

Rising dust pollution across Europe in a changing climate

<https://doi.org/10.1038/s41586-026-10743-w>

Received: 29 April 2025

Accepted: 1 June 2026

Published online: 15 July 2026

Open access

 Check for updates

Petros N. Vasilakos^{1✉}, Abhishek Upadhyay¹, Manousos I. Manousakas^{1,2}, Andrés Alastuey³, James D. Allan^{4,5}, Céla A. Alves⁶, Benjamin Bergmans⁷, Benjamin T. Brem¹, Sonia Castillo⁸, Theodoros Christoudias⁹, Cristina Colombi¹⁰, Sébastien Conil¹¹, Katja Dzepina^{11,12}, Anja Eichler^{1,13}, Konstantinos Eleftheriadis², Olivier Favez¹⁴, Michael Flynn⁴, Kristina Glojek^{3,12}, Stuart K. Grange^{15,16}, David C. Green¹⁷, Christoph Hueglin¹⁵, Jean-Luc Jaffrezo¹⁸, Theo M. Jenk^{1,13}, Jianhui Jiang¹⁹, Ekaterina Krymova²⁰, Franco Lucarelli²¹, Petra Makorič²², Dario Massabò²³, Nikolaos Mihalopoulos^{9,24,25}, Griša Močnik¹², Robin L. Modini¹, Claudia Mohr¹, Attilio Naccarato²⁶, Petra Pokorná²⁷, Paolo Prati²³, Nicole Probst-Hensch^{28,29}, André S. H. Prévôt¹, Xavier Querol³, Cristina Reche³, Jesús D. de la Rosa³⁰, Mark M. Scerri³¹, Jean Sciare⁹, Michael Sigl^{1,13}, Anja H. Tremper¹⁶, Rita Traversi³², Daniel Trejo Banos²⁰, Maria Tsagkaraki²⁴, Gaëlle Uzu¹⁷, Roberta Vecchi³³, Marta Via¹², Kees de Hoogh^{28,29}, Imad El-Haddad^{1✉} & Kaspar R. Daellenbach^{1✉}

Mineral desert dust is a major contributor to total atmospheric particulate matter¹. Desert dust outbreaks degrade air quality and can pose adverse health effects², including asthma exacerbation³ and increased mortality⁴. At some European locations, there has been a rise in the intensity and frequency of transported dust outbreaks from deserts in recent decades^{5–9}. However, it remains unclear whether this increase is consistent across Europe and whether desertification and aridity or shifts in atmospheric circulation are the main drivers behind this rise. Here we compile a database of daily dust metal concentrations from European sites, establishing robust elemental ratios for transported dust. Using this database, we develop a machine learning model to estimate daily PM₁₀ (particulate matter smaller than 10 µm) dust concentrations from 2012 to 2021, ranging from $2.09 \pm 1.05 \mu\text{g m}^{-3}$ across northern and central Europe to $5.28 \pm 2.65 \mu\text{g m}^{-3}$ across the south. In southern Europe, residents are exposed to transported dust events averaging $9.68 \pm 4.85 \mu\text{g m}^{-3}$, linked to a $0.67 \pm 0.02\%$ rise in daily mortality. Intensified dust intrusions over the past decade are linked to shifts in atmospheric circulation. Data from an Alpine ice core record shows a 110% increase in dust concentrations since pre-industrial times, mostly associated with North African desertification. As climate change accelerates land degradation and affects weather patterns, worsening dust pollution may pose increasing risks to public health and air quality goals.

Dust has a central role in the Earth and climate systems by interacting with solar and terrestrial infrared radiation¹⁰, serving as cloud condensation and ice nuclei¹¹, supplying iron and other nutrients to ecosystems¹²—particularly over oceans—and influencing atmospheric acidity¹³, which affects the partitioning of inorganic and organic vapours¹⁴. Desert dust outbreaks degrade air quality by raising PM₁₀ and PM_{2.5} (particulate matter smaller than 2.5 µm) levels^{15,16}, disrupt economic activities such as air traffic¹⁷ and pose adverse health effects², including asthma exacerbation³, stillbirths¹⁸, increased mortality⁴ and even the transport of pathogens¹⁹. Analyses of transported dust outbreaks from the Saharan and Middle East deserts, at specific European locations, reveal a rising intensity and frequency in recent decades^{5–9}. Regional and global modelling show that dust emissions and transport are tightly coupled to large-scale circulation patterns, including the North Atlantic Oscillation^{20–22} (NAO) and stratospheric intrusions²¹,

which would modulate the frequency of dust intrusions over Europe. Long-term reconstructions from Alpine ice cores²³ and recent assessments of dust-related health risks² further highlight the growing importance of dust under a changing climate, probably linked to increasing North African droughts. However, owing to the strong spatial and temporal variability in dust levels, it remains unclear whether this increase is consistent across Europe and whether desertification and aridity or shifts in atmospheric circulation are the main drivers behind this rise. Most existing studies rely on coarse-resolution dust transport models, satellite proxies or total PM₁₀ concentrations²⁴ to infer dust concentrations, approaches limited by uncertainties in emission parameterizations, spatial resolution or the lack of observational constraints. Consequently, there is no continent-wide, observation-driven quantification of dust levels, trends, drivers and impacts on air quality and health.

A list of affiliations appears at the end of the paper.

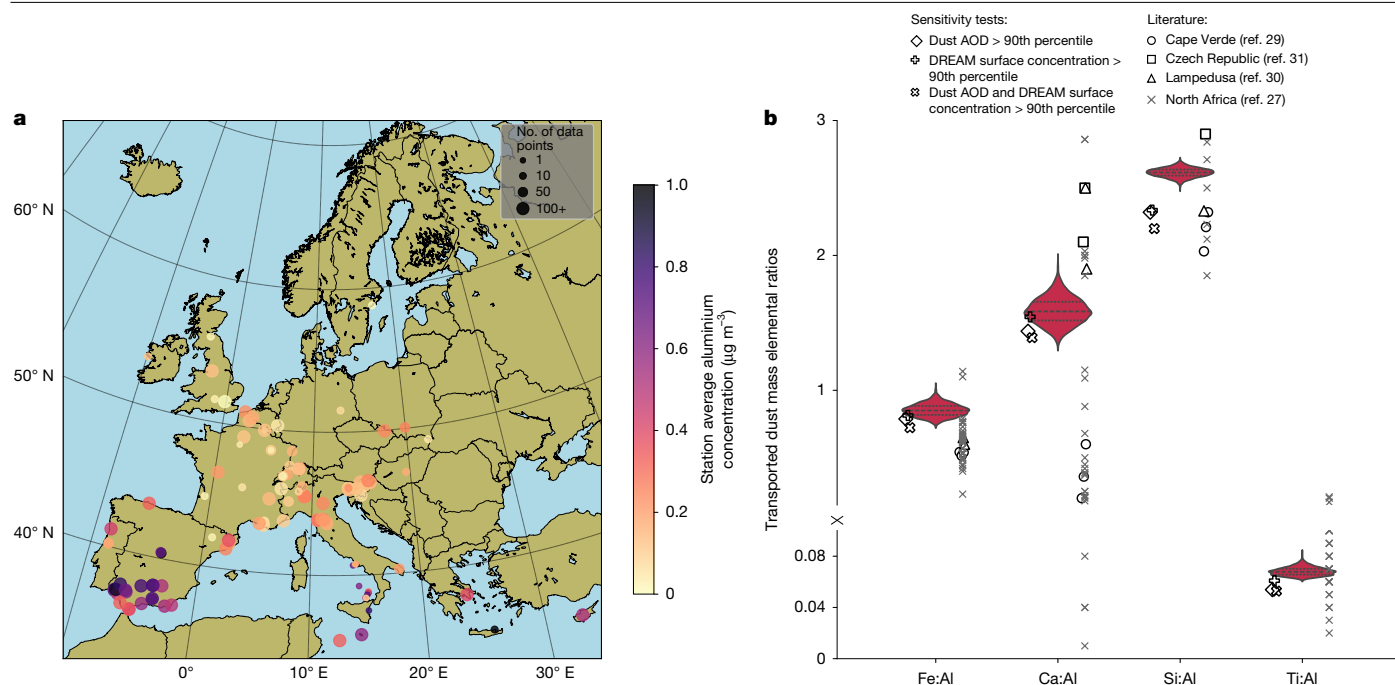


Fig. 1 | Dust observations and elemental ratios of transported dust in Europe.

a, Average aluminium concentrations for 103 locations in Europe. The size of the dot corresponds to the number of daily measurements and the colour corresponds to the average concentration of Al at this location. High Al concentrations show the large impact of transported dust on air quality in southern Europe. **b**, Violin plots of elemental ratios of transported dust, estimated by means of bootstrapped zero-intercept linear regressions applied

to daily data with increased transported dust concentrations (Al concentrations $>1 \mu\text{g m}^{-3}$ for Fe:Al and Ca:Al). The diamonds and outlined crosses and 'x's correspond to filtering to the 90th percentile of dust aerosol optical depth (AOD), DREAM dust surface concentration and both, respectively. Circles, squares and triangles represent literature values from Chiapello et al.²⁹, Loskot et al.³¹ and Marconi et al.³⁰, respectively, for dust arriving to Europe, whereas North Africa source estimates from Liu et al.²⁷ are denoted with the grey 'x's.

Here we address this gap by developing a data-driven, observation-constrained random forest (RF) model trained on the most extensive elemental dataset available in Europe. The model integrates several data products, including satellite-derived dust optical depth, land use information and a state-of-the-art physical dust model. The resulting European-wide daily dust PM_{10} concentrations for 2012–2021, combined with ice core observations, enable a comprehensive characterization of both the short-term variability and long-term drivers of dust, identifying the climatic modes controlling its transport and assessing its contribution to present air quality limit values and associated mortality from short-term exposure. We compile a uniquely comprehensive database of PM_{10} dust metal measurements (Al, Ti, Si, Ca and Fe) from 103 rural and urban sites across Europe, totalling about 18,500 daily measurements (Fig. 1a and Supplementary Table 1). We note that the spatial coverage of metal measurements is uneven across Europe, with limited data availability in northeastern Europe, the Balkans and Scandinavia—regions that are also climatologically less affected by Saharan dust. This data limitation underscores the need for further long-term supersite measurements in these areas, in line with present European air quality directives²⁴. Elements primarily associated with transported dust (Al, Si and Ti) show strong correlations throughout the dataset (Supplementary Fig. 1). By contrast, Ca and Fe, which are also influenced by local emissions such as brake wear, construction, soil erosion and road resuspension, show a larger variation in elemental ratios (Supplementary Fig. 1). Consistent with previous findings¹⁵ in Barcelona, Al and Ti serve as reliable tracers of desert dust, whereas Ca has substantial urban contributions. Although mean Al concentrations show no clear east–west differences, levels are clearly higher in the south than in the north, reflecting the susceptibility of these regions to Saharan dust intrusions in line with the literature²⁵ (Fig. 1a). Although dust is mostly present in the coarse mode, on the basis of the observed $\text{PM}_{2.5}:\text{PM}_{10}$ elemental ratios at locations for which both size

fractions are available, we estimate that 27.7% of dust is present as $\text{PM}_{2.5}$ (Supplementary Fig. 2).

Next we determine elemental ratios characteristic of transported dust by applying zero-intercept linear regression to a subset of data with high Al concentrations (here $>1 \mu\text{g m}^{-3}$; details in Methods section ‘Uncertainty of elemental ratios and dust estimate’)—indicative of increased transported dust contributions—and find Si:Al, Ti:Al, Ca:Al and Fe:Al ratios of 2.610 ± 0.033 , 0.068 ± 0.003 , 1.580 ± 0.099 and 0.850 ± 0.052 , respectively (Fig. 1b). Although standard linear regression would provide a better fit of the data, it would not be physically consistent, because, for pure dust, if one species was zero, then all would have to be zero and a non-zero-intercept regression violates that principle. These elemental ratios align well with reported values for dust transported to Europe^{25–31} and North Africa source estimates²⁷, acknowledging variability in elemental composition depending on the emission area, particularly for Ca. Despite this variability, given the extensive spatial and temporal coverage of our dataset across urban and rural environments, these ratios provide the most comprehensive observational characterization of the average elemental composition of transported dust in Europe so far.

Typically transported dust events, originating mainly from the Saharan and Middle East deserts, are predicted using three-dimensional models such as SKIRON, the Multiscale Online Nonhydrostatic Atmosphere Chemistry model (MONARCH), CHIMERE³² and the Barcelona Supercomputing Center Dust Regional Atmospheric Model (BSC-DREAM model, hereafter DREAM)³³. These models are essential for long-distance transport predictions as they simulate key atmospheric dust processes, including aeolian emission, transport and deposition. Satellite products, particularly dust optical depth at 550 nm (refs. 11,34), are also used to refine and validate model predictions. In fact, the European Commission’s working paper proposed a methodology combining modelling, remote sensing and particulate matter

concentration measurements to estimate daily dust contributions to particulate matter for compliance with limit values²⁴. Although effective at predicting extreme events, these models struggle with accurately resolving ground-level concentrations, often miss minor dust intrusions ($<10 \mu\text{g m}^{-3}$) and do not account for local sources of dust, such as wind-blown soil erosion, resuspension or agricultural activities³⁵. Recently, machine learning techniques have shown promise in predicting dust events, particularly in the Middle East^{36,37}. However, these models mostly rely on PM_{10} measurements, which include substantial contributions from anthropogenic and naturally formed particles³⁸, making them less accurate as dust-specific indicators.

Dust phenomenology

We developed a RF model for daily, ground-level dust concentrations ($10 \times 10 \text{ km}$), fusing the unique measurement database with reanalysis dust optical depth, a state-of-the-art physical dust model (DREAM³³), meteorological parameters (wind speed, temperature, wind direction and total precipitation) and land use data such as road coverage and population density (250 m-resolution land use data were used for each station) over continental Europe.

For quantifying dust, we opted for a single-tracer approach rather than a multitracer one, as most locations lack comprehensive tracer data, preventing consistent training of the RF model. Although Si would be an ideal tracer owing to its clear association with transported dust, Si measurements are notably less abundant than Al owing to common sampling methodologies using quartz filters. Given the strong correlation between Al and Si where both are available, we rely on Al as a transported dust proxy in the RF model (Supplementary Fig. 1). Specifically, we predict aluminium concentrations with the RF model and convert them to dust concentrations using the elemental ratios from the observational dataset ($\text{Ti}:\text{Al} = 0.068$, $\text{Si}:\text{Al} = 2.610$, $\text{Fe}:\text{Al} = 0.850$, $\text{Ca}:\text{Al} = 1.580$; Fig. 1b). This conversion assumes the elements being present as oxides, that is, Al_2O_3 , SiO_2 , CaO , Fe_2O_3 , FeO , TiO_2 and K_2O ($2.20\text{Al} + 2.49\text{Si} + 1.63\text{Ca} + 1.94\text{Ti} + 2.42\text{Fe}$)³⁹, which yields a dust estimate 13.5 times the Al concentration, similar to values reported in the literature²⁶. We validate the RF model by predicting daily Al levels at locations and years unseen by the model, using a rigorous leave-one-station-out (training the model on all but one station and testing it on the excluded station) and leave-one-year-out (training the model on all but one year and testing it on the excluded year) cross-validation. The RF model effectively captures daily dust concentrations in unseen training years (Supplementary Fig. 3b) and stations (Supplementary Fig. 3c), yet overestimates concentrations below $0.1 \mu\text{g m}^{-3}$ and underestimates the highest concentrations above $10 \mu\text{g m}^{-3}$ (Supplementary Fig. 3b,c), leading to an overall positive bias (fractional bias for leave-one-year-out: 30.1%; fractional bias for leave-one-station-out: 38.2%). When averaged annually across unseen locations, the model shows much better agreement with the measurements (fractional bias: 14.2%), making it well suited for long-term trend analysis (Supplementary Fig. 4). DREAM dust concentrations contribute the most to the predictions (RF weight 0.32), whereas the dust optical depth and meteorological variables also have an important effect (Supplementary Fig. 5). By contrast, land use variables have a small impact (RF combined weight < 0.18), consistent with most of Al arising from transported dust and not local sources^{40,41}, while at the same time indicating that, even for background dust, transported dust is its dominant component, even though it is a mixture of both local sources and desert dust from low-intensity events. The RF model accurately predicts multiyear dust concentrations and intrusions at the sites with the longest time series (Supplementary Fig. 6), without exhibiting a time-dependent bias (Supplementary Fig. 7). It outperforms DREAM by not only capturing high dust episodes but also moderate and low dust events (dust concentrations $< 20 \mu\text{g m}^{-3}$) (Supplementary Fig. 3). As a study case, the RF model captures the plume path and magnitude of the most severe and heavily

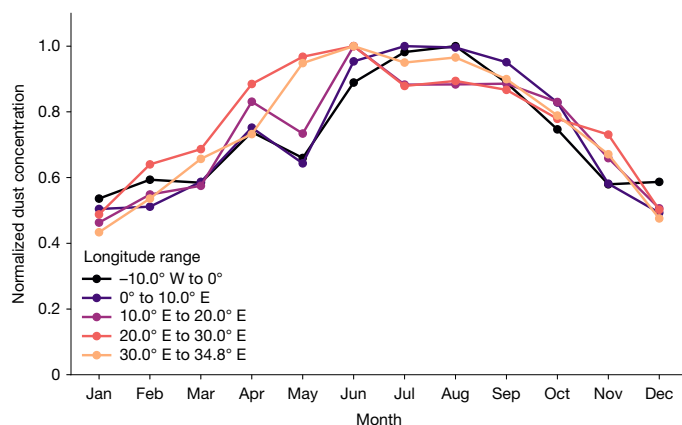


Fig. 2 | Dust seasonality in southern Europe. Monthly variation of normalized dust concentration across different longitudinal bands in southern Europe (land surface from 34° to 45°N) is presented. Normalized dust concentrations are calculated by dividing the mean monthly concentrations for the entire study decade (2012–2021) by the maximum mean monthly concentration within the respective longitudinal band.

studied Saharan dust events that hit the Balkans in April 2019, showing close agreement in plume path and ground-level concentrations ($30\text{--}80 \mu\text{g m}^{-3}$)⁴², supporting the suitability of the model for analysing dust frequency and intensity across Europe (Supplementary Fig. 8). By comparison, the physical DREAM model exhibits a strong negative bias at concentrations below $20 \mu\text{g m}^{-3}$ and a tenfold positive bias at concentrations above $50 \mu\text{g m}^{-3}$ (Supplementary Fig. 3a). Overall, the RF model outperforms DREAM in capturing spatiotemporal variations in dust loadings over Europe. The RF model shows high accuracy in regions with dense station clusters (for example, southern Spain), in which dust concentrations are highest. By contrast, its accuracy is expected to be lower in eastern Europe and Scandinavia owing to fewer measurements, although these areas are less affected by transported dust. Given the extensive dataset that powers the model, the lack of severe biases and the ability of the model to accurately capture intrusions both large and small, the RF model is uniquely suited for analysing dust in Europe.

Using the RF model, we generated a decadal European dust phenomenology for assessing transported dust, enabling an investigation into the trends in dust intrusion severity and frequency over the study period (2012–2021). For all of the analysis that follows, southern Europe is defined as all of the land surface of Europe below and including the latitude of Milan (45°N) and northern Europe as all land surface above that. In southern Europe, dust concentration shows a pronounced seasonal variation, with the lowest concentrations in winter (Fig. 2). However, seasonal patterns differ between the eastern and western Mediterranean areas (Fig. 2 and Supplementary Figs. 9 and 10). In the east, dust concentrations peak between March and May, a period during which levels in the west (for example, Spain) remain relatively low, and persist until October or November (Fig. 2), consistent with previous studies^{5,43}. Instead, the west experiences its dust peak in July–August, indicating a spatially dependent seasonality, with the east having a longer dust season than the west (Fig. 2 and Supplementary Figs. 9 and 10).

Yearly dust concentrations are predictably higher in southern Europe than in the north by a factor of 2.5 ($5.28 \pm 2.65 \mu\text{g m}^{-3}$ versus $2.09 \pm 1.05 \mu\text{g m}^{-3}$) (Fig. 3a). To assess changes in dust concentrations, we applied linear regression to the annual time series at each grid cell, using 100 bootstrap resamples. The reported change in dust concentration is the mean of the bootstrap slopes and cells in which zero lies within the interquartile range are masked in white (Fig. 3b, for regional means considered zero). For most of Europe, concentrations increase by $0.055 \pm 0.022 \mu\text{g m}^{-3} \text{ year}^{-1}$ between 2012 and 2021, with the largest increases over Italy and the Adriatic and Aegean seas

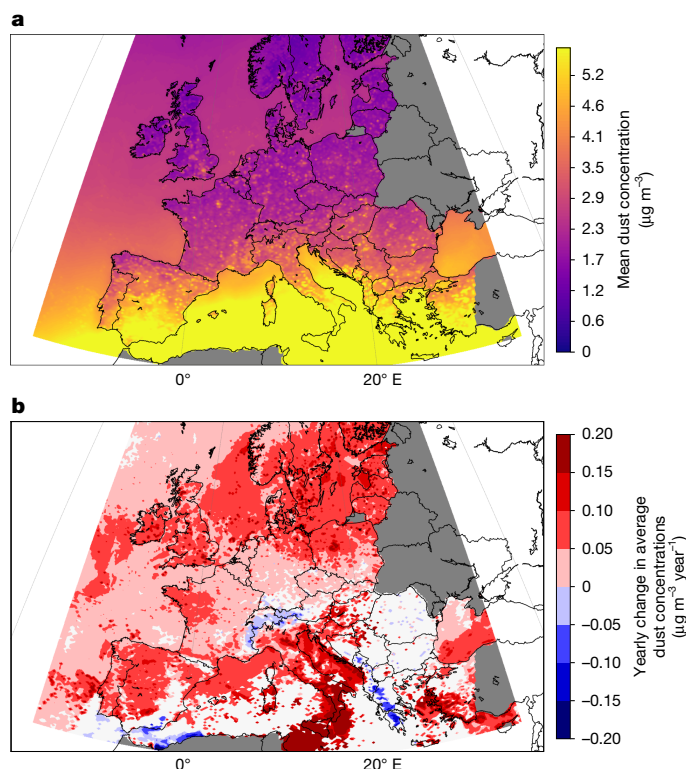


Fig. 3 | Dust concentrations and their changes in Europe during the period 2012–2021. a, b, Mean dust concentrations (a) and trend (b). Concentrations are reported in $\mu\text{g m}^{-3}$ and their corresponding trends in $\mu\text{g m}^{-3} \text{ year}^{-1}$. A trend is defined as the mean of a 100 times-bootstrapped linear regression for the variable of interest and shown on the map if there is no sign change between the 25th and 75th percentiles of the bootstrap results. Areas in which a sign change occurs are denoted with white (for regional means considered zero), whereas areas in which no land use data exist are denoted in grey.

at $0.074 \pm 0.030 \mu\text{g m}^{-3} \text{ year}^{-1}$ (Fig. 3b). Notably, in northern Europe, including Scandinavia, the magnitude of the increases is identical to that of the south ($0.055 \pm 0.022 \mu\text{g m}^{-3}$), suggesting that transported dust could become a growing concern for these regions in the future. A potential consideration for northern Europe is not only transported dust from the south but also Icelandic dust. Iceland is the most important dust source for the high-latitude areas of Europe and is also included in the model through DREAM⁴⁴, and therefore can contribute to the observed changes in concentration in northern Europe as well as to North African dust.

Daily dust levels exhibit pronounced variations, strongly affected by dust events. In accordance with present European Commission recommendations²⁴, we define transport events as deviations from a background (which includes both local and transported sources) varying over time for each location. For this, we determine for each day and each grid cell a concentration threshold (C_{th}) describing an upper concentration limit expected for background conditions (for details, see Methods section ‘Dust event detection’). This is a local and relative metric resulting in the number of detected transport events varying temporally and geographically. Southern Europe, including Spain, Italy and the eastern Mediterranean, experiences around 46 ± 7.82 dust intrusion days per year (Fig. 4a). However, compared with the average dust concentration (Fig. 3a), the north–south gradient is less pronounced (Fig. 4a). At higher altitudes, such as the Alps, Carpathians and Pyrenees, more dust intrusions are detected than in nearby lower-altitude areas, comparable with their number in the south. Conversely, dust concentrations during intrusions show a clear north–south gradient, similar to the average concentrations, with the highest levels

observed in the south, with an average of $11.00 \pm 5.51 \mu\text{g m}^{-3}$, gradually diminishing towards the north to an average value of $3.30 \pm 1.65 \mu\text{g m}^{-3}$ (Fig. 4c).

Changes in dust intrusions

The analyses of dust intrusion trends (Fig. 4b,d) reveal notable contrasts with average dust concentrations (Fig. 4a,c,e, analysis analogous to yearly means in Fig. 3). In northern Europe, including Scandinavia and the British Isles, background dust levels rise by $0.044 \pm 0.018 \mu\text{g m}^{-3} \text{ year}^{-1}$, nearly double the increase of $0.022 \pm 0.005 \mu\text{g m}^{-3} \text{ year}^{-1}$ during exceedances (Fig. 4f), suggesting that dust intrusions reaching these regions have not become more severe but rather the background dust levels have increased. It should be noted that, owing to the relative nature of the exceedance metric, the background also includes dust events of concentrations below the threshold and therefore it is expected that the observed increase is still attributed to transport. This interpretation is further supported by the fact that the increase during background conditions is observed over both land and sea—consistent with the spatial footprint expected for transported dust. Another potential driver is decreasing soil moisture owing to warming, albeit that the model-predicted AI seems insensitive to perturbations of soil parameters (Supplementary Fig. 11).

By contrast, southern Europe experiences a substantial rise in intrusion severity over the study period. In particular, dust concentrations during intrusion events increase by $1.40 \pm 0.33 \mu\text{g m}^{-3}$ over the decade in the Adriatic–Ionian region, the Aegean and much of the Balkans. The trend is even more pronounced in southern Italy, in which intrusion-related concentrations rise by $0.270 \pm 0.063 \mu\text{g m}^{-3} \text{ year}^{-1}$. These increases are mainly driven by more intense intrusion events, as changes in background concentrations remain comparatively small—approximately $0.032 \pm 0.013 \mu\text{g m}^{-3} \text{ year}^{-1}$ and $0.063 \pm 0.025 \mu\text{g m}^{-3} \text{ year}^{-1}$, respectively (Fig. 4f).

Meanwhile, the number of exceedance days (Fig. 4b) declines across Europe from 2012 to 2021, despite a decreasing threshold concentration for exceedance detection, suggesting a reduction in intrusion frequency. Eastern Europe, the Alpine regions and the eastern Mediterranean, including Greece, show fewer exceedance days but increased intrusion severity (Fig. 4b,d). In regions with the sharpest rise in exceedance concentrations, particularly southeastern Europe, the number of exceedance days generally remained unchanged. Overall, although the frequency of dust intrusions has not increased, their severity has greatly increased, along with a notable increase in background concentrations. Similar trends have been observed in the past with satellite data for the period 2000–2007, with increased intrusion severity and slightly reduced frequency⁴⁵.

Climatic and short-term drivers of dust

To assess whether the observed decadal trends reflect a sustained long-term increase in dust concentrations, we analysed Ca^{2+} concentration levels in Alpine ice cores from Colle Gnifetti (Monte Rosa, 4,450 m above sea level, $45^\circ 55' 46'' \text{ N}$; $07^\circ 52' 30'' \text{ E}$) for the period 1750–2020. The dataset combines three cores: one drilled in 2021 and two parallel cores recovered in 2003 (refs. 46,47). Because Al concentration data were only available until 1990, Ca^{2+} was used as a proxy for dust, an appropriate choice given the strong correlation between Ca^{2+} and Al during the overlapping period ($r = 0.87$, 1750–1990) and the absence of local sources at Colle Gnifetti. Although the ice core record reflects total dust deposition rather than atmospheric concentrations, it exhibits a consistent temporal pattern with the modelled dust concentrations over the nine overlapping years (2012–2020; Fig. 5a, inset), including a slight decline over the past decade—consistent with modelled trends across the Alps (Fig. 3b). Over the long term, Ca^{2+} concentrations preserved in the ice archive increased by

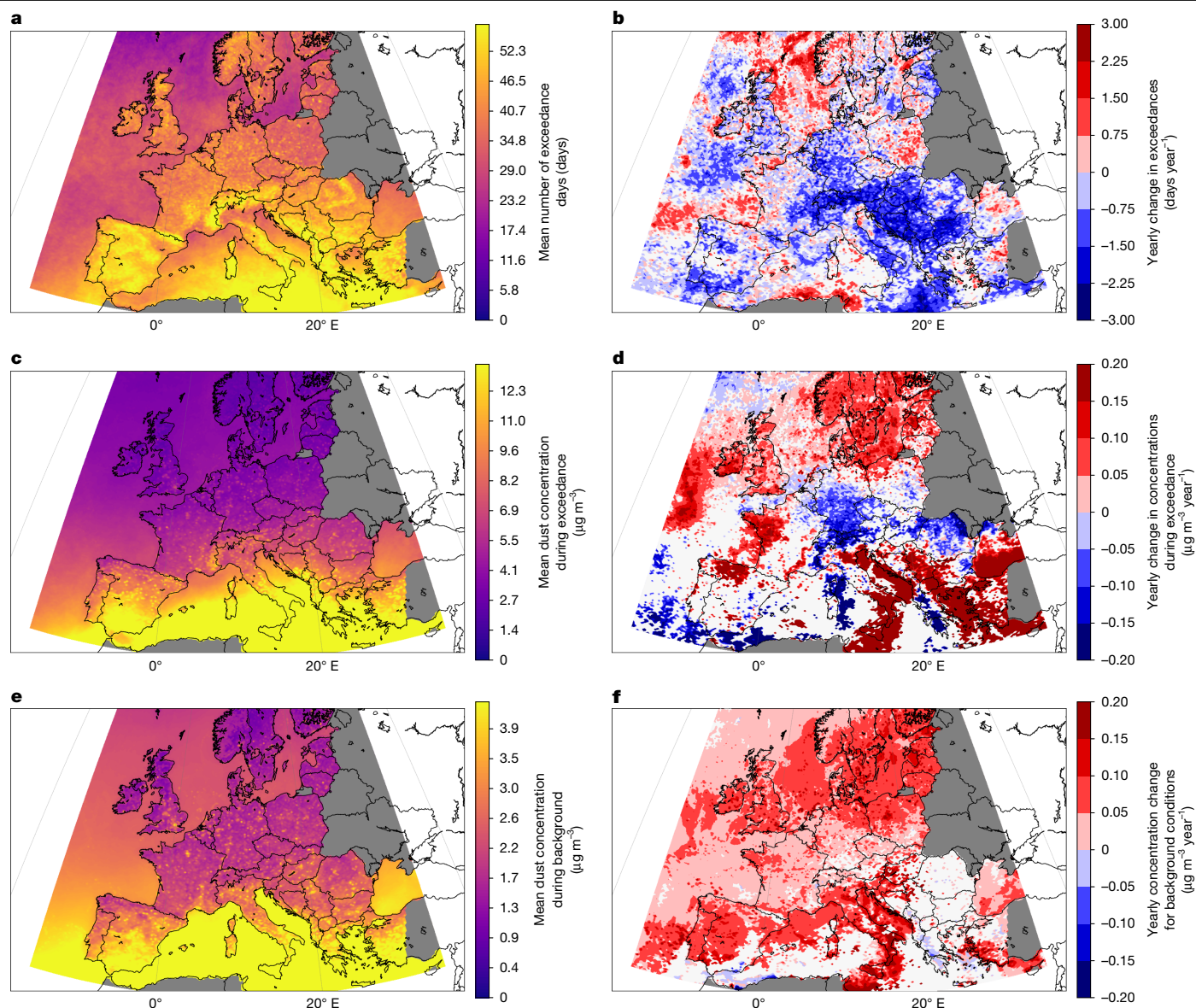


Fig. 4 | Dust episodes and their trends in Europe during the period 2012–2021. **a–f**, Mean number of days on which a dust exceedance occurs (**a**) and trend (**b**), mean dust exceedance concentrations (**c**) and trend (**d**) and mean dust concentrations during background (**e**) and trend (**f**). Concentrations are reported for PM₁₀ dust in $\mu\text{g m}^{-3}$ and their corresponding trends in $\mu\text{g m}^{-3} \text{ year}^{-1}$, whereas the number of exceedance days is reported as days,

with their corresponding trend reported as days year^{-1} . A trend is defined as the mean of a 100 times-bootstrapped linear regression for the variable of interest and shown on the map if there is no sign change between the 25th and 75th percentiles of the bootstrap results. Areas in which a sign change occurs are denoted with white (for regional means considered zero), whereas areas in which no land use data exist are denoted in grey.

approximately 110% between the pre-industrial period (1750–1850) and the past decade (2010–2020) (Fig. 5a). The trend derived from the three Colle Gnifetti cores is consistent with the multisite analysis in Kok et al.¹⁰, which reports a similar but less pronounced increase over the twentieth century. On interannual scales, Ca^{2+} concentrations in the ice core are correlated with the NAO (Spearman's rank correlation coefficient (R_s) = 0.48), consistent with the established influence of NAO-related circulation variability on dust transport^{9,20,23}. Also, the long-term trend in Ca^{2+} exhibits a strong negative correlation with the self-calibrating Palmer Drought Severity Index (PDSI)^{48,49}, which characterizes hydroclimatic variability over the primary dust source regions for the Swiss–Italian Alps, located in the northwestern Sahara (32.15°–35.66° N; –7.14°–11.27° E, following Coen et al.⁵⁰). Drier conditions—corresponding to lower PDSI values—are associated with higher Ca^{2+} levels in the ice core (R_s = –0.70). When restricting the analysis to Moroccan dust sources, the correlation strengthens further

(R_s = –0.78). These statistically robust associations indicate that long-term dust variability recorded in the Alpine archive is most strongly linked with persistent aridification trends in North Africa and with large-scale circulation variability, as represented by the NAO, which modulates precipitation and surface conditions over key dust source regions²¹ (Fig. 5a,b and Supplementary Fig. 12).

In Fig. 5d–f, we examine how circulation-oriented climate indices are statistically associated with dust inputs to Europe, alongside the dryness of the Sahara (Fig. 5c), considering a broad source region encompassing most of the Sahara (14.20°–30.17° N, –9.15°–32.52° E). The Saharan Oscillation Index (SaOI) represents the pressure difference between the Azores (37.79° N, –25.5° E) and Niamey (13.51° N, 2.10° E)⁵¹, whereas the Mediterranean Oscillation Index (MOI)⁵² corresponds to the normalized pressure difference between Algiers and Cairo. Both indices are closely related to the NAO^{51,52}. During 2012–2021, dust concentrations were strongly correlated with the SaOI and anticorrelated

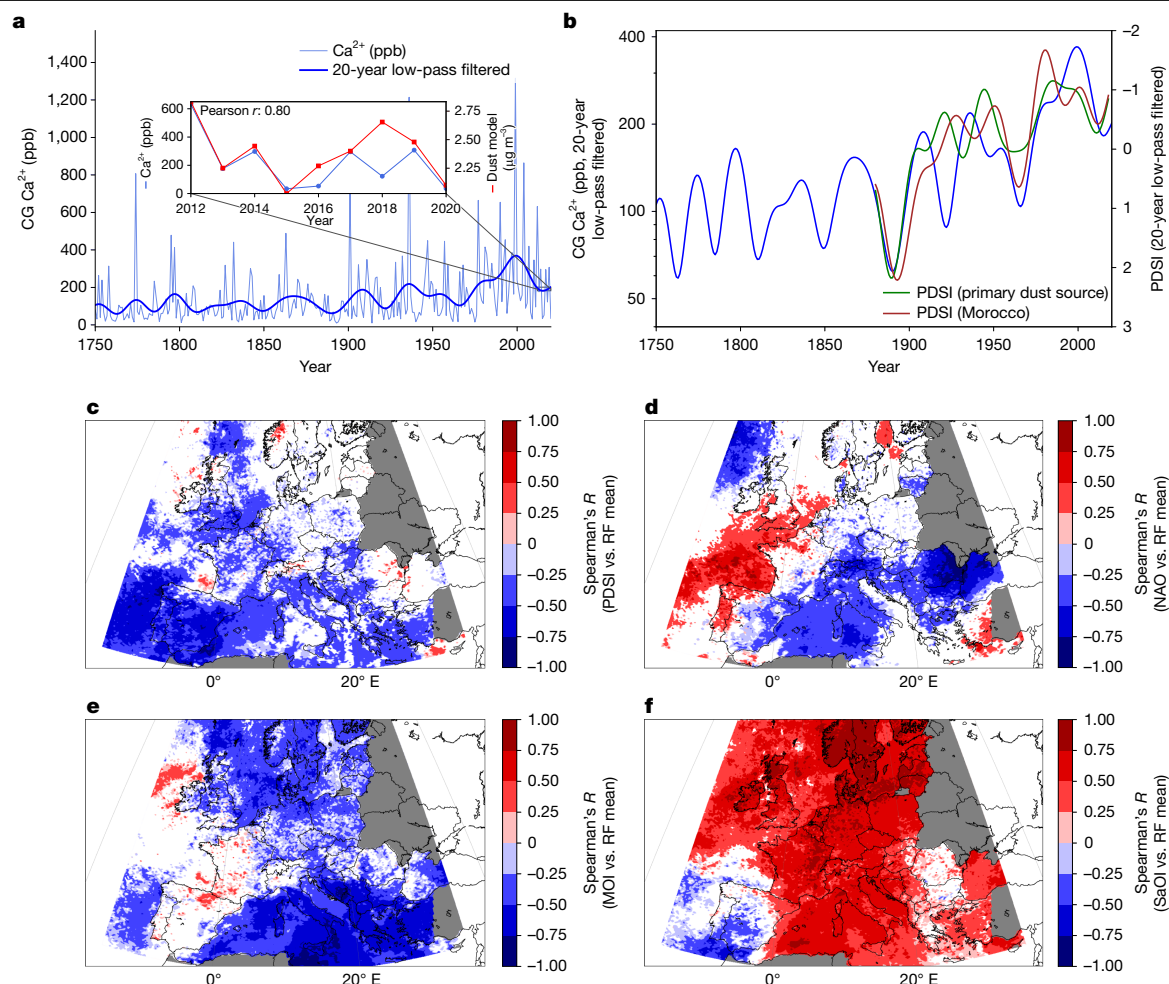


Fig. 5 | Long-term climatological trends and short-term variations in dust.

a,b, Time series of Colle Gnifetti (CG) ice core Ca^{2+} concentrations (annual, thin blue line; 20-year low-pass filtered, bold blue line) from 1750 to 2020 and comparison with model results for the period 2012–2020 (inset) (**a**), 20-year low-pass filtered Ca^{2+} concentrations (blue) and the PDSI in the main source region of dust for the Swiss–Italian Alps in the northwestern Sahara (-7.14° – 11.27° E, 32.15° – 35.66° N, green) and Morocco (-10° – 0° E, 30° – 35° N, brown) (**b**), along with the spatiotemporal correlation between annual mean dust concentrations in the period 2012–2021 and PDSI in the extended Sahara area (14.2° – 30.17° N, -9.15° – 32.52° E) (**c**), NAO index (**d**), MOI (**e**) and SaOI (**f**).

with the MOI (Fig. 5e,f), except in western Europe, in which correlations were weaker and occasionally of opposite sign. This pattern may reflect the dual influence of the MOI: although low MOI values are associated with enhanced westward dust transport, they also coincide with higher precipitation over western Europe⁵², probably offsetting increases in surface dust concentrations through wet deposition. Increasing SaOI values were most strongly correlated with enhanced Saharan dust advection towards Europe, whereas declining MOI values were associated with reduced eastward transport and a relative strengthening of dust inflow to central and southern Europe. These combined circulation tendencies are most strongly linked to intensified southerly winds from North Africa, consistent with the spatial dust patterns shown in Fig. 4d,f.

Similarly, the NAO exhibits a robust statistical relationship with dust transport. Positive NAO phases during 2012–2021 were associated with stronger cyclonic activity and westward transport, corresponding to higher dust levels over the Atlantic, France and Spain⁴¹ (Fig. 5d), whereas stronger westerlies under these conditions coincided with reduced dust transport into central and eastern Europe. Over the Aegean Sea,

Results for the correlations are presented as the mean of a 100 times-bootstrapped linear regression for the variable of interest and RF results and shown on the map if there is no sign change between the 25th and 75th percentiles of the bootstrap results. Areas in which a sign change occurs are denoted with white (for regional means considered zero), whereas areas in which no land use data exist are denoted in grey. This analysis is designed to reveal statistically significant associations, not causal relationships; however, the robustness of these correlations provides valuable insights into the drivers of the short-term variability in dust over Europe. ppb, parts per billion.

increased dust concentrations during positive NAO phases are most probably linked to intensified easterly winds in the far eastern Mediterranean⁴³.

Overall, the observed dust variability across Europe is most strongly correlated with concurrent trends in North African desertification and regional circulation patterns, emphasizing the dual influence of source region aridity and large-scale atmospheric dynamics on dust transport towards Europe.

Implications

By 2021, our model shows that, in southern Europe, transported dust alone already accounted for 31% of World Health Organization (WHO)⁵³ annual mean PM_{10} guideline values ($15 \mu\text{g m}^{-3}$) and 25.8% (calculated on the basis of observed $\text{PM}_{2.5}:\text{PM}_{10}$ elemental ratios) of $\text{PM}_{2.5}$ guideline values ($5 \mu\text{g m}^{-3}$), substantially hindering these regions to comply with regulations. Residents in the region experience 46 ± 7.82 dust episodes annually, with dust levels averaging $9.68 \pm 4.85 \mu\text{g m}^{-3}$ during these events. We estimate that short-term exposure to such dust

concentrations would be associated with a $0.67\% \pm 0.02\%$ contribution to daily mortality and a $0.73\% \pm 0.04\%$ contribution to daily respiratory hospitalizations during dust events among patients older than 15 years, as in Stafoggia et al.⁴ (Methods; using global relative risk coefficients leads to similar results presented in Supplementary Table 6). The impact on younger patients is even more pronounced, with a contribution of $2.47\% \pm 0.07\%$ in daily respiratory hospitalizations during dust events, highlighting more severe impacts of dust on vulnerable populations⁴. Although the increase in dust concentrations observed during the 10-year study period was not sufficient to cause substantial changes in health outcomes, continued increases in dust levels could exacerbate these health risks. We demonstrate that dust deposited in the Alps has increased by 110% since pre-industrial times owing to desertification, whereas shifts in synoptic circulation patterns have intensified intrusions, raising the annual dust PM₁₀ concentrations by $2.2 \pm 0.51 \mu\text{g m}^{-3}$ in the south of Italy over the past decade. With climate change accelerating desertification and altering atmospheric circulation, dust pollution will increasingly threaten public health and hinder efforts to meet WHO and European Union air quality targets—a direct feedback of climate change on air quality.

Online content

Any methods, additional references, Nature Portfolio reporting summaries, source data, extended data, supplementary information, acknowledgements, peer review information; details of author contributions and competing interests; and statements of data and code availability are available at <https://doi.org/10.1038/s41586-026-10743-w>.

- Kok, J. F. et al. Contribution of the world's main dust source regions to the global cycle of desert dust. *Atmos. Chem. Phys.* **21**, 8169–8193 (2021).
- Li, T. T. et al. Sand and dust storms: a growing global health threat calls for international health studies to support policy action. *Lancet Planet. Health* **9**, e34–e40 (2025).
- Kanatani, K. T. et al. Desert dust exposure is associated with increased risk of asthma hospitalization in children. *Am. J. Respir. Crit. Care Med.* **182**, 1475–1481 (2010).
- Stafoggia, M. et al. Desert dust outbreaks in Southern Europe: contribution to daily PM₁₀ concentrations and short-term associations with mortality and hospital admissions. *Environ. Health Perspect.* **124**, 413–419 (2016).
- Cuevas-Agulló, E. et al. Sharp increase in Saharan dust intrusions over the western Euro-Mediterranean in February–March 2020–2022 and associated atmospheric circulation. *Atmos. Chem. Phys.* **24**, 4083–4104 (2024).
- Salvador, P., Pey, J., Pérez, N., Querol, X. & Artíñano, B. Increasing atmospheric dust transport towards the western Mediterranean over 1948–2020. *npj Clim. Atmos. Sci.* **5**, 34 (2022).
- Varga, G., Rostási, A., Meiramova, A., Dagsson-Waldhauserová, P. & Gresina, F. Increasing frequency and changing nature of Saharan dust storm events in the Carpathian Basin (2019–2023) – the new normal? *Hung. Geogr. Bull.* **72**, 319–337 (2023).
- Mousavi, S. V. et al. Future dust concentration over the Middle East and North Africa region under global warming and stratospheric aerosol intervention scenarios. *Atmos. Chem. Phys.* **23**, 10677–10695 (2023).
- Liu, G. Y., Li, J. & Ying, T. The shift of decadal trend in Middle East dust activities attributed to North Tropical Atlantic variability. *Sci. Bull.* **68**, 1439–1446 (2023).
- Kok, J. F. et al. Mineral dust aerosol impacts on global climate and climate change. *Nat. Rev. Earth Environ.* **4**, 71–86 (2023).
- Brunner, C. et al. The contribution of Saharan dust to the ice-nucleating particle concentrations at the High Altitude Station Jungfraujoch (3580 m a.s.l.), Switzerland. *Atmos. Chem. Phys.* **21**, 18029–18053 (2021).
- Wyatt, N. J. et al. Phytoplankton responses to dust addition in the Fe–Mn co-limited eastern Pacific sub-Antarctic differ by source region. *Proc. Natl Acad. Sci. USA* **120**, e222011120 (2023).
- Vasilakos, P., Russell, A., Weber, R. & Nenes, A. Understanding nitrate formation in a world with less sulfate. *Atmos. Chem. Phys.* **18**, 12765–12775 (2018).
- Vasilakos, P., Hu, Y. T., Russell, A. & Nenes, A. Determining the role of acidity, fate and formation of IEPOX-derived SOA in CMAQ. *Atmosphere* **12**, 707 (2021).
- Querol, X. et al. PM₁₀ and PM_{2.5} source apportionment in the Barcelona Metropolitan area, Catalonia, Spain. *Atmos. Environ.* **35**, 6407–6419 (2001).
- Querol, X. et al. Monitoring the impact of desert dust outbreaks for air quality for health studies. *Environ. Int.* **130**, 104867 (2019).
- Baddock, M. C., Strong, C. L., Murray, P. S. & McTainsh, G. H. Aeolian dust as a transport hazard. *Atmos. Environ.* **71**, 7–14 (2013).
- Li, P. F. et al. Stillbirths associated with particle pollution are disproportionately contributed by sand dust: findings from 52 low- and middle-income countries. *Environ. Sci. Technol.* **58**, 15971–15983 (2024).
- Vergadi, E., Rouva, G., Angeli, M. & Galanakis, E. Infectious diseases associated with desert dust outbreaks: a systematic review. *Int. J. Environ. Res. Public Health* **19**, 6907 (2022).
- Shao, Y. P., Klose, M. & Wyrwoll, K. H. Recent global dust trend and connections to climate forcing. *J. Geophys. Res. Atmos.* **118**, 11,107–11,118 (2013).
- Dai, Y. et al. Stratospheric impacts on dust transport and air pollution in West Africa and the Eastern Mediterranean. *Nat. Commun.* **13**, 7744 (2022).
- Moulin, C., Lambert, C. E., Dulac, F. & Dayan, U. Control of atmospheric export of dust from North Africa by the North Atlantic Oscillation. *Nature* **387**, 691–694 (1997).
- Clifford, H. M. et al. A 2000 year Saharan dust event proxy record from an ice core in the European Alps. *J. Geophys. Res. Atmos.* **124**, 12882–12900 (2019).
- European Commission. Commission Staff Working Paper Establishing Guidelines for Demonstration and Subtraction of Exceedances Attributable to Natural Sources under the Directive 2008/50/EC on Ambient Air Quality and Cleaner Air for Europe Commission Staff Working Paper 6771/11 <https://data.consilium.europa.eu/doc/document/ST%206771%202011%20INIT/EN/pdf> (European Commission, 2011).
- Alastuey, A. et al. Geochemistry of PM₁₀ over Europe during the EMEP intensive measurement periods in summer 2012 and winter 2013. *Atmos. Chem. Phys.* **16**, 6107–6129 (2016).
- Borai, M. et al. Elemental ratios as tracers of the sources of mineral dust in north-eastern Sahara. *Int. J. Environ. Sci. Technol.* **21**, 1875–1888 (2024).
- Liu, X., Turner, J. R., Hand, J. L., Schichtel, B. A. & Martin, R. A global-scale mineral dust equation. *J. Geophys. Res. Atmos.* **127**, e2022JD036937 (2022).
- Nava, S. et al. Saharan dust impact in central Italy: an overview on three years elemental data records. *Atmos. Environ.* **60**, 444–452 (2012).
- Chiapello, I. et al. Origins of African dust transported over the northeastern tropical Atlantic. *J. Geophys. Res. Atmos.* **102**, 13701–13709 (1997).
- Marconi, M. et al. Saharan dust aerosol over the central Mediterranean Sea: PM₁₀ chemical composition and concentration versus optical columnar measurements. *Atmos. Chem. Phys.* **14**, 2039–2054 (2014).
- Loskot, J. et al. Impact of Saharan dust on particulate matter characteristics in an urban and a natural locality in Central Europe. *Sci. Rep.* **14**, 32002 (2024).
- Vautard, R., Bessagnet, B., Chin, M. & Menut, L. On the contribution of natural Aeolian sources to particulate matter concentrations in Europe: testing hypotheses with a modelling approach. *Atmos. Environ.* **39**, 3291–3303 (2005).
- Basart, S., Pérez, C., Nickovic, S., Cuevas, E. & Baldasano, J. M. Development and evaluation of the BSC-DREAM8b dust regional model over Northern Africa, the Mediterranean and the Middle East. *Tellus B* **64**, 18539 (2012).
- Lemmouchi, F. et al. Machine learning-based improvement of aerosol optical depth from CHIMERE simulations using MODIS satellite observations. *Remote Sens.* **15**, 1510 (2023).
- Leung, D. M. et al. A new process-based and scale-aware desert dust emission scheme for global climate models – part II: evaluation in the Community Earth System Model version 2 (CESM2). *Atmos. Chem. Phys.* **24**, 2287–2318 (2024).
- Ebrahimi-Khusfi, Z., Taghizadeh-Mehrjardi, R. & Mirakbari, M. Evaluation of machine learning models for predicting the temporal variations of dust storm index in arid regions of Iran. *Atmos. Pollut. Res.* **12**, 134–147 (2021).
- Jafari, R., Amiri, M. & Jebali, A. Machine learning-driven scenario-based models for predicting desert dust sources in central plays of Iran. *Catena* **234**, 107618 (2024).
- Daellenbach, K. R. et al. Sources of particulate-matter air pollution and its oxidative potential in Europe. *Nature* **587**, 414–419 (2020).
- Chow, J. C., Lowenthal, D. H., Chen, L. W. A., Wang, X. L. & Watson, J. G. Mass reconstruction methods for PM_{2.5}: a review. *Air Qual. Atmos. Health* **8**, 243–263 (2015).
- Gini, M., Manoussakis, M., Karydas, A. G. & Eleftheriadis, K. Mass size distributions, composition and dose estimates of particulate matter in Saharan dust outbreaks. *Environ. Pollut.* **298**, 118768 (2022).
- Vasilatou, V. et al. Long term flux of Saharan dust to the Aegean sea around the Attica region, Greece. *Front. Mar. Sci.* **4**, 42 (2017).
- Peshev, Z. Y., Deleva, A. D., Vulkova, L. A. & Dreischuh, T. N. Lidar observations of massive Saharan dust intrusion above Sofia, Bulgaria, in April 2019. *J. Phys. Conf. Ser.* **1859**, 012031 (2021).
- Pey, J., Querol, X., Alastuey, A., Forastiere, F. & Stafoggia, M. African dust outbreaks over the Mediterranean Basin during 2001–2011: PM₁₀ concentrations, phenomenology and trends, and its relation with synoptic and mesoscale meteorology. *Atmos. Chem. Phys.* **13**, 1395–1410 (2013).
- Cvetkovic, B. et al. Fully dynamic high-resolution model for dispersion of Icelandic airborne mineral dust. *Atmosphere* **13**, 1345 (2022).
- Gkikas, A., Hatzianastassiou, N. & Mihalopoulos, N. Aerosol events in the broader Mediterranean basin based on 7-year (2000–2007) MODIS C005 data. *Ann. Geophys.* **27**, 3509–3522 (2009).
- Huber, C. J. et al. High-altitude glacier archives lost due to climate change-related melting. *Nat. Geosci.* **17**, 110–113 (2024).
- Sigl, M. et al. 19th century glacier retreat in the Alps preceded the emergence of industrial black carbon deposition on high-alpine glaciers. *Cryosphere* **12**, 3311–3331 (2018).
- Liu, Y. et al. A new physically based self-calibrating Palmer drought severity index and its performance evaluation. *Water Resour. Manag.* **29**, 4833–4847 (2015).
- Dai, A., Bui, S., Lenssen, N. & National Center for Atmospheric Research Staff. *The Climate Data Guide: Palmer Drought Severity Index (PDSI)* <https://climateatdataguide.ucar.edu/climate-data/palmer-drought-severity-index-pdsi> (National Science Foundation, 2023).
- Coen, M. C. et al. Saharan dust events at the Jungfraujoch: detection by wavelength dependence of the single scattering albedo and first climatology analysis. *Atmos. Chem. Phys.* **4**, 2465–2480 (2004).
- Khomsi, K., Najmi, H., Chelhaoui, Y. & Souhaili, Z. The contribution of large-scale atmospheric patterns to PM₁₀ pollution: the new Saharan Oscillation Index. *Aerosol Air Qual. Res.* **20**, 1038–1047 (2020).
- Palutikof, J. in *Mediterranean Climate. Regional Climate Studies* (ed. Bolle, H. J.) 125–132 (Springer, 2003).
- World Health Organization. *WHO Global Air Quality Guidelines: Particulate Matter (PM_{2.5} and PM₁₀), Ozone, Nitrogen Dioxide, Sulfur Dioxide and Carbon Monoxide* (WHO, 2021).

Publisher's note Springer Nature remains neutral with regard to jurisdictional claims in published maps and institutional affiliations.



Open Access This article is licensed under a Creative Commons Attribution 4.0 International License, which permits use, sharing, adaptation, distribution and reproduction in any medium or format, as long as you give appropriate credit to the original author(s) and the source, provide a link to the Creative Commons licence, and indicate if changes were made. The images or other third party material in this article are included in the article's Creative Commons licence, unless indicated otherwise in a credit line to the material. If material is not included in the article's Creative Commons licence and your intended use is not permitted by statutory regulation or exceeds the permitted use, you will need to obtain permission directly from the copyright holder. To view a copy of this licence, visit <http://creativecommons.org/licenses/by/4.0/>.

© The Author(s) 2026

¹PSI Center for Energy and Environmental Sciences, Paul Scherrer Institute, Villigen, Switzerland. ²Environmental Radioactivity & Aerosol Technology for Atmospheric and Climate Impact Lab, INRaSTES, National Centre for Scientific Research "Demokritos", Athens, Greece. ³Institute of Environmental Assessment and Water Research (IDAEA-CSIC), Barcelona, Spain. ⁴Department of Earth and Environmental Sciences, The University of Manchester, Manchester, UK. ⁵National Centre for Atmospheric Science, The University of Manchester, Manchester, UK. ⁶Department of Environment and Planning, CESAM - Centre for Environmental and Marine Studies, University of Aveiro, Aveiro, Portugal. ⁷Institut Scientifique de Service Public - ISSeP, Liège, Belgium. ⁸Andalusian Institute for Earth System Research, IISTA-CEAMA, University of Granada, Granada, Spain. ⁹Climate and Atmosphere Research Centre (CARE-C), The Cyprus Institute, Nicosia, Cyprus. ¹⁰Agenzia Regionale per la Protezione

dell'Ambiente Lombardia (ARPA Lombardia), Milan, Italy. ¹¹ANDRA DISTEC/EES Observatoire Pérenne de l'Environnement, Bure, France. ¹²Centre for Atmospheric Research (CRA), University of Nova Gorica, Ajdovščina, Slovenia. ¹³Oeschger Centre for Climate Change Research (OCCR), University of Bern, Bern, Switzerland. ¹⁴Institut National de l'Environnement Industriel et des Risques (INERIS), Verneuil-en-Halatte, France. ¹⁵Laboratory for Air Pollution and Environmental Technology, Swiss Federal Laboratories for Materials Science and Technology (EMPA), Dübendorf, Switzerland. ¹⁶Climate and Environmental Physics, Physics Institute, University of Bern, Bern, Switzerland. ¹⁷Environmental Research Group, School of Public Health, Imperial College London, London, UK. ¹⁸University Grenoble Alpes, CNRS, IRD, INP-G, INRAE, IGE (UMR 5001), Grenoble, France. ¹⁹Global Institute for Urban and Regional Sustainability, School of Ecological and Environmental Sciences, East China Normal University, Shanghai, China. ²⁰Swiss Data Science Center, EPFL and ETH Zurich, Zurich, Switzerland. ²¹Department of Physics and Astronomy, University of Florence and INFN, Florence, Italy. ²²Laboratory for Environmental and Life Sciences, University of Nova Gorica, Nova Gorica, Slovenia. ²³Dipartimento di Fisica, Università di Genova, Genoa, Italy. ²⁴Environmental Chemical Processes Laboratory, Department of Chemistry, University of Crete, Heraklion, Greece. ²⁵Institute for Environmental Research and Sustainable Development, National Observatory of Athens, Athens, Greece. ²⁶Department of Chemistry and Chemical Technologies, University of Calabria, Rende, Italy. ²⁷Institute of Chemical Process Fundamentals, Czech Academy of Sciences, Prague, Czech Republic. ²⁸Swiss Tropical and Public Health Institute, Allschwil, Switzerland. ²⁹University of Basel, Basel, Switzerland. ³⁰CIQSO-Center for Research in Sustainable Chemistry, Associate Unit CSIC-University of Huelva "Atmospheric Pollution", Huelva, Spain. ³¹Institute of Earth Systems, University of Malta, Msida, Malta. ³²Department of Chemistry "Ugo Schiff", University of Florence, Florence, Italy. ³³Department of Physics, Università degli Studi di Milano, Milan, Italy. ^{E2}e-mail: petros.vasilakos@psi.ch; imad.el-haddad@psi.ch; kaspar.daellenbach@psi.ch

Methods

Dataset and database construction

We compiled an extensive database of daily average measurements of PM₁₀ elements related to dust (Al, Ti, Si, Ca and Fe) from all over Europe, to estimate dust concentrations. The bulk of the data was compiled ad hoc for this study, whereas some were extracted from the EBAS database⁵⁴. The most frequently measured metal that is prevalent in mineral dust was aluminium; measurement data for other dust metals found purely in dust such as silicon and titanium were much scarcer. By contrast, local sources such as road dust as well as non-exhaust traffic and construction emissions affect calcium and iron, making them unsuitable for estimating transported dust in this study⁴⁰. Metal concentrations were measured with both offline filter-based techniques (inductively coupled plasma (ICP), particle-induced X-ray emission (PIXE) and offline X-ray fluorescence (XRF))⁵⁵ and online techniques (online XRF X-ray induced acoustic computed tomography (XACT))^{55,56}. Al measurements from these techniques are directly comparable for the purposes of trend analysis, as numerous inter-laboratory comparison studies have shown that these techniques yield highly consistent results for particulate elemental concentrations^{57–61}. Also, grouping model RF cross-validation results by training technique (XRF-based, ICP-based or PIXE; Supplementary Fig. 13) does not indicate any inherent biases between the techniques. Great care was taken in curating the database by flagging data below detection limits and working with data providers to ensure high data quality. After data were aggregated from all providers, various methods were used to ensure data quality. Low concentration values at the detection limit were discarded. To that end, concentration values below a detection cut-off of Al = 0.003 µg m⁻³ were removed. Also, values that were repeated more than three times in sequence or repeated more than five times (at a third decimal precision) were interpreted as data that were replaced by the detection limit. Further data with a time resolution >1 day were discarded, whereas hourly data were aggregated to daily averages.

Uncertainty of elemental ratios and dust estimate

We quantified the impact of uncertainty in the elemental ratios—defined here as one standard deviation from the bootstrap results shown in Fig. 1b (Si:Al 2.610 ± 0.033, Ti:Al 0.068 ± 0.003, Ca:Al 1.580 ± 0.099 and Fe:Al 0.890 ± 0.052)—on the dust estimates using uncertainty propagation based on the formula: 2.2Al + 2.49Si + 1.63Ca + 1.94Ti + 2.42Fe. This results in a total dust estimate of (13.5 ± 0.377) × Al. Notably, the largest contributions to the overall uncertainty originate from Fe and Ca, which are also influenced by substantial local sources. Zero-intercept fits were used for element–element plots (for example, Ca concentration versus Al concentration) to estimate dust phase mass-based elemental ratios under the assumption that—after blank correction and removal of non-dust contributions—both analytes originate from the same dust endmember. In that case, when the dust contribution vanishes (Al → 0), the co-emitted element should also vanish (Ca → 0), implying a physically meaningful intercept of zero. Also, we tested a linear model with an intercept using bootstrap analysis and found that zero lies within one standard deviation of the mean intercept (intercept ± $\sigma_{\text{intercept}}$), supporting the use of a zero-intercept model. Given this result, along with the physical consistency and practical advantages of a zero-intercept regression—such as enabling the use of average ratios to estimate other components—we opted for the zero-intercept model. A summary of the relevant statistics is provided in Supplementary Table 2.

Elemental ratios on days with low amounts of transported dust can be substantially affected by local sources. With a rather high aluminium concentration threshold of >1 µg m⁻³ for determining the transported dust elemental ratios, we aimed at minimizing the impact of local sources on the elemental ratios. Also, we performed a sensitivity analysis varying the threshold, showing that the elemental ratios vary with cut-offs above 0.5 µg m⁻³ for Ca:Al and 0.75 µg m⁻³ for Fe:Al within

one standard deviation of the value obtained with a cut-off of 1 µg m⁻³. Further, we investigated model-driven criteria for dust episodes instead of using a static aluminium concentration cut-off using the DREAM and optical depth values and only including data points that were higher than the 90th percentile in those (each alone or both at the same time), as a proxy for dust events. This approach resulted in similar values (Fig. 1b and Supplementary Fig. 14).

The uncertainty σ_{mean} in modelled yearly mean dust concentrations $\text{dust}_{\text{mean}}$ (including exceedance and background concentrations) is estimated by accounting for two main components: (1) the uncertainty in the derived dust:Al ratio ($\text{RE}_{\text{ratio}} = 0.377/13.5 = 2.7\%$) and (2) the relative root mean square error of the modelled annual mean dust concentrations ($\text{RRMSE}_{\text{annual mean}} = 0.45$; Supplementary Fig. 3a), as determined through a leave-one-out performance analysis against chemically derived dust concentrations from in situ measurements (equation (1)).

$$\sigma_{\text{mean}} = \sqrt{(\text{RRMSE}_{\text{annual mean}} \times \text{dust}_{\text{mean}})^2 + (\text{RE}_{\text{ratio}} \times \text{dust}_{\text{mean}})^2} \quad (1)$$

$$= \sqrt{(\text{RRMSE}_{\text{annual mean}})^2 + \text{RE}_{\text{ratio}}^2} \times \text{dust}_{\text{mean}}$$

The robustness of dust concentration trends over the 10-year study period was evaluated by bootstrapping the linear regression of annual dust concentration time series at each grid cell (100 bootstrap resamples). The reported trend t_{dust} represents the mean of the bootstrap-derived slopes, in which grid cells where the interquartile range includes zero are masked in white (zero trends are assumed for regional means in such cases). The overall uncertainty $\sigma_{\text{trend,conc}}$ in the trend (equation (2)) includes both uncertainty in the dust:Al ratio and the bootstrap-based variability in fitting the trend (RE_{boot} , domain mean relative uncertainty excluding grid cells with non-significant trends), whereas the precision in the yearly dust prediction is already accounted for by the latter (see equation (1)):

$$\sigma_{\text{trend,conc}} = \sqrt{(\text{RE}_{\text{ratio}} \times t_{\text{dust}})^2 + (\text{RE}_{\text{boot}} \times t_{\text{dust}})^2} \quad (2)$$

$$= \sqrt{(\text{RE}_{\text{ratio}})^2 + (\text{RE}_{\text{boot}})^2} \times t_{\text{dust}}$$

with $\sqrt{(\text{RE}_{\text{ratio}})^2 + (\text{RE}_{\text{boot}})^2}$ equal to 0.5, 0.23 and 0.4 for the trend in mean, exceedance and non-exceedance concentrations, respectively.

Analogously, we estimate the uncertainty σ_{ed} of the number of exceedance days and related trends. We use counting statistics to describe the uncertainty of the number of days ($\text{RE}_d = 0.17$ being the domain mean relative uncertainty based on $\sqrt{\text{days}/\text{days}}$) (equation (3)).

$$\sigma_{\text{ed}} = \text{RE}_d \times \text{days} \quad (3)$$

The domain mean relative uncertainty $\sigma_{\text{trend,day}}$ (excluding grid cells with non-significant trends) in the trend of exceedance days t_{days} is estimated by means of the bootstrap-based relative variability ($\text{RE}_{\text{db}} = 0.27$) in fitting the trend (equation (4)), which also accounts for the precision in the yearly prediction in exceedance days (see equation (3)).

$$\sigma_{\text{trend,day}} = \text{RE}_{\text{db}} \times t_{\text{days}} \quad (4)$$

Machine learning model

We created a dust predicting model using RF^{34,36,62–64}, which is an ensemble learning method that constructs several decision trees during training and outputs the mean prediction of all individual trees. Each tree in the RF is trained on a bootstrapped subset of the data and node splits occur on the basis of a randomly selected number of features, leading to a reduction in overfitting and variance. RF was selected because it provided a balance between bias and variance and owing to its robustness to noisy data and feature correlation, especially given that aluminium concentrations can span many orders of magnitude, from a few

nanograms to tens of micrograms. We use the sklearn Python package RandomForestRegressor, which is based on the Breiman implementation of RFs^{64,65}. The model hyperparameters including the number of trees (n_estimators), maximum depth (max_depth) and minimum samples per split (min_samples_split) were tuned and optimized to ensure the best possible validation (parameters used in this work are n_estimators=150, max_depth=None, criterion='friedman_mse', bootstrap=True, random_state=66). RF also permits the derivation of the importance of each individual feature, allowing for the physical interpretation of the results.

Model inputs

The most influential input that drives the machine learning model are the daily transported dust fields provided by DREAM. DREAM is an Eulerian model that resolves the equations for aeolian dust emission and dispersion. It includes all of the relevant processes for dust emission and transport, from convection to deposition. A detailed description of the model can be found in ref. 33. We used the surface dust concentration model variable (SCONC_dust; $\mu\text{g m}^{-3}$) at a model resolution of 0.33° by 0.33° . As well as DREAM, other model inputs included CERRA sub-daily regional reanalysis data for Europe, specifically surface level temperature, precipitation, wind speed and wind direction⁶⁶, along with satellite column integrated dust optical depth (550 nm) from CAMS global reanalysis (EAC4)⁶⁷ for the years 2012–2021 (Supplementary Table 3). Finally, coordination of information on the environment (CORINE) annual land-use data at 200 m resolution for the whole of Europe were also used as RF model inputs for the training dataset (that is, stations), including the fractions of 14 land use types⁶⁸ (for example, urban fraction, natural green, agriculture and barren land), population density⁶⁹, altitude⁷⁰ and road length and category⁷¹. Vegetation is already accounted for in the model through the natural green land use variable; however, seasonal vegetation changes were not explicitly included. A comprehensive list of variables can be found in Supplementary Table 4. To improve spatial coverage, all data were harmonized by cubic interpolation, to a 10×10 km-resolution grid extending from the northern part of Tunisia to the edge of Norway, similar to the grid used in ref. 38. For grid cells with measurements over the sea, all land-use categories are set to 0 and the water fraction is set to 1.

Although RF is inherently resilient to potential collinearities⁷², to evaluate potential redundancy between DREAM-modelled dust and meteorological predictors, we examined the contribution of each variable to the RF model performance using both Shapley Additive Explanations (SHAP) values and variable importance metrics. DREAM dust emerged as the single most influential predictor, accounting for roughly 30% of the explained variance in model outputs, whereas individual meteorological variables (for example, wind speed) contributed substantially less (<5% each). This dominance of DREAM dust in both metrics, combined with the distinct physical meaning of this variable—representing long-range dust emission and transport processes rather than local meteorology—demonstrates that it provides independent explanatory power rather than redundant information.

Model performance

The model was optimized by hyperparameter testing, with the optimal parameters being n_estimators=150, max_depth=None, criterion='friedman_mse', bootstrap=True and random_state=66. The criterion for picking the optimal model was the average Pearson r value for all of the sites from the leave-one-station-out validation, that is, the model configuration that had the highest r across all sites. It was chosen over other absolute metrics such as root mean square error owing to the measurements spanning many orders of magnitude (Supplementary Figs. 2 and 3), which would lead to sites in the south with larger average concentrations, overpowering sites in the north. It is noteworthy that all configurations attempted yielded very similar results, indicating that the method was insensitive to the hyperparameters.

The RF model was trained using observations from 2013–2021, whereas reconstructed dust concentrations cover the analysis period 2012–2021. The year 2012 therefore represents an out-of-sample prediction. Notably, 'year' was not included as a predictor variable, preventing the model from learning or imposing artificial temporal trends. Consequently, trends in reconstructed dust concentrations arise solely from variations in the meteorological and environmental predictors rather than from the model structure itself.

A leave-one-year-out validation, that is, removing an entire year from the training dataset and predicting it, was also conducted, yielding similar results to the leave-one-station-out validation, indicating that the model carries no inherent spatiotemporal biases (Supplementary Tables 5 and 7 and Supplementary Figs. 7 and 15). To further validate the model, and test its generalizability, we conducted an extra temporal validation for the year 2012, which was not included in the training of the model²⁵. Also, these sites, with the exception of Athens and Finokalia, were also not included in the training dataset, meaning that the spatial and temporal predictive abilities of the model were tested at the same time. The year was not used as a predictor variable, preventing the model from learning or imposing artificial temporal trends. Model performance for the unused year 2012 is very similar to that of the other years used in training (Supplementary Table 7 and Supplementary Fig. 15), reinforcing the robustness of the model.

Model training is completed within less than 5 min, whereas producing dust fields for 10 years and for the entire domain takes 53 min on 60 processors (two cores per processor).

Dust event detection

For the detection of dust transport events, we determine for each day and each grid cell a concentration threshold (C_{th}) describing an upper concentration limit expected for background conditions. This methodology is based on the working paper by the European Commission for event identification²⁴. When the observed dust concentration exceeds this threshold concentration, the day in question is classified as a transport event for this grid cell. This daily threshold concentration for each grid cell of the domain is based on the rolling median of the 30 preceding and 30 following days ($C_{Med,30\text{ prec.}}$, $C_{Med,30\text{ foll.}}$), along with their median absolute deviation ($MAD_{30\text{ prec.}}$, $MAD_{30\text{ foll.}}$). Specifically for an exceedance to occur, the dust concentration needs to be higher than a threshold value, defined as (equation (5)):

$$C_{th} = \frac{(C_{Med,30\text{ prec.}} + C_{Med,30\text{ foll.}})}{2} + 3 \frac{(MAD_{30\text{ prec.}} + MAD_{30\text{ foll.}})}{2}. \quad (5)$$

Health impacts from short-term exposure to dust

To estimate the increase in mortality owing to short-term exposure to dust, we first calculated the fractional increase in mortality resulting from dust, based on the increase in risk for all-cause mortality (denoted IR) owing to acute dust exposure provided in ref. 4, per $10 \mu\text{g m}^{-3}$ increase in dust exposure. The corresponding IRs were $0.65\% \pm 0.02\%$ for all-cause mortality with a 0–1 lag day, $0.70\% \pm 0.02\%$ for respiratory hospitalizations of the older than 15 years age group with 0–5 lag days and $2.47\% \pm 0.07\%$ for respiratory hospitalizations for the 0–14 years age group with 0–5 lag days. This increase is the population weighted to account for the differences in population density. Assuming an exponential dose–response function, the percent increase in mortality or hospitalizations for a given grid area then becomes (equation (6)):

$$\begin{aligned} &\% \text{ increase in mortality} \\ &= 100\% \times \sum_{i=1}^{\text{no. of points}} \frac{\text{population of cell } i}{\text{total population}} \\ &\quad \times \left(1 - \frac{1}{e^{IR \times \text{exceedance concentration}_i}} \right), \end{aligned} \quad (6)$$

in which no. of points is equal to the total number of grid cells in the area, population of cell i is the total population within the grid cell, IR is the increase in risk for all-cause mortality and exceedance concentration, i is the average exceedance concentration over the decade for that grid cell. For this calculation, the south is defined as all of the land surface of Europe below, and including, Milan (45° N), for which we calculate an increase of $0.67\% \pm 0.02\%$ in all-cause mortality for exceedance days, a $0.73\% \pm 0.04\%$ increase in daily respiratory hospitalizations among the older than 15 years age group and a $2.55\% \pm 0.07$ rise in daily respiratory hospitalizations among the 0–14 years age group. The mortality calculations were repeated using the IRs from ref. 73 and shown in Supplementary Table 6.

Data availability

The full datasets shown in the figures and tables are publicly available at <https://doi.org/10.5281/zenodo.19236528> (ref. 74). Source data are provided with this paper.

Code availability

The random forest model was built in Python v3.9.10 using scikit-learn (<https://scikit-learn.org/stable/>) and is publicly available at <https://doi.org/10.5281/zenodo.19236528> (ref. 74).

54. Torseth, K. et al. Introduction to the European Monitoring and Evaluation Programme (EMEP) and observed atmospheric composition change during 1972–2009. *Atmos. Chem. Phys.* **12**, 5447–5481 (2012).
55. Manousakas, M. et al. XRF characterization and source apportionment of PM10 samples collected in a coastal city. *X-Ray Spectrom.* **47**, 190–200 (2018).
56. Furger, M. et al. Elemental composition of ambient aerosols measured with high temporal resolution using an online XRF spectrometer. *Atmos. Meas. Tech.* **10**, 2061–2076 (2017).
57. Gini, M. et al. Inter-laboratory comparison of ED-XRF/PIXE analytical techniques in the elemental analysis of filter-deposited multi-elemental certified reference materials representative of ambient particulate matter. *Sci. Total Environ.* **780**, 146449 (2021).
58. Saitoh, K., Sera, K., Gotoh, T. & Nakamura, M. Comparison of elemental quantity by PIXE and ICP-MS and/or ICP-AES for NIST standards. *Nucl. Instrum. Methods Phys. Res. B* **189**, 86–93 (2002).
59. Lucarelli, F. et al. Is PIXE still a useful technique for the analysis of atmospheric aerosols? The LABEC experience. *X-Ray Spectrom.* **40**, 162–167 (2011).
60. Chiari, M. et al. Comparison of PIXE and XRF analysis of airborne particulate matter samples collected on Teflon and quartz fibre filters. *Nucl. Instrum. Methods Phys. Res. B* **417**, 128–132 (2018).
61. Cadeo, L. et al. Intercomparison of online and offline XRF spectrometers for determining the PM₁₀ elemental composition of ambient aerosol. *Atmos. Meas. Tech.* **18**, 6435–6448 (2025).
62. Aryal, Y. Evaluation of machine-learning models for predicting aeolian dust: a case study over the Southwestern USA. *Climate* **10**, 78 (2022).
63. Wang, W., Samat, A., Abuduwaili, J., De Maeyer, P. & Van de Voorde, T. Machine learning-based prediction of sand and dust storm sources in arid Central Asia. *Int. J. Digit. Earth* **16**, 1530–1550 (2023).
64. Breiman, L. Random forests. *Mach. Learn.* **45**, 5–32 (2001).
65. Breiman, L. Arcing classifiers. *Ann. Stat.* **26**, 801–824 (1998).
66. Schimanke S. et al. *CERRA Sub-Daily Regional Reanalysis Data for Europe on Model Levels from 1984 to Present* (Climate Data Store, accessed 12 September 2024); <https://doi.org/10.24381/cds.622a565a>.
67. Inness, A. et al. The CAMS reanalysis of atmospheric composition. *Atmos. Chem. Phys.* **19**, 3515–3556 (2019).
68. Copernicus Land Monitoring Service. *CORINE Land Cover Dataset* (Copernicus, accessed 12 September 2024); <https://land.copernicus.eu/pan-european/corine-land-cover>
69. Eurostat. *Database - Population and Demography* (Eurostat, accessed 31 October 2025); <https://ec.europa.eu/eurostat/web/population-demography/demography-population-stock-balance/database>.
70. CGIAR-CSI. *STRM 90m DEM Digital Elevation Database* (Consortium for Spatial Information, accessed 29 May 2025); <http://srtm.csi.cgiar.org/>.
71. OpenStreetMap <https://www.openstreetmap.org/> (OpenStreetMap Foundation, 2026).
72. Tomaschek, F., Hendrix, P. & Baayen, R. H. Strategies for addressing collinearity in multivariate linguistic data. *J. Phon.* **71**, 249–267 (2018).
73. Pouri, N., Karimi, B., Kolivand, A. & Mirhoseini, S. H. Ambient dust pollution with all-cause, cardiovascular and respiratory mortality: a systematic review and meta-analysis. *Sci. Total Environ.* **912**, 168945 (2024).
74. Vasilakos, P. N., El Haddad, I., & Dällenbach, K. R., Code and data for 'Rising dust pollution across Europe in a changing climate'. *Zenodo* <https://doi.org/10.5281/zenodo.19236528> (2026).

Acknowledgements We would like to express our deep thanks to many people in the AASQA France for the sampling of all of these samples and to people in several laboratories in France, including IMT NE (under the coordination of L. Alleman) and IGE and the Air O Sol analytical plateau for the ICP-MSMS analysis of these samples by S. Darfeuill. We also acknowledge K. Styszko for filter collection during a previous project.

Author contributions P.N.V., I.E.-H. and K.R.D. designed and conceptualized the study. M.I.M., A.A., C.A.A., B.B., B.T.B., S. Castillo, C.C., S. Conil, K.D., A.E., K.E., O.F., M.F., K.G., S.K.G., D.C.G., C.H., J.-L.J., T.M.J., J.J., E.K., F.L., P.M., D.M., N.M., G.M., R.L.M., A.N., P. Pokorná, P. Prati, N.P.-H., A.S.H.P., X.Q., C.R., J.D.d.I.R., M.M.S., J.S., M.S., A.H.T., R.T., D.T.B., M.T., G.U., R.V., K.R.D., I.E.-H., C.M., J.D.A., K.d.H., M.V. and T.C. provided measured dust elemental data. P.N.V. and A.U. conducted the database construction. P.N.V., A.U., E.K. and D.T.B. performed the model development and modelling. P.N.V., A.U., I.E.-H., K.R.D., M.I.M., A.E., T.M.J. and M.V. conducted the data analysis. P.N.V., I.E.-H., K.R.D. and M.I.M. interpreted the results and wrote the manuscript. All authors reviewed the manuscript and provided feedback.

Funding The publication was made possible with funding from the ARMOUR project provided by the Swiss Federal Office for the Environment (BAFU); the SCENE project; the Swiss Data Science Center (SDSC) collaborative projects grant (C20-08); the National Key Research and Development Program of China (2023YFC3710400) and the National Natural Science Foundation of China (42207122); CESAM by FCT (UID/50017/2025 (<https://doi.org/10.54499/UID/50017/2025>) and LA/P/0094/2020 (<https://doi.org/10.54499/LA/P/0094/2020>)); the PATOS project (Particolato Atmosferico in Toscana) financed by Tuscany Region; the KALOS project financed by Calenzano (Florence) municipality. Samples in France were collected within many research and air quality assessment programmes, including the programmes CARA (financed by the Ministry of Environment within the LCSQA), DECOMBIO, CAMERA, SOURCES and QAMECS (all financed by Ademe), QAMECS (financed by University Grenoble Alpes), OPE – Andra (financed by Andra) and support from Atmo AuRA, Atmo Sud, Atmo Grand Est, Atmo Haut de France and Atmo Normandie for the sampling and analyses. Samples in Granada were collected within many research projects at the University of Granada and the Spanish Ministry of Science and Innovation. The IDAEA-CSIC thanks the support from the Spanish Ministry of Environment (MITERD), from MICIU through the AIRPHONEMA project (PID2022-142160OB-I00/MCIN/AEI/10.13039/501100011033/FEDER EU), from Generalitat de Catalunya (Direcció General de Qualitat Ambiental i Canvi Climàtic and AGAUR, 2017 SGR41) and from the Madrid Council. We also acknowledge support to the University of Huelva, by grants PID2021-126986OB-I00 and PID2024-157355OB-I00, funded by MICIU/AEI/10.13039/501100011033, and the Environmental Agency of the Andalusian Government. The research leading to these results was supported by the Ministry of Education, Youth and Sports of the Czech Republic as the Large Research Infrastructure Support Project – ACTRIS Participation of the Czech Republic (ACTRIS-CZ, LM2023030). This research has been supported by the Marie Skłodowska-Curie COFUND Postdoctoral Programme grant agreement No.101081355-SMASH, the Slovenian Research and Innovation Agency (programme grants no. P1-0385, I-0033) and the Municipality of Kanal ob Soči. The UK measurements were supported by the Natural Environmental Research Council (NERC) through the Integrated Research Observation System for Clean Air (OSCA) project (Manchester grant ref. NE/T001984/1, London grant ref. NE/T001909/2), part of the Clean Air Strategic Priorities Fund. Open Access funding provided by Lib4RI – Library for the Research Institutes within the ETH Domain: Eawag, Empa, PSI & WSL.

Competing interests The authors declare no competing interests.

Additional information

Supplementary information The online version contains supplementary material available at <https://doi.org/10.1038/s41586-026-10743-w>.

Correspondence and requests for materials should be addressed to Petros N. Vasilakos, Imad El-Haddad or Kaspar R. Daellenbach.

Peer review information *Nature* thanks Yves Balkanski, Xiao-Hui Bi and Yang Liu for their contribution to the peer review of this work. Peer reviewer reports are available.

Reprints and permissions information is available at <http://www.nature.com/reprints>.