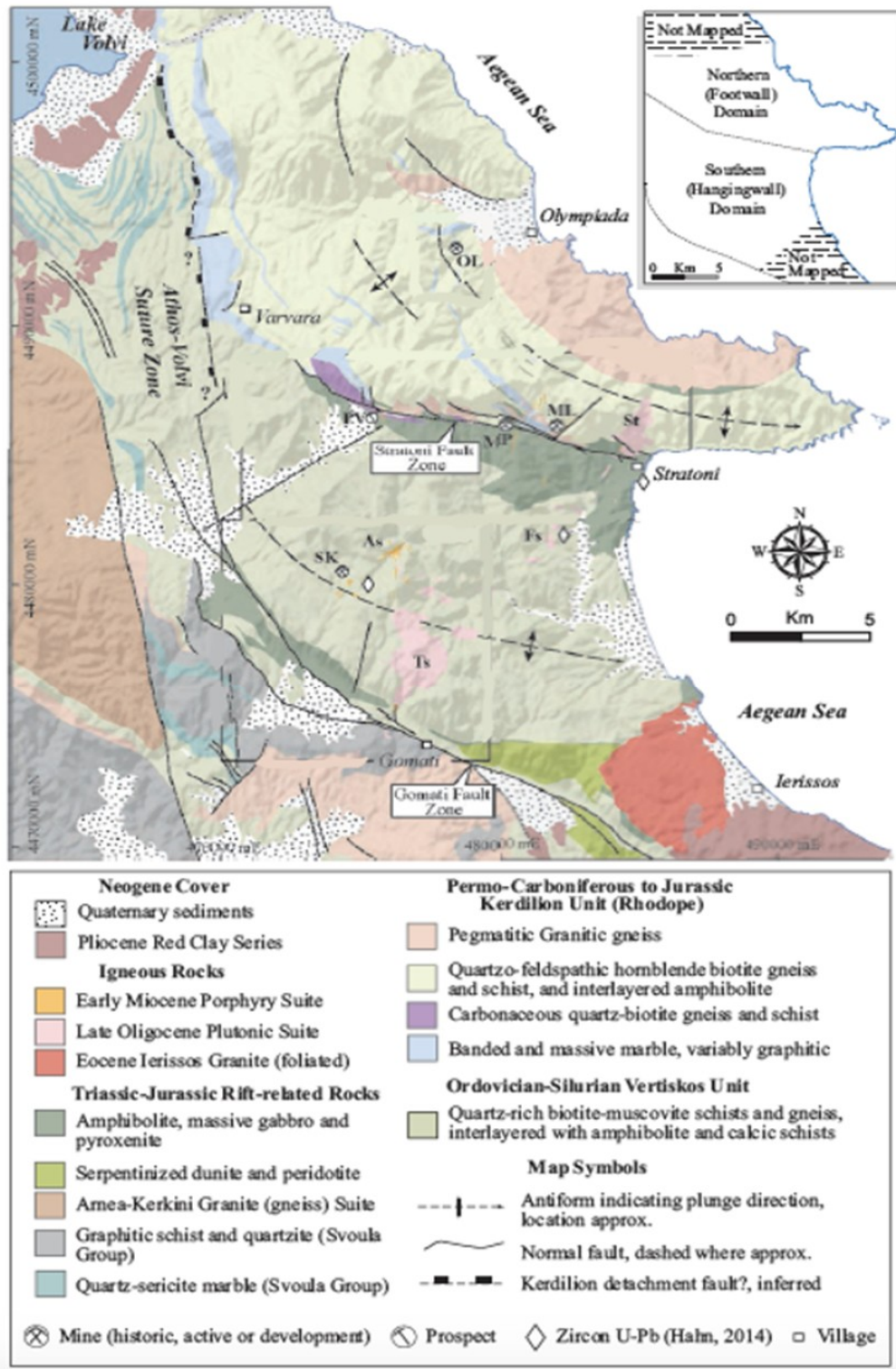


Figure 2. Geological map of the area of NE Calkidiki [2 and references therein] .

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Previous works on mineralogy and geochemistry of ores, derived from both Stratioti (MademLakkos-Mavres Peters) and Olympias mines, have been presented in the literature (e.g [2, 3, 9-15]). However, there is very limited literature about the flotation concentrates (e.g. [-20]) and particularly with respect to their mineral chemistry issues and moreover trace and critical element content.

92 1.3 Scope of the present study

93 As mentioned above, there are few published works about the concentrates, produced by
94 Kassandra mines since 70s. Additionally, on the best of our knowledge, there are no published data
95 about trace and critical elements in these hydrometallurgy (flotation) products. Thus, the scope of
96 the present study, was to demonstrate new results concerning: i) the mineral chemistry and the
97 formulae of the sulfide minerals into the concentrates, ii) the trace and critical element content,
98 specifically REE, Sb, Bi, Ge, V, Ga, Co (e.g. BGS), iii) the actinide element content (U, Th) as well as
99 their natural radioactivity. Radioisotopes present in the environment can be classified as naturally
100 occurring and are components of the earth's crust since its formation (e.g. ^{238}U , ^{235}U , ^{232}Th and their
101 decay products as well as ^{40}K), cosmogenic radioisotopes (radioisotopes that are produced by the
102 interaction between cosmic radiation and the atmosphere (e.g. ^{14}C , ^{10}Be , ^{44}Ti and ^{22}Na) and finally
103 artificially produced radionuclides that are produced in nuclear reactors (e.g. ^{90}Sr and ^{137}Cs). Natural
104 radionuclides can be found in soil, rocks, water, air, food, building materials, etc. The study of
105 natural radioactivity present in geological materials and ores is an important subject in
106 environmental radiological protection as it provides the possibility to assess any associated health
107 hazard. In this paper, the products of the Kassandra mines are studied for their natural radioactivity.
108 This involves not only the products themselves, but the associated risk from mine tailings (surface
109 water) and dissolution from underground water. Moreover, the results are explained by bulk
110 geochemistry of the samples.

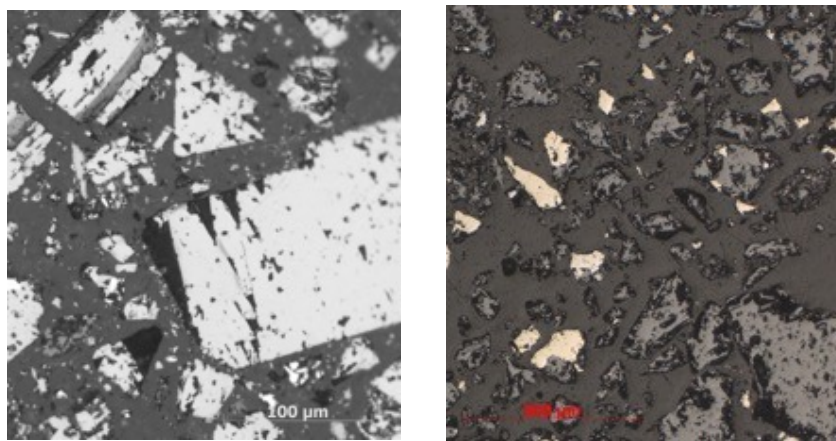
111 2. Materials and Methods

112 2.1 Samples

113 Representative composite pyrite/As-pyrite/arsenopyrite (Py-(As)Py-AsPy), galena (PbS) and
114 sphalerite (ZnS) concentrates -in powdered form-, from the flotation plants of Olympiada and
115 Stratoni (Kassandra mines, Chalkidiki, N. Greece), were supplied by Hellas Gold S.A.

116 2.2 Point analyses

117 Scanning electron microscopy (SEM) images of free mineral grains and microprobe analyses
118 (EPMA) on polished mineral grains (after examination in optical microscope -see Figure 3-) were
119 obtained using a J EOL 8200 electron probe micro-analyzer equipped with a wavelength dispersive
120 spectrometer (WDS). Analytical conditions were: 15kV accelerating voltage, 15 nA beam current, 2
121 μm beam diameter with a counting time of 20 s on the peaks and 10 s on the background.



122 **Figure 3.** Optical images (reflected light) of polished minerals grains in Stratoni PbS and ZnS
123 concentrates.

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125 2.3 Bulk analyses

126 Major and trace elements, in the powdered concentrates, were analyzed using a Perkin Elmer
127 ICP-OES and a Perkin Elmer Sciex Elan 9000 ICP-MS following a $\text{LiBO}_2/\text{LiB}_4\text{O}_7$ fusion and HNO_3
128 digestion of the fused solid sample.

129 2.4 Gamma-ray spectroscopy

130 The samples after oven-dried at 60°C to constant weight, were measured using two
131 high-resolution gamma ray spectrometry systems. The first one consisted of an HPGe coaxial
132 detector with 42% efficiency and 2.0 keV resolution at 1.33 MeV photons, shielded by 4" Pb, 1mm Cd
133 and 1mm Cu and the second one consisted of a LEGe planar detector with 0.7 keV resolution at 122
134 keV photons, shielded by 3.3" Fe-Pb, 1mm Cd and 1mm Cu. The first spectrometry system with the
135 High Purity Ge detector was used to measure the majority of the natural radionuclides examined in
136 this study, except ^{238}U . The second one with the Low Energy planar Ge detector was used so as to
137 determine only the concentration of ^{238}U , considering the low energy γ -ray of 63 keV emitted by its
138 daughter ^{234}Th .

139 The ^{40}K content was obtained using its 1461 keV γ -ray. The ^{232}Th content was calculated as the
140 weighted mean value of ^{228}Ra concentration (measured as ^{228}Ac , using 911, 968 and 338 keV γ -rays)
141 and ^{228}Th concentration (measured as decay products in equilibrium, i.e. ^{212}Pb , using 238 and 300 keV
142 γ -rays, ^{212}Bi , using 727 keV γ -ray and ^{208}Tl , using 2614, 583 and 860 keV γ -rays). The determination of
143 ^{226}Ra content was based on measurement of ^{222}Rn decay products being in equilibrium. The
144 measurement of ^{226}Ra from its own γ -ray at 186.25 keV introduces some problems because of the
145 adjacent photo peak of ^{235}U at 185.75 keV, so that the isotopic ratio between ^{235}U and ^{238}U was
146 considered being the natural one, i.e. 0.0072 and secular equilibrium between ^{238}U and ^{226}Ra had to be
147 assumed. Accuracy in the measurements of ^{226}Ra concentrations by ^{222}Rn decay products depended
148 on the integral trapping of radon gas in the sample volume, so a small addition (~2%) of charcoal in
149 powder form (less than 400 μm in size) was mixed with the sample before sealing it hermetically and
150 storing it in a freezer during ^{222}Rn in-growth period [21] . The efficiency calibration of the gamma
151 spectrometry systems was performed with the radionuclide specific efficiency method in order to
152 avoid any uncertainty in gamma ray intensities as well as the influence of coincidence summation
153 and self-absorption effects of the emitting gamma photons. A set of high quality certified reference
154 materials (RGU-1, RGTh-1, RGK-1) [22] was used, with densities similar to the average beach sands
155 measured after pulverization. Cylindrical geometry was used assuming that the radioactivity is
156 homogenously distributed in the measuring samples. The samples were measured up to 200.000 s in
157 order to achieve a Minimum Detectable Activity of 12 Bq kg^{-1} for ^{40}K , 4 Bq kg^{-1} for ^{232}Th , 2 Bq kg^{-1} for
158 ^{228}Th , 2 Bq kg^{-1} for ^{226}Ra and 21 Bq kg^{-1} for ^{238}U , with 33% uncertainty. The total uncertainty of the
159 radioactivity levels was calculated by propagation of the systematic and random errors of
160 measurements. The systematic errors in the efficiency calibration ranges from 0.3–2% and the
161 random errors of the radioactivity measurements extend up to 19 %, except in the ^{238}U measurement,
162 where the error extends up to 50% for activities measured lower 10 Bq kg^{-1} .

163 3. Results and Discussion

164 3.1 Mineral chemistry

165 The SEM and EPMA data, concerning the mineral chemistry of the sulfide minerals (major
166 phases) into the concentrates from the flotation plants of Stratoni and Olympiada mines, are given in
167 Figures 4-6 and Tables 1-3.

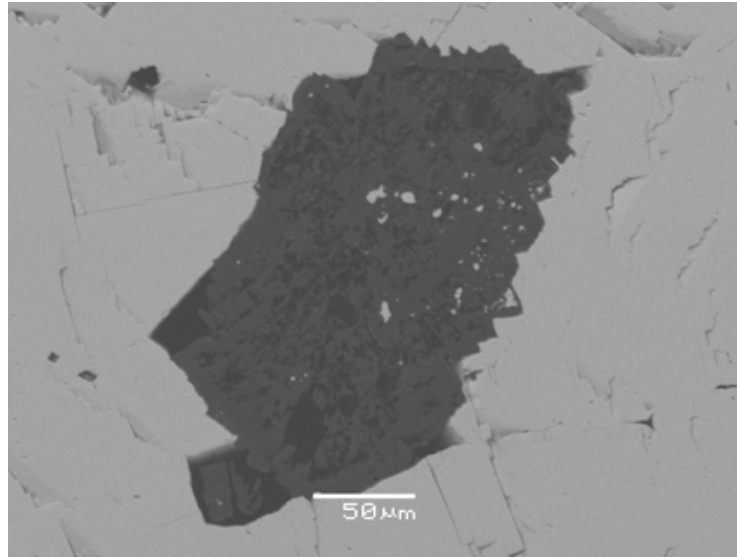
168 The chemical formulae of the major sulfide minerals were calculated as following:

- 169 • Galena (Stratoni): $\text{Pb}_{0.98-0.99}\text{S}$
- 170 • Sphalerite (Stratoni): $\text{Zn}_{0.79-0.85}\text{Fe}_{0.12-0.17}\text{Mn}_{0.00-0.01}\text{S}$
- 171 • Pyrite (Olympiada): $\text{Fe}_{1.02-1.05}\text{As}_{0.00-0.03}\text{S}_2$
- 172 • Arsenopyrite (Olympiada): $\text{FeAs}_{0.85-0.88}\text{S}_{1.07-1.13}$

173 In addition, the frequent presence of boulangerite was confirmed in the Olympiada concentrate.
174 The EPMA revealed the following chemical formula:

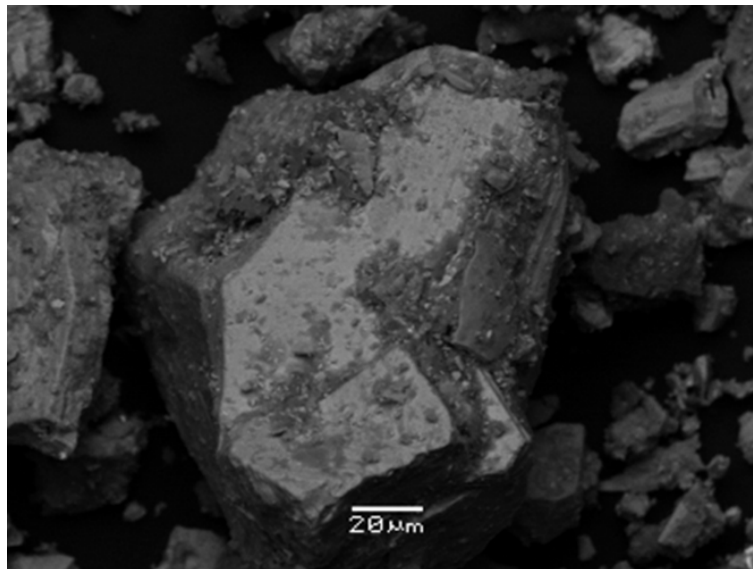
175 • Boulangerite (Olympiada): $\text{Pb}_{5.18-5.25}\text{Sb}_{4.21-4.45}\text{As}_{0.06-0.15}\text{Fe}_{0.04-0.15}\text{Zn}_{0.00-0.06}\text{Mn}_{0.01-0.02}\text{S}_{11}$

176 All sulfide minerals studied were found to exhibit typical/expected chemical compositions in
177 major elements. Ongoing research on these samples is also targeting the characterization of noble
178 metals [19, 23] .



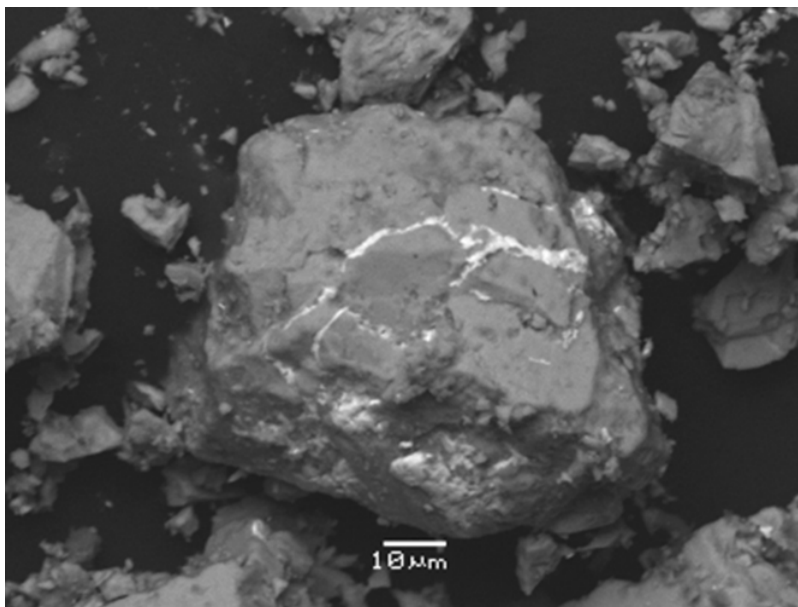
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180 **Figure 4.** BSE image of polished galena, with carbonate mineral inclusions, in Stratoni PbS
181 concentrate.



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183 **Figure 5.** BSE image of free pyrite grain (darker), including arsenopyrite (brighter), in Olympiada
184 Py-(As)Py-AsPy concentrate.



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Figure 6. BSE image of free sphalerite grain, with galena veins, in Stratoni ZnS concentrate.

Table 1. EPMA analyses of mineral phases present in Stratoni ZnS concentrate.

Phase Endmember AnalysisNo.	Sphalerite (Zn,Fe)S						Galena PbS		Arsenopyrite FeAsS	Pyrite FeS ₂		
	2	3	4	15	16	17	10	11	1	5	6	7
As	0.03	0.00	0.00	0.00	0.01	0.02	0.00	0.01	42.08	1.08	1.40	1.24
Fe	8.35	7.05	10.07	8.37	8.22	7.48	0.46	0.08	36.07	47.54	47.56	47.12
Mn	0.58	0.61	0.38	0.43	0.57	0.49	0.00	0.00	0.00	0.06	0.00	0.02
Pb	0.00	0.06	0.14	0.05	0.08	0.02	85.91	84.86	0.08	0.12	0.02	0.16
S	34.37	34.33	34.59	34.52	34.42	34.53	13.33	13.38	22.37	52.76	52.58	53.12
Zn	56.84	59.13	55.58	57.04	57.71	58.24	0.79	0.28	1.33	0.07	0.53	0.28
Total	100.18	101.18	100.75	100.41	101.01	100.79	100.48	98.61	101.94	101.63	102.09	101.95
Ionsbasedon:	1 (S)						1 (S)		1 (Fe)	2 (S)		
As	---	---	---	---	---	---	---	---	0.87	0.02	0.02	0.02
Fe	0.14	0.12	0.17	0.14	0.14	0.12	0.02	---	1.00	1.03	1.04	1.02
Mn	0.01	0.01	0.01	0.01	0.01	0.01	---	---	---	---	---	---
Pb	---	---	---	---	---	---	1.00	0.98	---	---	---	---
S	1.00	1.00	1.00	1.00	1.00	1.00	1.00	1.00	1.08	2.00	2.00	2.00
Zn	0.81	0.84	0.79	0.81	0.82	0.83	0.03	0.01	0.03	---	0.01	0.01

Table 2. EPMA analyses of mineral phases present in Stratoni PbS concentrate

Phase Endmember AnalysisNo.	Galena PbS						Arsenopyrite FeAsS				Pyrite FeS ₂			
	1	4	5	6	13	14	2	7	8	9	3	10	11	12
As	0.04	0.01	0.03	0.03	0.03	0.02	41.53	42.43	41.60	43.41	1.03	0.11	2.03	1.92
Fe	0.00	0.00	0.00	0.00	0.00	0.00	36.52	36.22	36.55	36.04	47.12	47.71	47.19	47.03
Mn	0.02	0.02	0.00	0.04	0.03	0.01	0.01	0.02	0.04	0.00	0.03	0.04	0.00	0.05
Pb	86.33	85.87	85.66	85.76	85.92	86.12	0.17	0.10	0.16	0.00	0.16	0.14	0.15	0.21
S	13.61	13.42	13.56	13.54	13.52	13.67	23.72	22.46	23.02	22.05	53.15	53.69	52.04	52.67
Zn	0.00	0.02	0.12	0.00	0.08	0.00	0.18	0.06	0.08	0.00	0.02	0.08	0.00	0.05
Total	100.01	99.33	99.37	99.37	99.58	99.82	102.13	101.29	101.45	101.50	101.50	101.76	101.41	101.93
Ionsbasedon:	1 (S)						1 (Fe)				2 (S)			
As	---	---	---	---	---	---	0.85	0.87	0.85	0.90	0.02	---	0.03	0.03
Fe	---	---	---	---	---	---	1.00	1.00	1.00	1.00	1.02	1.02	1.04	1.03
Mn	---	---	---	---	---	---	---	---	---	---	---	---	---	---
Pb	0.98	0.99	0.98	0.98	0.98	0.98	---	---	---	---	---	---	---	---
S	1.00	1.00	1.00	1.00	1.00	1.00	1.13	1.08	1.10	1.07	2.00	2.00	2.00	2.00
Zn	---	---	---	---	---	---	---	---	---	---	---	---	---	---

Table 3. EPMA analyses of mineral phases present in Olympiada Py-(As)Py-AsPy concentrate

Phase Endmember AnalysisNo.	Sphalerite (Zn,Fe)S	Boulangerite Pb ₅ Sb ₄ S ₁₁			Arsenopyrite FeAsS						Pyrite FeS ₂					
	1	13	14	15	2	3	16	17	18	20	5	6	7	19	21	22
As	0.00	0.22	0.58	0.35	41.02	41.35	41.07	41.22	41.43	41.18	2.03	0.02	1.43	0.89	1.68	1.24
Fe	11.28	0.11	0.43	0.15	36.41	36.77	36.52	36.55	36.85	36.57	47.42	47.82	47.25	47.55	47.74	47.33
Mn	0.61	0.03	0.06	0.06	0.03	0.02	0.03	0.03	0.22	0.23	0.00	0.00	0.00	0.00	0.00	0.00
Pb	0.10	54.76	54.38	56.25	0.04	0.07	0.05	0.06	0.05	0.04	0.17	0.14	0.12	0.15	0.13	0.16
S	34.55	17.99	17.66	18.23	23.47	22.73	22.84	23.15	22.36	22.94	52.03	53.90	52.45	53.11	52.17	52.41
Zn	53.64	0.21	0.05	0.00	0.10	0.16	0.13	0.15	0.17	0.15	0.00	0.04	0.00	0.03	0.00	0.02
Sb	n/a ¹	27.08	27.12	26.51	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a
Total	100.17	100.40	100.27	101.55	101.07	101.10	100.64	101.16	101.08	101.10	101.65	101.92	101.25	101.74	101.72	101.16
Ionsbasedon:	1 (S)	11 (S)			1 (Fe)						2 (S)					
As	---	0.06	0.15	0.09	0.84	0.84	0.84	0.84	0.84	0.84	0.03	---	0.02	0.01	0.03	0.02
Fe	0.19	0.04	0.15	0.05	1.00	1.00	1.00	1.00	1.00	1.00	1.05	1.02	1.03	1.03	1.05	1.04
Mn	0.01	0.01	0.02	0.02	---	---	---	0.00	0.01	0.01	---	---	---	---	---	---
Pb	---	5.18	5.24	5.25	---	---	---	---	---	---	---	---	---	---	---	---
S	1.00	11.00	11.00	11.00	1.12	1.08	1.09	1.10	1.06	1.09	2.00	2.00	2.00	2.00	2.00	2.00
Zn	0.76	0.06	0.01	---	---	---	---	---	---	---	---	---	---	---	---	---
Sb	n/a	4.36	4.45	4.21	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a

¹n/a=not analyzed.

192 3.2 Geochemistry

193 The bulk chemical compositions (ICP-OES & MS) of the studied Kassandra mines concentrates
 194 are given in Table 4. It is obvious that basic and noble metals (Pb, Zn, Ag, and Au) being exploited by
 195 Hellas Gold S.A., show high concentrations, as well as Sb and Ga (Sb: >0.2 wt.% in PbS concentrate;
 196 Ga: 25 ppm in ZnS concentrate). On the other hand, there are no considerable contents of Bi, Co, V,
 197 and REE. Considering other elements, As is found in elevated concentrations (>1 wt.% in
 198 Py-(As)Py-AsPyOlympiada concentrate and almost 1 wt.% in StratoniPbS and ZnS concentrates)
 199 along with Cd (specifically in ZnS concentrate). Moreover, actinides occur in very low
 200 concentrations (U <2 ppm and Th<0.5 ppm in all concentrates).

201 The enrichment and depletion of the elements studied can be revealed from the normalization
 202 to the Upper Crust (UCC) (Figure 7), the Primitive Mantle (Figure 8) and the Chondrite (Figure 9).
 203 REE's and other elements like Cs, Rb, Co, Ni, Ba and V are depleted. As expected, major elements
 204 like Pb, Zn and Cu are enriched, as well as other trace elements like Mo, As, Sb, Se, Sn, Cd, Hg, Rb
 205 and U. It should be noted that the enrichment in these trace elements, relative to UCC, Primitive
 206 Mantle and Chondrite, is strictly geochemical and it is not associated to practical mining and
 207 metallurgical issues. For instance, if we consider U, the bulk natural radioactivity of the samples is
 208 negligible, as discussed below.

 209 **Table 4.** Trace and critical elements concentration of the studied Kassandra mines concentrates.

Element	MDL (ppm)	Py-(As)Py-AsPy Conc.	PbS Conc.	ZnS Conc.
		OLYMPIADA	STRATONI	STRATONI
Pb	0.1	5235.9	>10000.0	>10000.0
Zn	1	>10000	>10000	>10000
Ag	0.1	22.5	>100.0	>100.0
Au	0.0005	16.9	1.1	1.0
Cu	0.1	710.8	1035.9	2191.1
As	0.5	>10000.0	>10000.0	9476.9
Sb	0.1	712.9	>2000.0	748.2
Bi	0.1	0.1	2.0	0.2
Cd	0.1	54.5	77.7	>2000.0
Ni	0.1	9.3	6.6	3.1
Co	0.2	1.7	0.5	0.2
Hg	0.01	0.15	0.49	10.98
Tl	0.1	1.2	37.2	3.4
Se	0.5	2.7	8.3	<0.5
Be	1	<1	<1	<1
Th	0.2	0.5	<0.2	<0.2
U	0.1	1.4	2.0	1.1
Sn	1	43	131	218
Mo	0.1	2.6	29.9	5.7
Ga	0.5	4.5	1.5	25.4
V	8	12	9	<8
Nb	0.1	1.0	<0.1	<0.1
Ta	0.1	<0.1	<0.1	<0.1

W	0.5	1.8	1.0	0.7
Ba	1	15	8	2
Cs	0.1	0.4	0.2	0.3
Hf	0.1	0.2	<0.1	<0.1
Rb	0.1	6.9	2.1	3.1
Sr	0.5	7.1	1.5	1.9
Zr	0.1	3.9	2.0	1.9
Y	0.1	0.5	0.2	0.1
La	0.1	1.2	<0.1	0.4
Ce	0.1	2.0	0.8	1.1
Pr	0.02	0.20	0.09	0.10
Nd	0.3	0.8	<0.3	<0.3
Sm	0.05	0.12	<0.05	<0.05
Eu	0.02	0.03	<0.02	<0.02
Gd	0.05	0.16	<0.05	0.08
Tb	0.01	0.02	<0.01	<0.01
Dy	0.05	0.10	<0.05	<0.05
Ho	0.02	<0.02	<0.02	<0.02
Er	0.03	0.04	<0.03	<0.03
Tm	0.01	0.01	<0.01	<0.01
Yb	0.05	<0.05	<0.05	<0.05
Lu	0.01	0.01	<0.01	<0.01

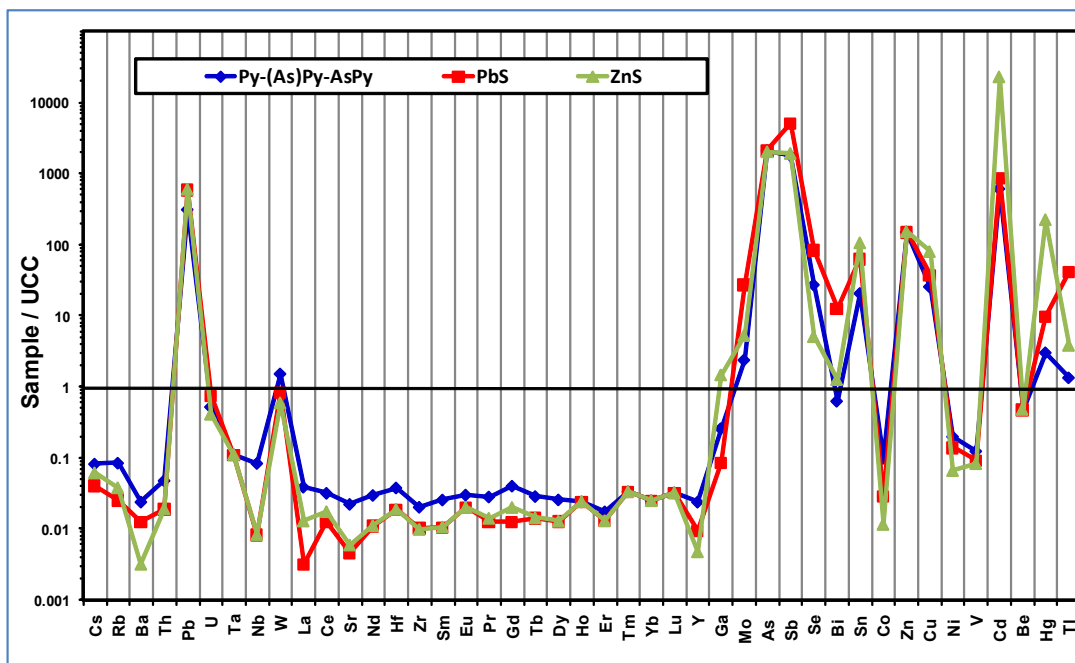


Figure 7. Spider diagram of Sample/Upper Continental Crust (UCC).

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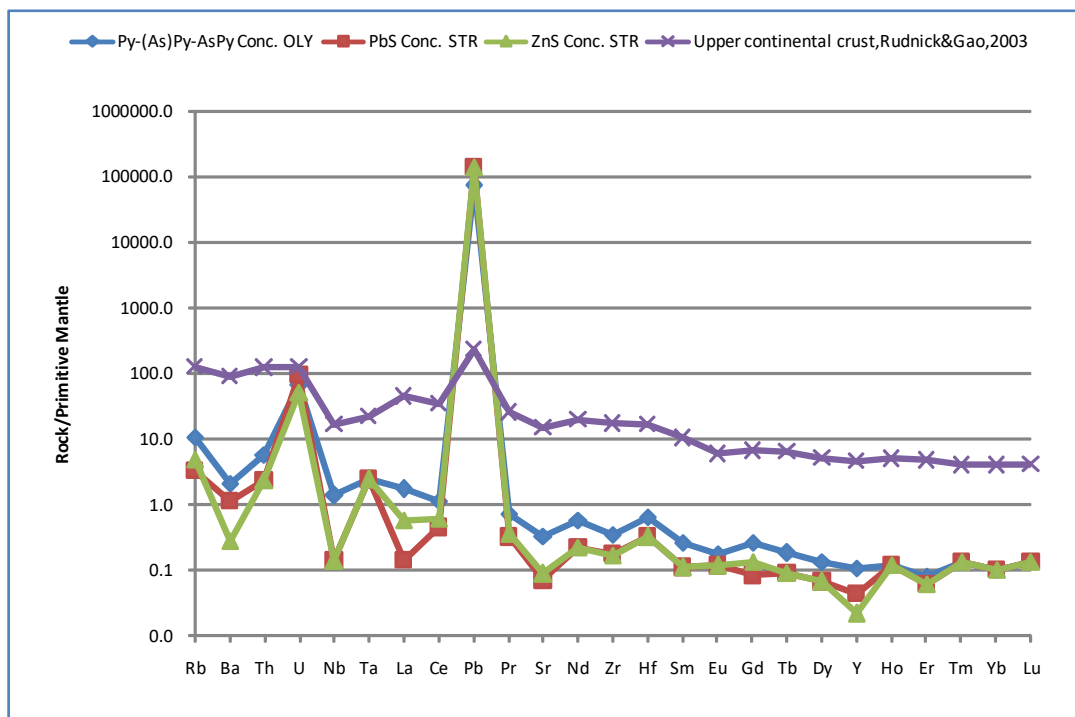


Figure 8. Spider diagram of Sample/Primitive Mantle.

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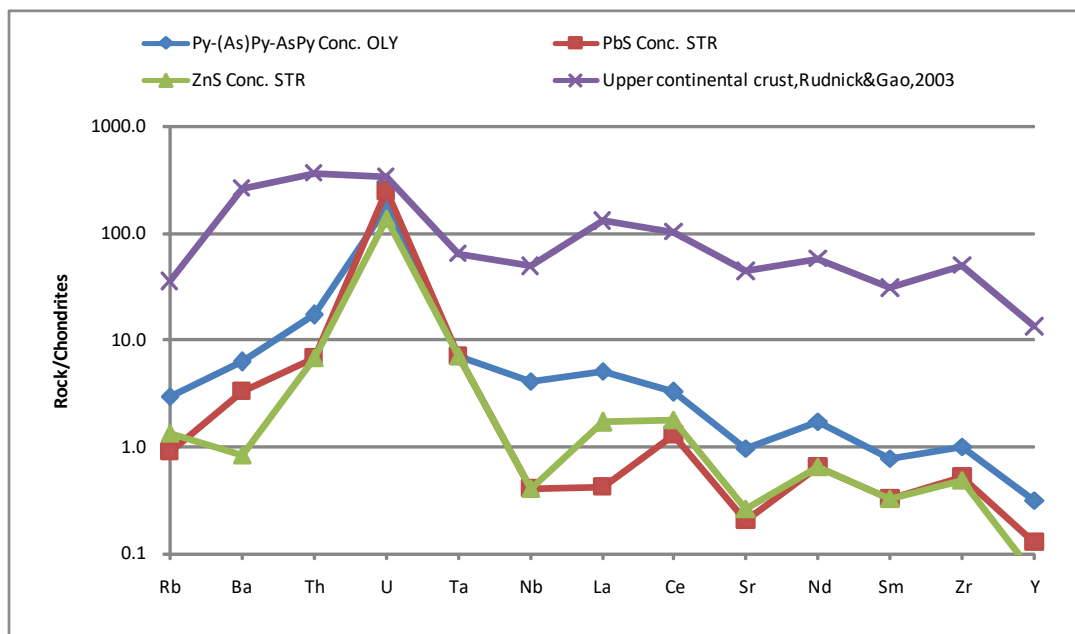


Figure 9. Spider diagram of Sample/Chondrite.

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216 3.3 Actinide elements and natural radioactivity

217 The concentrations of the natural radionuclides, detected by gamma-ray spectroscopy, are
218 given in Table 5.

219 **Table 5.** Activity concentrations of ^{238}U , ^{226}Ra , ^{232}Th , ^{228}Th and ^{40}K (Bq.kg^{-1}), along with the respective
220 standard errors ($\pm\sigma$).

Sample	^{238}U -series				^{232}Th -series				^{40}K		^{226}Ra	^{232}Th
	^{238}U		^{226}Ra		^{232}Th		^{228}Th		^{40}K		^{226}Ra	^{232}Th
	Bq.kg^{-1}	$\pm\sigma$	Bq.kg^{-1}	$\pm\sigma$	Bq.kg^{-1}	$\pm\sigma$	Bq.kg^{-1}	$\pm\sigma$	Bq.kg^{-1}	$\pm\sigma$	ppm	ppm
OLY-C-(FeAsS)	B.D.L.	-	24.8	0.3	4.1	0.7	3.3	0.2	51.5	2.9	2.0	1.0
STR-(PbS)	B.D.L.	-	22.0	0.3	1.0	0.5	B.D.L.	-	10.8	2.1	1.8	0.2
STR-(ZnS)	B.D.L.	-	19.1	0.4	1.3	0.8	B.D.L.	-	18.2	2.8	1.5	0.3

221 The concentrations of the radionuclides of ^{238}U , ^{232}Th -series and ^{40}K are small and close to the
222 detection limit of gamma-ray spectroscopy. These small concentrations are mainly due to the small
223 ability of the chemical components of the sulphides to be substituted by the measured radionuclides.
224 Moreover, low concentrations of these radionuclides have been detected in the Stratoni granitic
225 bodies [24] . Similar conclusions on the content of sulphides have been previously reported by [25,
226 26] . However, the previous researchers mention that high U concentrations may be present in the
227 late accessory mineral phases deposited in microfissures.

228 These values are by far lower than a typical granitic rock used as building material [26] .
229 Therefore, the radioactive dose to humans from these materials is insignificant.

230 4. Conclusions

231 The results of the present study can be summarized as follows:

- 232 • Except basic (Pb, Zn, and potentially Cu) and noble (Ag, Au) elements in Kassandra mines
233 concentrates, being exploited and commercialized by Hellas Gold S.A., it can be argued that
234 there are also significant concentrations of Sb and Ga (Sb: >0.2 wt.% in PbS concentrate; Ga: 25
235 ppm in ZnS concentrate), but no substantial contents of Bi, Co, V, and REE.
- 236 • Concerning other elements, of course it well-known that As occurs in rather high concentrations
237 (>1 wt.% in Py-(As)Py-AsPy Olympias concentrate and almost 1 wt.% in Stratoni PbS and ZnS
238 concentrates), as well as Cd (specifically in ZnS concentrate).
- 239 • There are negligible concentrations of actinides (U <2 ppm and Th <0.5 ppm in all concentrates),
240 minimizing the possibility of increased natural radioactivity. The concentrations of natural
241 radionuclides are by far lower than a typical granitic rock used as building material
242 (Papadopoulos et al. 2013). Therefore, the radioactive dose to humans from these materials is
243 insignificant.

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