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OPEN Non-destructive analysis of a mixed H₂O-CO₂ fluid in experimental noble-metal capsule by means of freezing and high-energy synchrotron X-ray diffraction

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High-pressure high-temperature syntheses that involve volatile-bearing aqueous fluids are typically accomplished by enclosing the samples in gas-tight welded shut noble-metal capsules, from which the bulk volatile content must be extracted to be analyzed with mass spectroscopy, hence making the analysis non-replicable. Here we describe a novel non-destructive method that ensures the identification and the quantitative estimate of the volatiles directly in the sealed capsule, focusing on fluid H₂O-CO₂ mixtures equilibrated with graphite at conditions of geological interest (1 GPa, 800 °C). We used a high-energy (77 keV) synchrotron X-ray radiation combined with a cryostat to produce X-ray diffraction patterns and X-ray diffraction microtomographic cross-sections of the volatilebearing samples down to -180 °C, thus encompassing the conditions at which crystalline phases-solid CO₂ and clathrate (CO₂ hydrate)-form. The uncertainty of the method is < 15 mol%, which reflects the difference between the volatile proportion estimated by both Rietveld refinement of the diffraction data and by image analysis of the microtomograms, and the reference value measured by quadrupole mass spectrometry. Therefore, our method can be reliably applied to the analysis of frozen H₂O-CO₂ mixtures and, moreover, has the potential to be extended to experimental fluids of geological interest containing other volatiles, such as CH₄, SO₂ and H₂S.

Carbon dioxide, the main Earth's greenhouse gas, is cycled between the solid Earth and the atmosphere over long timescales (> 1 Ma) by a variety of natural processes. Diffuse degassing and volcanic emissions are the final link in a long chain of devolatilization reactions occurring in the Earth's interior¹. The oxidation of the organic carbonaceous matter (soon transformed into graphitic carbon) and the decarbonation of sedimentary carbonate minerals² in the subducted lithosphere are probably the most important processes supplying CO₂ to the deep carbon cycle. This has been shown for instance in high-pressure high-temperature experiments reproducing the environmental conditions where CO₂-releasing reactions are thought to occur at depth³⁻⁵. Recent advances in the synthesis and analysis of CO₂-bearing aqueous fluids produced by oxidation of graphite at controlled redox conditions allowed to explore a wide range of compositional systems. In the C-O-H, the most simple system describing carbon-bearing aqueous fluids, it has been demonstrated that the composition of fluids in equilibrium with pure graphite is consistent with thermodynamic predictions⁶⁻⁸. However, an unpredictable oxidation behavior of graphite has been described in more complex systems containing contain silicates³, carbonates⁵, disordered forms of graphite⁶ or disordered redox-buffering materials (e.g., NiO⁸). In these experimental studies, the redox state during the fluid synthesis has invariably been constrained using the so-called double-capsule technique⁹, which has been proven to be robust and reliable. Double capsules consist of an inner welded shut capsule, usually made of a noble-metal alloy permeable to hydrogen (e.g., Pt- or Pd-bearing alloys) and containing the experimental charge, and an outer welded shut capsule containing a redox-buffering mineral assemblage (usually metal + metal oxide or metal oxide + metal silicate, where the metal has different oxidation states) soaked in water. During the run at high pressure and temperature, redox reactions in the buffering mineral assemblage coupled with the dissociation of water constrain the fO_2 and the fH_2 in the outer capsule. Because of the permeability to

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 H_2 , the same fH_2 is expected in the inner capsule as well. When dealing with graphite oxidation, fH_2 controls the amount of CO_2 produced in the inner capsule containing graphite and water (see "Methods").

The composition of the final fluid is typically measured by puncturing the inner experimental capsule, conveying the volatiles towards a quadrupole mass spectrometer⁷, so that the fluid phase is eventually lost and unavailable for replicated analyses. Moreover, to set up the instrument for quantitative analyses, the species forming the fluid (e.g., H₂O, CO₂, CO, CH₄) must be established in advance. In this study, we will describe a non-destructive method to qualitatively and quantitatively analyze the fluids contained in a noble-metal experimental capsule that relies on freezing the sample and analyzing it by X-ray diffraction, using a high-energy synchrotron source. This is made possible due to the fact that the solid H₂O-CO₂ mixture is crystalline when the temperature is below the freezing points of all the species that compose the fluid¹⁰. We will show that the 77 keV X-ray source, available at the ID15A line of the European Synchrotron Radiation Facility (ESRF; Grenoble, France), can effectively be used to obtain diffraction data and micro-tomographic cross sections of the inner capsule containing graphite and a H₂O-CO₂ fluid. We will then compare the results with the conventional (destructive) quadrupole mass spectrometry analysis of the fluid to estimate the reliability of our method.

Methods

Synthesis of fluids and sample preparation. Mixed H_2O-CO_2 fluids were generated by oxidation of graphite powder (obtained from spectroscopic pure rods gently hand-ground in boron carbide mortar) in ultrapure MilliQ water. Experiments were buffered using the double-capsule technique⁹ to prevent the direct contact of the sample with the buffering assemblages. An inner H_2 -permeable $Au_{60}Pd_{40}$ capsule (outer diameter OD=2.3 mm; inner diameter ID=2.0 mm; length ≈ 8 mm) and an outer Au capsule (OD=3 mm; ID=2.8 mm) were used. The outer capsule contained the redox buffer fayalite-magnetite-quartz (FMQ) soaked in water, constraining fH_2 . Quartz (SiO_2) was obtained from powdered natural crystals; fayalite (Fe_2SiO_4) and magnetite (Fe_3O_4) have been synthesized at 1100 °C in a gas-mixing furnace in reducing atmosphere ($CO_2:CO=10:1$), starting from stoichiometrically weighted reagent-grade Fe_2O_3 (Sigma-Aldrich) and amorphous SiO_2 obtained from hydrolyzed tetraethyl orthosilicate (Sigma-Aldrich). At equilibrium conditions, as long as all the buffer phases are present – which has been confirmed by scanning electron microscopy in almost identical runs performed at the same P–T conditions $^{3.6}$ – the chemical potential of hydrogen is expected to be homogeneous in the inner and in the outer capsules. In the outer capsule, the hydrogen fugacity (fH_2) is constrained through the reactions:

$$3Fe_2SiO_4(fayalite) + 2H_2O = 2Fe_3O_4(magnetite) + 3SiO_2(quartz) + 2H_2(1 GPa, 800 °C)$$
 (1)

In the inner capsule, the equilibration of the COH fluid is accomplished by the fH_2 -dependent reaction³:

$$C(graphite) + 2H_2O = CO_2 + 2H_2$$
 (2)

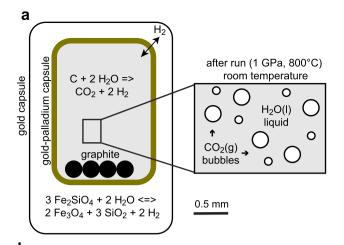
As a consequence, the initial CO_2 -free aqueous fluid adjusts its CO_2/H_2O fraction until the equilibrium with the fH_2 imposed by the buffer is reached. In a similar fashion, the oxygen fugacity (fO_2) is constrained directly in the outer capsule by the FMQ buffer and indirectly in the inner capsule because of the water dissociation reaction:

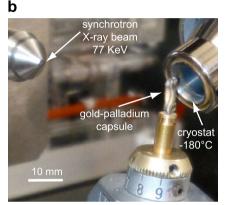
$$2H2O = 2H2 + O2$$
 (3)

In the inner capsule, however, the fO_2 will be slightly lower compared to that imposed by the FMQ buffer in the outer capsule since the fluid is not pure H_2O but a H_2O-CO_2 mixture, with a consequent declined fugacity for H_2O (and O_2)¹¹.

Experiments were performed at 1 GPa at 800 °C using an end-loaded piston-cylinder apparatus. Capsules were embedded in MgO rods (Norton Ceramics) and inserted in graphite furnaces surrounded by NaCl and borosilicate glass (Pyrex*). At the top of the assembly, a pyrophyllite–steel plug was placed to ensure the electrical contact. Temperatures were measured with K-type thermocouples that have an estimated uncertainty of \pm 5 °C. An alumina disk was placed at the top of the capsule to avoid the direct contact with the thermocouple. Pressure calibration of the apparatus is based on the quartz to coesite transition¹² that guarantees an uncertainty of \pm 0.01 GPa. Samples were first brought to the run pressure, then heated to 800 °C, with a ramp of 100 °C/min. Run duration was 92 h, in order to approach equilibrium between graphite and CO₂-bearing aqueous fluid at 800°C³. Eventually, the experiments were quenched by turning off the power supply, resulting in a cooling rate of > 40 °C/sec.

X-ray diffraction of the frozen fluid. The inner Au–Pd capsule (Fig. 1a,b) containing graphite and H_2O-CO_2 fluid (Fig. 1c) was exposed by peeling off the outer gold capsule and removing the redox buffer (FMQ), then mounted on a goniometer head on the sample stage at the beamline ID15A of the ESRF (Fig. 1b). A Debye transmission geometry was used, and the signal was collected with a Dectris Pilatus3 X 2 M detector with a Cd–Te sensor. The beam had an energy of 77 keV (λ =0.1610 Å) and a spot size of 0.05 × 0.05 mm on the sample. Such an elevated beam energy coupled with a high energy threshold of the detector minimized the absorption effects from the capsule (Au X-ray absorption edge is located at ~80.7 keV) and allowed a significant reduction of the background (mainly arising from Au₆₀Pd₄₀ fluorescence), respectively. The X-ray diffraction effects were collected every 2 min while cooling the sample with a cryostat from room temperature down to – 90 °C and from – 90 to – 180 °C with rates of – 3 °C/min and – 6 °C/min, respectively (Fig. 2a). The beam was focused orthogonally to the axis of the rotating capsule and multiple diffraction pattern collections were made at different z-axis positions. For each temperature, multiple scans with z offsets of 0.05 mm in 1 mm z interval were collected. Phase abundances were quantitatively estimated by performing Rietveld refinement and the propor-





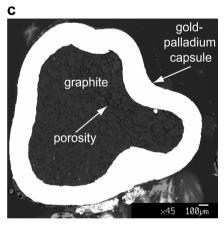


Figure 1. Experimental and analytical setup. (a) Sketch of the experimental capsule, consisting of an outer gold capsule containing an fH_2 – fO_2 buffering assemblage (FMQ = fayalite + magnetite + quartz + water) and an inner gold–palladium capsule (brown thick line), permeable to hydrogen, containing graphite + water. At run conditions of 1 GPa, 800 °C and fH_2 ^{FMQ}, the oxidation of graphite produces CO_2 . (b) The inner gold–palladium capsule containing graphite and a mixed H_2O-CO_2 fluid is mounted on a goniometer head and cooled down to - 180 °C to allow the fluid to freeze into solid phases; X-ray diffraction effects inside the noble-metal capsule are collected using a high-energy 77 keV synchrotron source at ID15A of the ESRF. (c) Back-scattered electron image of a cross section of the inner capsule (white), mounted in epoxy resin and polished after synchrotron and QMS measurements; the former presence of a fluid phase can be inferred from the porosity between the graphite grains.

tions obtained from multiple acquisitions were used to calculate average values and the related uncertainties. Rietveld analysis was performed with GSAS+EXPGUI software^{13,14}. Crystal structure of graphite, solid CO₂, ice and clathrate were taken from ICSD database. Background, phase scale factors, lattice parameters and two peak

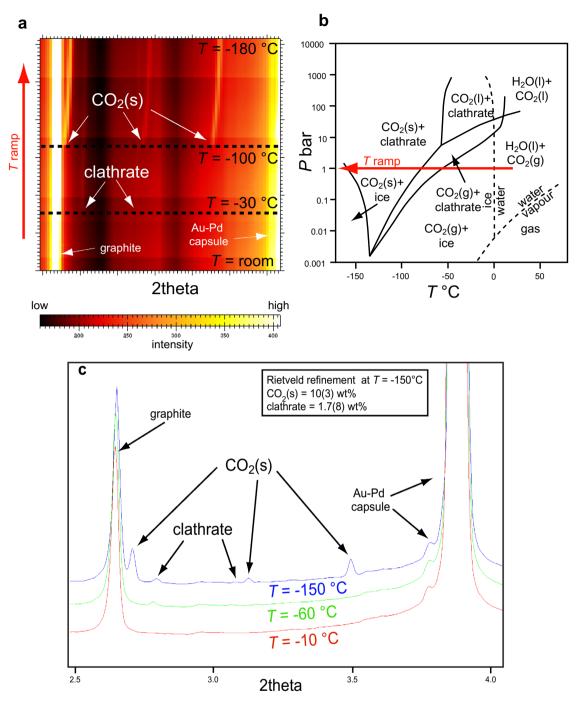


Figure 2. X-ray diffraction data collected during the cooling ramp from room temperature down to $-180\,^{\circ}\text{C}$. (a) Clathrate and solid CO $_2$ diffraction effects are detected below $-30\,^{\circ}\text{C}$ and $-100\,^{\circ}\text{C}$, respectively. (b) Low-temperature phase diagram for the H_2O-CO_2 system redrawn after 10 ; the experimental cooling ramp is shown by a red arrow. (c) Representative X-ray diffraction patterns collected at $-10\,^{\circ}\text{C}$, $-60\,^{\circ}\text{C}$ and $-150\,^{\circ}\text{C}$; the Rietveld refinement of the $-150\,^{\circ}\text{C}$ pattern yields 10 ± 3 wt% CO $_{2(\text{solid})}$ and 1.7 ± 0.8 wt% clathrate (a 100 wt% closure is imposed on the sum of CO $_{2(\text{solid})}$ + clathrate + graphite).

profile function parameters (the constant Gaussian term and a term for Lorentian of a pseudo-Voigt function), were refined.

X-ray diffraction microtomography¹⁵ was used to acquire a cross section of the middle part of the capsule at $T=-180\,^{\circ}\text{C}$ (Fig. 3), well below the freezing points of both $H_2\text{O}$ and CO_2 . An RGB image showing the spatial distribution of the phases in the capsule was obtained by assigning to the color channels the characteristic peaks of solid CO_2 , clathrate and graphite (red, green and blue, respectively, in Fig. 3; cf. also Supplementary Fig. 1). Background removal was applied to eliminate noise. Pixels of the RGB image were also classified using a K-means clustering method (Wolfram Mathematica).

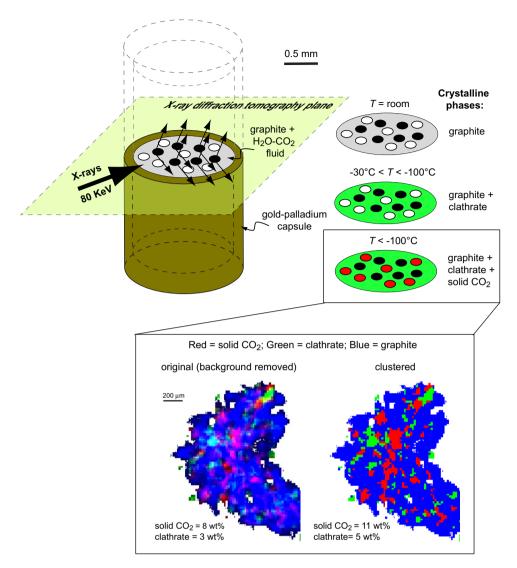


Figure 3. Schematic representation of X-ray diffraction microtomography of the experimental capsule (left) and the phase transitions observed in the H_2O-CO_2 fluid upon cooling (right). The inset shows the microtomogram collected at -180 °C, i.e., when the fluid is converted into solid crystalline clathrate and $CO_{2(s)}$: red, green and blue channels have been attributed to the characteristic peaks of solid CO_2 , clathrate and graphite, respectively, and combined into an RGB image after background removal (left) or clustering (right); the estimated mass percentages of solid CO_2 and clathrate are reported.

Quadrupole mass spectrometry. The capsule-piercing technique⁷ was employed for the analysis of quenched volatiles: the inner capsule was first punctured in vacuum conditions in a Teflon® reactor maintained at $T \approx 90$ °C (at this temperature the liquid water contained in the sample evaporates completely); then, the gases were conveyed by an ultrapure Ar flux to an EXTORR 0–200 amu, Mod. XT200, quadrupole mass spectrometer (QMS), equipped with a secondary electron multiplier. The pressure and temperature conditions in the reactor were monitored by means of high-resolution pressure gauges (with a precision of ± 1 mbar) and a K-type thermocouple, respectively. Gas mixtures of known composition and ultrapure water were utilized as standards for the calibration of the QMS. The stability and the reproducibility of the analyses over time were monitored by periodically analyzing the volatile products released by thermal decomposition at 250 °C of 1 mg oxalic acid dihydrate contained in test capsules. In this study the technique enabled to quantify micromolar quantities of the volatiles H_2O and CO_2 with uncertainties of 1.2% and 0.45%, respectively (Supplementary Table 1).

Results

By cooling the sample, the volatiles contained in the capsule, composed of a mixture of water and gaseous CO_2 , crystallized into a mixture of clathrate $(CO_2 \cdot nH_2O)$ and solid CO_2 (Fig. 2b). From room temperature to $-30\,^{\circ}C$, only the peaks of graphite and of the Au–Pd alloy are detectable in the diffraction patterns; at $T = -30\,^{\circ}C$ the peaks of a phase ascribable to clathrate appear and at $T \le -100\,^{\circ}C$ they coexist with the peaks of solid CO_2 . The diffraction patterns remain almost unchanged down to $-180\,^{\circ}C$, with the only exception of a subtle shift of the

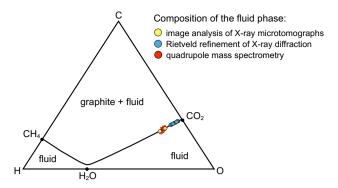


Figure 4. Fluid composition estimates (in wt%) plotted on a ternary C-O-H phase diagram (calculated at 1 GPa and 800 °C). The results of Rietveld refinement of single X-ray z-scans (blue dots) and image analysis (yellow dots) are compared with the reference fluid composition measured by QMS (red cloud). The black thick line represents all the possible compositions of graphite-saturated C-O-H fluids.

solid CO_2 peaks towards higher diffraction angles. Apart from a slightly anticipated crystallization of water to clathrate (predicted at $T \approx -60$ °C) and a slightly retarded crystallization of $CO_{2(gas)}$ to solid CO_2 (predicted at $T \approx -80$ °C), our findings agree with proposed phase diagrams in the H_2O-CO_2 system¹⁰ (Fig. 2b), assuming that the pressure in the capsule is close to room conditions (red arrow in Fig. 2b). Representative X-ray diffraction patterns, collected at T = -10 °C, -60 °C and -150 °C, are displayed in Fig. 2 c. Rietveld analyses showed slight inhomogeneities in the relative abundance of solid CO_2 and clathrate (Supplementary Table 1), which can be averaged to 10 ± 3 wt% and 1.7 ± 0.8 wt%, respectively.

Microtomography of a single cross-section shows the micron-scale distribution in the capsule of the formerly volatile phases (Fig. 3); clathrate and solid CO_2 fill the open porosity between graphite grains. By counting the average RGB pixel values, where $R = \text{solid } CO_2$, G = clathrate and B = graphite, the volume percentages of solid CO_2 and clathrate are estimated to 10 vol% and 6 vol%, respectively. Mass percentages are then calculated to be 8 wt% $CO_{2(\text{solid})}$ and 3-4 wt% clathrate, by considering that the density of solid CO_2 is 1.68 g/cm^3 , while that of clathrate is in the range $1.18-1.36 \text{ g/cm}^3$, depending on whether n is assumed to be 5.75 or 7^{16} . A second way to estimate volume percentages is by means of image analysis (clustering), which allows to attribute every pixel of the microtomogram to a single phase. With this method, the volume percentages of solid CO_2 and clathrate are 14 vol% and 9 vol%, respectively, which correspond to mass percentages of 11 wt% $CO_{2(\text{solid})}$ and 5-6 wt% clathrate.

Validation of the method. Destructive QMS analysis of the volatiles extracted from the experimental capsule provides the reference values for the abundances of CO_2 and H_2O to be compared with the estimates obtained from X-ray diffraction of the frozen fluid. Mass spectrometry measurements give $12.66\pm0.01~\mu mol~CO_{2(gas)}$ and $7.80\pm0.02~\mu mol~H_2O_{(vapour)}$. By considering the presence of 7.87~mg of graphite in the capsule, these values can be converted to mass percentages: $6.50\pm0.03~mt$ wt% CO_2 and $1.64\pm0.02~mt$ Wt% H_2O , respectively, with a $CO_2/(H_2O+CO_2)_{mass}=0.80\pm0.05$. A complication arises when the mass percentages of solid CO_2 and clathrate derived by both Rietveld refinement of X-ray diffraction patterns and image analysis of the microtomogram are compared with the mass percentages of CO_2 and CO_2 are therefore: (i) CO_2 and CO_2 and CO_2 and CO_2 are therefore: (i) CO_2 and CO_2 and CO

In Fig. 4 it is shown the comparison between the $CO_2/(H_2O+CO_2)$ mass ratios obtained from X-ray diffraction (blue dots), QMS (red cloud) and image analyses (yellow dots). Following this approach, the volatile components proportion from microtomogram analysis $[CO_2/(H_2O+CO_2)\sim 0.8]$ is consistent to that from QMS (0.80 ± 0.05) within uncertainties. Rietveld refinement of X-ray diffraction data, considering spatial variability, low absolute content of clathrate ≈ 1 wt%, possible uncertainties related to clathrate structure, and especially possible non randomly oriented crystallite and/or large crystal grains, all critical parameters in quantitative analysis 17,18 , provide also consistent values from 0.85 to 0.95: the average value of ~ 0.92 is only 15% higher than the reference QMS value.

Conclusions

We demonstrate that high-energy synchrotron X-rays can be used to analyze – qualitatively and quantitatively-a frozen fluid phase sealed into a noble-metal experimental capsule. Temperatures below – $100\,^{\circ}$ C are required to freeze the H_2O+CO_2 fluid mixture into crystalline solid CO_2 and clathrate. The total amount of fluid obtained both by (i) Rietveld refinement of X-ray diffraction patterns (12 ± 3.5 wt%) and (ii) image analysis of the X-ray diffraction microtomographic Sect. (11-17 wt%) agree within uncertainties with mass spectrometry measurements (8.14 ± 0.05 wt%). This protocol, tested for H_2O-CO_2 mixtures, is likely suitable for all those fluid species that can be frozen using a cryostat. For instance, methane (freezing point – $182\,^{\circ}$ C), sulfur dioxide ($-72\,^{\circ}$ C) and

hydrogen sulfide (– 85.5 °C) could be suitable species of geological interest¹⁹. Our results open new perspectives when non-destructive analysis of experimental fluids is required.

Data availability

The authors declare that all data are available within the manuscript.

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References

- Houghton, R. A. The contemporary carbon cycle in *Treatise on Geochemistry* Vol. 8 (ed. Schlesinger, W. H.) 473–513 (Elsevier, 2004).
- 2. Plank, T. & Manning, C. E. Subducting carbon. Nature 574, 343-352 (2019).
- 3. Tumiati, S. et al. Silicate dissolution boosts the CO₂ concentrations in subduction fluids. Nat. Commun. 8, 1-11 (2017).
- 4. Tiraboschi, C., Miozzi, F. & Tumiati, S. Carbon-saturated COH fluids in the upper mantle: a review of high-pressure and high-temperature ex situ experiments. *Eur. J. Mineral.* **34**, 59–75 (2022).
- 5. Tumiati, S. *et al.* Subducted organic matter buffered by marine carbonate rules the carbon isotopic signature of arc emissions. *Nat. Commun.* **13**, 1–10 (2022).
- 6. Tumiati, S. *et al.* Dissolution susceptibility of glass-like carbon versus crystalline graphite in high-pressure aqueous fluids and implications for the behavior of organic matter in subduction zones. *Geochim. Cosmochim. Acta* **273**, 383–402 (2020).
- 7. Tiraboschi, C., Tumiati, S., Recchia, S., Miozzi, F. & Poli, S. Quantitative analysis of COH fluids synthesized at HP–HT conditions: an optimized methodology to measure volatiles in experimental capsules. *Geofluids* 16, 841–855 (2016).
- Miozzi, F. & Tumiati, S. Aqueous concentration of CO₂ in carbon-saturated fluids as a highly sensitive oxybarometer. Geochem. Persp. Let. 16, 30–34 (2020).
- 9. Eugster, H. P. & Skippen, G. B. Igneous and metamorphic reactions involving gas equilibria. Res. Geochem. 2, 492-520 (1967).
- 10. Longhi, J. Phase equilibria in the system CO₂–H₂O I: New equilibrium relations at low temperatures. *Geochim. Cosmochim. Acta* **69**, 529–539 (2005).
- 11. Luth, R. W. Natural versus experimental control of oxidation state: effects on the composition and speciation of C-O-H fluids. *Am. Mineral.* **74**, 50–57 (1989).
- Bose, K. & Ganguly, J. Quartz-coesite transition revisited; reversed experimental determination at 500–1200 degrees C and retrieved thermochemical properties. Am. Mineral. 80, 231–238 (1995).
- 13. Larson, A.C. & Von Dreele, R.B. General structure analysis system (GSAS). Los Alamos National Laboratory Report LAUR, 86–748 (2004).
- 14. Toby, B. H. EXPGUI, a graphical user interface for GSAS. J. Appl. Cryst. 34, 210-213 (2001).
- 15. Vaughan, G. B. et al. ID15A at the ESRF-a beamline for high speed operando X-ray diffraction, diffraction tomography and total scattering. *J. Synchrotron Radiat.* 27, 515–528 (2020).
- 16. Circone, S. et al. CO₂ hydrate: synthesis, composition, structure, dissociation behavior, and a comparison to structure I CH₄ hydrate. *J. Phys. Chem. B* **107**, 5529–5539 (2003).
- 17. Zevin, L. S. & Kimmel, G. Quantitative X-ray Diffractometry (Springer, 1995).
- 18. León-Reina, L. et al. Accuracy in Rietveld quantitative phase analysis: a comparative study of strictly monochromatic Mo and Cu radiations. J. Appl. Cryst. 49, 722–735 (2016).
- 19. Piccoli, F. et al. Subducting serpentinites release reduced, not oxidized, aqueous fluids. Sci. Rep. 9, 1-7 (2019).

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Author contributions

S.T. conceived this experimental study, wrote the manuscript, performed the Q.M.S. and the image analysis. M.M. and M.D.M. are responsible for X-ray data collection and reduction. A.A. assisted during the synchrotron measurements and was involved in the first stage of the research. L.T. collaborated to the final version of the manuscript.

Competing interests

The authors declare no competing interests.

Additional information

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